



1 **The Effects of Wind Farm Wakes on Freezing Sea Spray in the Mid-** 2 **Atlantic Offshore Wind Energy Areas**

3
4 David Rosencrans^{1,2}, Julie K. Lundquist^{1,2,3}, Mike Optis^{2,4}, and Nicola Bodini²

5 ¹Department of Atmospheric and Oceanic Sciences, University of Colorado, Boulder, 80303, USA

6 ²National Renewable Energy Laboratory, Golden, 80401, USA

7 ³Renewable and Sustainable Energy Institute, Boulder, 80303, USA

8 ⁴Veer Renewables, Courtenay, V9N 9B4, Canada

9 *Correspondence to:* David Rosencrans (David.Rosencrans@Colorado.edu)

10

11 **Abstract**

12 The U.S. is expanding its wind energy fleet offshore where winds tend to be strong and consistent. In the mid-
13 Atlantic, strong winds, which promote heat transfer and wind-generated sea spray, paired with cold temperatures can
14 cause ice on equipment when plentiful moisture is available. Near-surface icing is induced by a moisture flux from
15 sea spray, which poses a risk to vessels and crews. Ice accretion aloft occurs when liquid precipitation is present and
16 can reduce turbine blade performance and introduce extra load and fatigue on the turbine. Thus, it is crucial to
17 understand the icing hazard across the mid-Atlantic. We analyze Weather Research and Forecasting model
18 numerical weather prediction simulations at coarse temporal resolution over a 20-year period to assess freezing
19 events over the long-term record and at finer granularity over the 2019–2020 winter season to identify the post-
20 construction turbine impacts. Over the 2019–2020 winter season, results suggest that sea-spray–induced icing can
21 occur up to 66 hours per month at 10 m at higher latitudes. Freezing events during this season typically occur during
22 cold air outbreaks, which are the introduction of cold continental air over the warmer maritime surface and last a
23 total duration of 253 hours. Over the 20-year period, all cold air outbreak events coincide with freezing conditions,
24 although not all freezing events are cold enough to signify a cold air outbreak. Further, we assess the impacts of
25 wind plant installation on icing using the fine-scale simulation data set. Wakes from large wind plants reduce the
26 wind speed, which mitigates the chance for freezing. Conversely, the near-surface turbine-induced introduction of
27 cold air in frequent wintertime unstable conditions enhances the risk for freezing. Overall, the turbine–atmosphere
28 interaction causes a net mitigation of freezing hours within the wind plant areas, with a reduction up to 17 hours at
29 20 m in January 2020.

30

31 **1 Introduction**

32 The offshore wind energy industry is undergoing rapid growth to supply emissions-free energy to the electrical
33 grid. Across the mid-Atlantic outer continental shelf (OCS), the installed capacity could reach 30 GW by 2030
34 (White House, 2021). Capacity expansion into relatively cold offshore regions will subject turbines to harsher
35 wintertime conditions, which necessitate an understanding of the hazards that marine icing poses to offshore wind
36 turbines, service vessels, and crew safety.

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38 Ice accretion reduces the aerodynamic efficiency of the turbine blade, which hinders energy capture and annual
39 energy production (Battisti et al., 2006; Kraj and Bibeau, 2010; Wei et al., 2020). Ice can remain on the rotors even
40 after freezing conditions end, as slow natural processes such as ice shedding and melting extend the limitation to
41 energy yield (Gao and Hong, 2021). Some observations indicate that excessive icing can reduce torque enough that
42 blade rotation stops entirely, causing up to 80 % reduced power production for a single turbine (Gao and Hu, 2021).
43 Investigating a 2050 future capacity expansion scenario, Novacheck et al. (2021) found that onshore icing events
44 could reduce wind energy generation by 7 % and 10 % over two case-study events using 65 % variable renewable
45 energy penetration. For this scenario, Minnesota and Wisconsin would experience 75 % reduced wind energy
46 generation during a daylong case study period, causing reliance on ramped up local gas generation and interregional
47 transmission to meet the load. Faster winds in cold air outbreaks (CAO) enhance wind-energy supply during high-
48 load cold-weather events, although the combination of cold temperatures and slow wind speeds following frontal
49 passage pose severe challenges for utility grid planners. While turbine blade icing is well studied (Martini et al.,
50 2021; Contreras Montoya et al., 2022), icing near the turbine base is not. Some turbines have icing detection and
51 mitigation technology included at added cost, although current strategies need improvement (Madi et al., 2019).
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53 The leading causes for low-level offshore icing are wave impact and wind induced sea spray (Dehghani-Sanij et
54 al., 2017). Sea spray provides nuclei for ice clouds at high latitudes where airborne dust is sparse, being lofted by
55 bursting bubbles and droplets from white-capped waves (Russell, 2015; Dehghani-Sanij et al., 2017). Ice
56 accumulation from spray raises the center of gravity of ships, which can cause loss of stability and lead to capsizing
57 (Guest and Luke, 2005). Ice accumulation is believed to have caused the recent losses of three ships, including 1)
58 the *Destination*, which sank near St. George Island, Alaska in 2017 (Destination likely sank after accumulating ice
59 in heavy freezing spray, report says, 2023); 2) the *Scandies Rose*, which sank southeast of Kodiak, Alaska, in 2019
60 (NTSB announces the probable cause of the sunken Scandies Rose, 2023); and 3) the *Onega*, which sank in the
61 Barents Sea in 2020 (Icing believed to cause sinking of fishing boat in Barents Sea, 17 missing, 2023). To mitigate
62 ice-induced accidents, inclement weather forecasts are furnished for coastal waters. A Coastal Waters Forecast,
63 delivered by the National Weather Service, will contain a “freezing spray advisory” if freezing water droplets can
64 accumulate on vessels due to a combination of sea surface temperature (SST), wind speed, air temperature, and
65 vessel motion (Glossary - NOAA’s National Weather Service, 2023). At rates greater than 2 cm h^{-1} , the advisory
66 becomes a “heavy freezing spray watch”.

67
68 Wind turbines can modify the amount and severity of freezing conditions via competing effects. Enhanced
69 turbulence caused by spinning blades transports temperatures (either warmer or colder) from aloft to lower altitudes
70 within the rotor-swept region or near the surface. In stable stratification, warmer potential temperatures are
71 transported downward, which introduces a near-surface warming effect, and vice versa in unstable conditions (Fitch
72 et al., 2013; Rajewski et al., 2013; Xia et al., 2016; Siedersleben et al., 2018; Tomaszewski and Lundquist, 2020).
73 However, recent research suggests taller turbines may reverse this phenomenon (Golbazi et al., 2022). As the winter
74 months feature more frequent unstable stratification along the U.S. East Coast (Bodini et al., 2019), turbine-induced



75 cooling may increase the potential for near-surface icing. In contrast, turbines harness momentum from the flow,
76 which reduces the downwind wind speed (Nygaard, 2014; Platis et al., 2018; Schneemann et al., 2020). A reduction
77 in wind speed conversely reduces the potential for icing (Dehghani-Sanij et al., 2017). Thus, it is crucial to
78 understand how large-scale wind deployment across the mid-Atlantic will modify the regularity and intensity of
79 freezing conditions.

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81 Herein, we employ numerical weather prediction modeling to quantify the baseline offshore icing risk and the
82 wind plant post-production effects. Section 2 outlines the modeling setup and discusses the techniques for discerning
83 freezing events and cold air outbreak. Section 3 reports results for the spatiotemporal icing risk, causal factors, and
84 the adjustments by wind plants. Section 4 offers concluding remarks and discussion.

85

86 **2 Methods**

87 **2.1 NOW-23**

88 We explore annual variability of freezing conditions using the 2023 National Offshore Wind (NOW-23) dataset
89 (NREL, 2020; Bodini et al., 2023). This dataset provides an offshore wind resource spanning all offshore regions of
90 the United States at 5 min resolution for up to 22 years. For the mid-Atlantic region, NOW-23 was validated against
91 observations from three ZephIR ZX300M floating lidars (Pronk et al., 2022). We acquire model output at an hourly
92 temporal resolution for the 20-year period from 01 January 2000 to 31 December 2020. NOW-23 employs the
93 Weather Research and Forecasting (WRF) model (Powers et al., 2017) version 4.2.1. A parent domain feeds into an
94 inner nested domain with horizontal grid resolutions of 6 km and 2 km, respectively. Both domains have a vertical
95 grid resolution of 5 m near the surface with stretching to 45 m aloft, using 61 vertical levels up to a 50 hPa top. The
96 European Centre for Medium Range Weather Forecasts 5 Reanalysis (ERA5) dataset supplies hourly initial and
97 boundary conditions at a 30 km resolution to WRF (Hersbach et al., 2020). NOW-23 employs the MYNN2
98 planetary boundary layer and surface layer (Nakanishi and Niino, 2006) schemes, eta microphysics (Ferrier et al.,
99 2002), the Noah Land Surface Model (Tewari et al., 2004), the rapid radiative transfer model for shortwave and
100 longwave radiation (Iacono et al., 2008), and the Kain–Fritsch cumulus parameterization (Kain, 2004) in the
101 outmost domain only.

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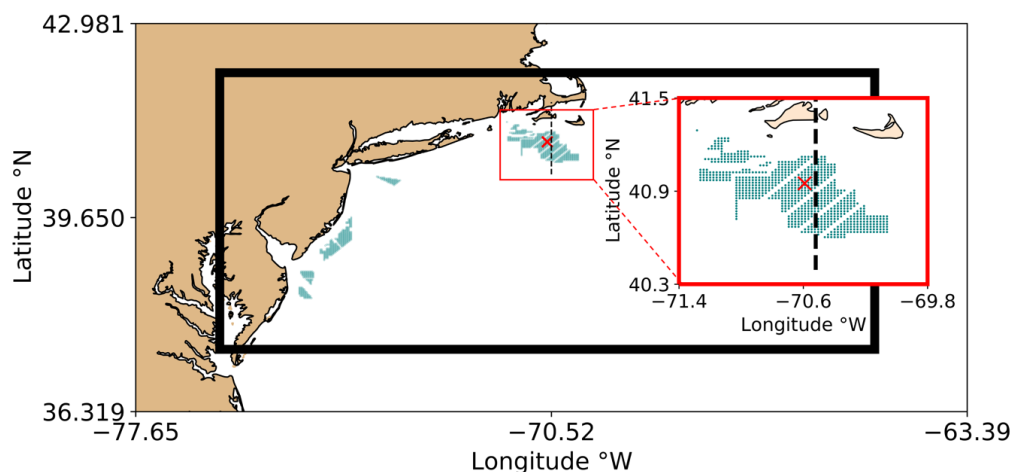
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2.2 NOW-WAKES

104 We explore the seasonal variability and impacts of wind plants on icing conditions using high-fidelity numerical
105 weather prediction simulations over the period 01 September 2019 to 31 August 2020. These validated WRF version
106 4.2.1 simulations are described in detail in Rosencrans et al. (2023) but are summarized here for the reader's
107 convenience. This period is chosen for the availability of lidar measurements for validation of the wind speed
108 profile. A parent domain hosts an inner nest with horizontal grid resolutions of 6 km and 2 km, respectively (Figure
109 1). Both domains have a vertical grid resolution of 10 m near the surface with stretching aloft, using 54 vertical
110 levels up to a 50 hPa top. The inner domain outputs data at an instantaneous history file frequency of 10 minutes.
111 Constant time steps are set to 18 s and 6 s in the outer and inner domains, respectively. Initial and boundary
112 conditions are also supplied by the hourly 30 km ERA5 dataset (Hersbach et al., 2020). Lower boundary conditions



113 are provided as SST by the UK Met Office Operational Sea Surface Temperature and Sea Ice Analysis dataset
114 (Donlon et al., 2012) and show good agreement during validation against mid-Atlantic bight buoys (Redfern et al.,
115 2023). Physics parameterizations include the MYNN2 planetary boundary layer and surface layer (Nakanishi and
116 Niino, 2006), the Noah Land Surface Model (Niu et al., 2011), the New Thompson microphysics (Thompson et al.,
117 2008), the rapid radiative transfer model for longwave and shortwave radiative transfer (Iacono et al., 2008), and the
118 Kain–Fritsch Cumulus (Kain, 2004) schemes. The Kain–Fritsch cumulus parameterization applies to the parent
119 domain only.
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122 **Figure 1. Modeling domains. The entirety of the outer domain with inner domain is shown, outlined by the black**
123 **rectangle. The red square is zoomed in on the Rhode Island–Massachusetts (RIMA) block to enhance visibility. Turbines**
124 **are shown as teal dots. The red “X” indicates the point of interest (POI) where time series are acquired. The dashed black**
125 **line is a cross section extending through the RIMA block.**

126
127 We incorporate the effects of wind turbines using the WRF wind farm parameterization (WFP) (Fitch et al.,
128 2012). WFP simulations feature wind plant layouts of the lease areas and include 1,418 turbines (Figure 1, Table 1).
129 The WFP incorporates the effects of turbines by implementing a drag-induced deceleration of wind flow and an
130 addition of turbulence at model levels intersecting the rotor area. We execute WFP simulations adding both 0 % and
131 100 % turbulent kinetic energy (TKE) (Rosencrans et al., 2023), although a smaller value of 25 % agrees better with
132 neutrally stratified large-eddy simulations (Archer et al., 2020). The differences in icing results between 0 % and
133 100 % added TKE are slight, so we report those from 100 % added TKE only. This work utilizes 12 MW GE
134 Haliade wind turbines with a 138 m hub height and 215 m rotor diameter, which is scaled by Beiter et al. (2020)
135 from a 15 MW reference turbine. We carry out separate simulations using both no wind farms (NWF) and wind
136 farms (WFP) for the full year-long period from 01 September 2019 to 31 August 2020 (Table 1).
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Table 1. List of simulations characterized by turbine characteristics.

Simulation Type	Turbine Type	Period	Added TKE	# Turbines
No Wind Farms (NWF)	N/A	09/2019–08/2020	N/A	0
Wind Farm Param. (WFP)	12 MW	09/2019–08/2020	0 %	1,418
Wind Farm Param. (WFP)	12 MW	09/2019–08/2020	100 %	1,418

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2.3 Freezing hours detection

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Ice accretion occurs when supercooled water freezes upon contact with objects. The largest contributions to sea spray icing are provided by the bursting of bubbles and advection of spray from white-capped waves (Dehghani-Sanij et al., 2017). Further aloft, supercooled water can be introduced by liquid precipitation and fog. In the presence of moisture, there are three key variables that dictate offshore freezing conditions: wind speed, SST, and air temperature (Overland et al., 1986; Overland, 1990; Guest and Luke, 2005; Dehghani-Sanij et al., 2017; Line et al., 2022).

We define input freezing conditions following the most liberal thresholds defined by the latter studies (Guest and Luke, 2005; Dehghani-Sanij et al., 2017; Line et al., 2022), which produce more freezing events, to compensate for a negative wind speed bias in unstable stratification (Rosencrans et al., 2023), which mitigates freezing occurrence. These criteria require 1) wind speeds in excess of 9 m s^{-1} , 2) air temperatures below -1.7° C , and 3) SST less than 7° C . The skin temperature (WRF output variable “TSK”) is used because the SST field inherits coarse blocks of missing data around coastlines from the ERA5 dataset. The resulting spatial maps are masked by the land use (WRF output variable “LU_INDEX”) to ensure freezing conditions over land are not counted. The number of 10 min timestamps where these criteria are met each month are recorded for all simulations.

Sea spray-induced icing can affect structural integrity, blade aerodynamics, and crew safety. As sea spray often lofts to between 5 and 20 m above sea level (Dehghani-Sanij et al., 2017), we detect possible icing conditions at the lowest model level of 10 m and at 20 m. We further consider riming conditions at the 138 m hub height by including an additional criterion for the presence of liquid rainwater, as precipitation-induced ice can generate a considerably higher ice accretion rate than fog-induced icing (Gao and Hong, 2021).

2.4 Ice accumulation rate

A predictability function assesses the likelihood for freezing in the presence of sea spray. We assess the predictability of icing conditions at the point of interest (POI) in the Rhode Island/Massachusetts (RIMA) block (Figure 1) separately from the NOW-WAKES and the NOW-23 datasets. The predictability (*PPR*) for sea spray-induced ice formation follows:

$$PPR = \frac{V_a(T_f - T_a)}{1 + 0.4(T_s - T_f)} \quad (1)$$



169 where V_a is the wind speed, T_f is the temperature threshold of -1.7°C , T_a is the air temperature, and T_s is the SST
 170 (Guest and Luke, 2005; Overland et al., 1986; Overland, 1990). A humidity variable is not present in Eq. (1) due to
 171 the assumption that sea spray introduces a constant source of moisture during fast winds, that of which is required
 172 for nonzero PPR. A group of successive timestamps with nonzero PPR are considered the same event. Separate
 173 flagged timestamps occurring within 24 hours of each other span the same synoptic regime (Winters et al., 2019),
 174 and so the entire duration between the two flagged timestamps is considered one event.

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176 **Table 2. Icing rate by PPR. Rows delineate icing predictability (PPR), icing class, and ice accretion rate. Columns**
 177 **delineate the icing rate per PPR range. From Guest and Luke (2005).**

PPR	<0	0–22.4	22.4–53.3	53.3–83.0	>83.0
Icing Class	None	Light	Moderate	Heavy	Extreme
Icing Rate [cm h^{-1}]	0	<0.7	0.7–2.0	2.0–4.0	>4.0

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179 The magnitude of PPR can determine the rate of ice accretion (Table 2). The ice accretion rates are a general
 180 guideline developed for 20 to 75 m vessels; specific rates depend on the type of ship, its load, and its handling
 181 characteristics (Guest and Luke, 2005). For instance, a larger ship requires faster winds and taller waves for sea-
 182 spray–induced ice to accumulate on a higher deck but is more vulnerable to the prevailing wind direction due to
 183 reduced maneuverability. It is not known how these icing rates would apply to wind turbines or to the vehicles used
 184 to access offshore wind turbines.

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186 2.5 Cold air outbreak detection

187 Freezing conditions can be stimulated by the advection of cold continental air over a warmer maritime surface.
 188 The resulting temperature profile induces instability, which causes filamentary convective rolls that align to make
 189 cloud “streets” with parallel columns of ascending and descending air that transform into open convective cells
 190 further offshore (Geerts et al., 2022). An approach proposed by Vavrus et al. (2006) identifies a CAO by the
 191 magnitude and duration of anomalous air temperature, which we apply at the POI (Figure 1). This strategy requires
 192 that the near-surface temperature be at least 2 standard deviations below the wintertime average following Eq. (2):

$$193 \quad T < \bar{T} - 2(\sigma) \quad (2)$$

194 where T is the 10 m temperature, \bar{T} is the average 10 m temperature during the wintertime period, and σ is the
 195 standard deviation. The wintertime period spans November through March at a 10 min frequency to account for all
 196 non-zero-freezing predictability events. Again, successive timestamps with detected CAO are considered a single
 197 event, and separate events occurring within a 24 h span are conglomerated into the same event.

198

199 2.6 Atmospheric stability

200 Turbulence from wind turbines modifies the near-surface temperature based on the atmospheric stability or
 201 stratification. We calculate the modeled atmospheric stability using the Obukhov Length (L) (Monin and Obukhov,



202 1954) (Eq. 3), which delineates the height above the surface at which buoyant turbulence equals mechanical shear
203 production of turbulence, at a point centered on the RIMA block of lease areas:

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$$L = -\frac{u_*^3 \overline{\theta_v}}{\kappa g (\overline{w' \theta_v'})}$$
 (3)

205 where u_* (UST in WRF output) is the friction velocity, θ_v is the virtual potential temperature, κ is the von Kármán
206 constant of 0.4, g is gravitational acceleration of 9.81 m s^{-1} , and $\overline{w' \theta_v'}$ (HFX in WRF output) is the surface dynamic
207 heat flux converted into kinematic heat flux. Negative lengths between 0 m and -500 m imply unstable stratification
208 due to a positive heat flux (Archer et al., 2016). Conversely, lengths between 0 m and 500 m imply stable
209 stratification due to a negative heat flux. Lengths approaching negative or positive infinity imply neutral
210 stratification, as buoyancy is no longer a dominating factor. Each 10 min timestamp from the NWF run is assigned a
211 stability classification from November 2019 to March 2020.

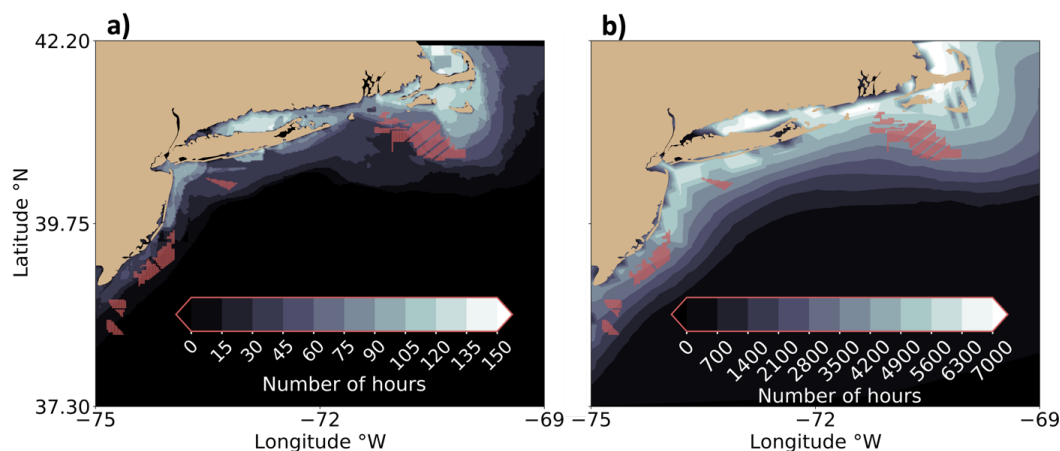
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213 **3 Results**

214 **3.1 Spatial variability of freezing conditions**

215 The percentage of occurrence of freezing conditions exhibits regional variability. The commonality of freezing
216 increases toward higher latitudes and near the coast where cold continental air advects over the ocean during the
217 winter (Figure 2). In general, the spatial icing pattern during the 2019–2020 winter season (Figure 2a) matches well
218 with the pattern over the 20-year period (Figure 2b). Freezing conditions shadow the mid-Atlantic coast but occur
219 less often along the New Jersey Bight where wind speeds decrease and temperatures warm. The commonality of
220 freezing conditions extends furthest offshore southeast of Nantucket and enhances in the Long Island Sound; both
221 regions feature local minima in mean January 2020 SST less than 5° C . To the north, the cyclonic current in the Gulf
222 of Maine transports cold surface water southward. East of Cape Cod, this current bifurcates, and a branch feeds cold
223 fresh water into the mid-Atlantic. Predominant northerly winter winds instigate onshore Ekman transport towards
224 the coast, which is favorable for downwelling (Shcherbina and Gawarkiewicz, 2008b). However, downwelling is not
225 always supported, as the mixed layer stratification is dominated by salinity (Shcherbina and Gawarkiewicz, 2008a),
226 leaving a cold pool near the surface.

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Figure 2. The number of hours freezing conditions occur during (a) the November 2019 to March 2020 period at 10 m in NWF and (b) the November to March period from 2000 to 2020 in NOW-23. Lighter contouring indicates more freezing hours. Red dots represent turbine locations.

Freezing conditions exhibit seasonal variability in NWF, starting at 0 hours in November, increasing through the winter, and falling to 0 again by April at all heights (Figure 3 and Figure. A1–A3). At the 10 m altitude, freezing conditions occur most often in January, up to 66 hours, with an offshore spatial extent of 57,420 km², or roughly 12 times the area of the wind plants. At 20 m, freezing conditions also occur most often in January, up to 70 hours, covering a total area of 61,924 km², or roughly 13 times the area of the wind plants. The 20 m height experiences more freezing hours than the 10 m height because average wind speeds are at least 0.25 m s⁻¹ faster around Nantucket, Cape Cod, and the Long Island Sound. The 138 m hub height has a smaller maximum of 29 hours during January in the Gulf of Maine, with a band extending south from Cape Cod, posing no threat to the lease areas (Figure A3). Although wind speeds increase aloft, the regularity of liquid water is not as consistent, as is near-surface sea spray.

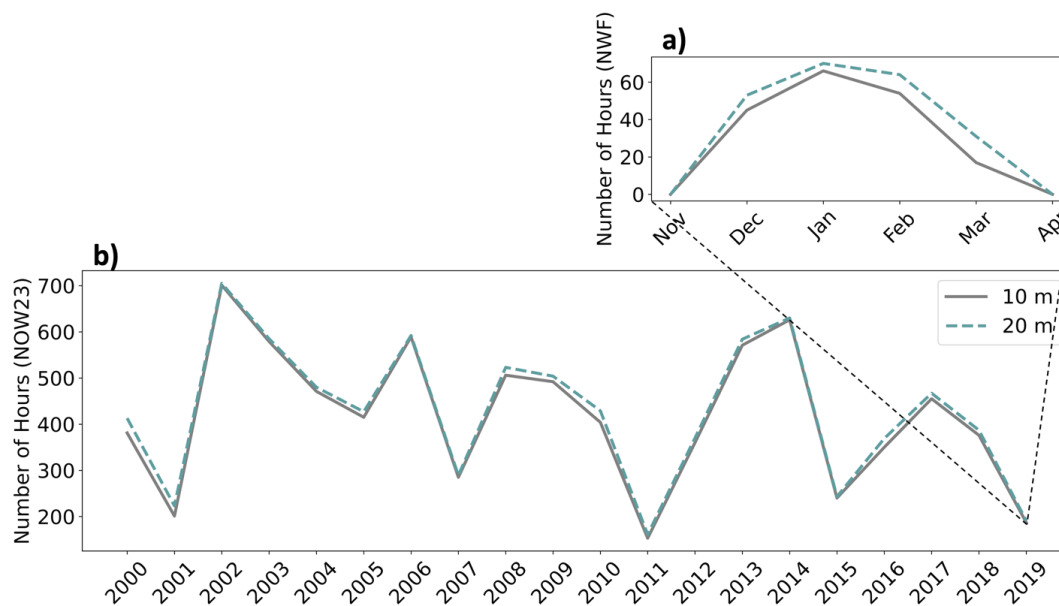


Figure 3. The maximum number of freezing hours over the OCS (a) seasonally and (b) annually from 2019 to 2020.

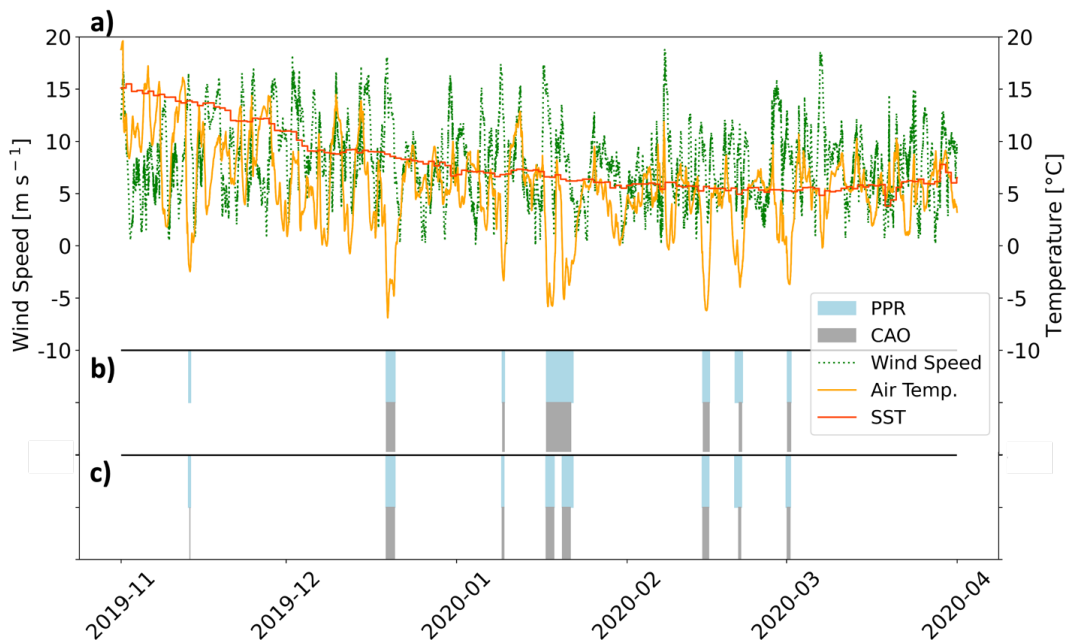
The 2019–2020 winter season was the mildest compared to other winters (Figure 3b). This winter season had the fewest number of freezing hours compared to other winters over the 20-year period, reaching 182 hours in NWF or 187 hours in NOW-23 at 10 m. At 20 m, the 2019–2020 winter season contains 218 hours in NWF or 191 hours in NOW-23. The greatest number of freezing hours occurs during the 2002–2003 season, with 701 total hours at 10 m and 705 hours at 20 m. While the 20 year slope shows a decrease, it is not statistically significant using the Mann–Kendall (M–K) test (Hussain and Mahmud, 2019). P-values for the maximum number of hours (found across the OCS) (Figure 3b) and for the number of freezing hours at the POI (Figure 1) are 0.20 and 0.12, respectively. We additionally applied the seasonal M–K test (Hirsch et al., 1982) to account for upward and downward trends throughout the year on monthly mean PPR, monthly maximum PPR, and the monthly total number of freezing hours at the POI. Neither test returned a statistically significant trend.

3.2 Freezing conditions and cold air outbreak

Investigating all events with a non-zero freezing PPR at the POI (Figure 1) reveals similar synoptic trends. We identify seven events with freezing sea spray (FSS) conditions with a total duration of 253 hours from November 2019 to March 2020. All times during the 2019–2020 winter period with nonzero PPR contain light ice accumulation of less than 0.7 cm h^{-1} (Table 2). During each FSS event, higher relative pressure resided to the southwest throughout the Great Plains, Appalachia, or the Great Lakes with lower relative pressure to the northeast around Nova Scotia and Newfoundland. The largest pressure gradient forces occurred during the two January events of up to $4 \times 10^{-3} \text{ Pa m}^{-1}$, or roughly 4 times the pressure gradient force required for a 10 m s^{-1} geostrophic



266 wind in the midlatitudes (Parish et al., 2007). Most events feature either an outflow boundary or cold front. This
267 pressure regime directs quasi-geostrophic flow near the surface toward the southeast, introducing cold continental
268 air offshore. During the winter, the prevailing wind direction is northwesterly across the mid-Atlantic OCS (Bodini
269 et al., 2019) because regions of land mass feature higher surface pressure than the surrounding ocean and the
270 Bermuda High retreats to the east.
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272 **Figure 4. (a) Time series of wind speed (green dotted), 10 m air temperature (orange), and SST (red) from November**
273 **2019 to April 2020 at the downwind edge of Vineyard Wind (Figure 1). Light-blue shading indicates the duration of**
274 **nonzero PPR, and gray shading indicates the duration of detected CAO from (b) NWF and (c) NOW-23.**
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277 Most offshore freezing events coincide with CAO. We detect six CAO events in NWF with a total duration of
278 200 hours (Figure 4b). The typical durations of CAO events are 3 hours shorter than FSS, with 78.9 % of flagged
279 FSS timestamps having CAO present. Overall, six of the seven FSS events occur in conjunction with a CAO (Figure
280 4b), with November air temperatures not cold enough to be flagged as CAO candidates. We note that in NWF, all
281 seven FSS events coincide with CAO at the northeast edge of the RIMA block which is nearer to the introduction of
282 cold continental air.

283
284 Common between events are fast wind speeds and cold 10 m air temperatures; SST plays a secondary role for
285 its weak variability (Figure 4a). The average wind speed during FSS events is 10 m s^{-1} with gusts exceeding 15 m
286 s^{-1} during four events. Nonzero PPR does not occur until after the wind speed peaks, when cold air temperatures



287 sweep in, averaging minimum temperatures of -5°C (Figure 4a). This wind speed–temperature dynamic poses
288 challenges for grid planners because wind energy generation reduces during periods of peak heating load.

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290 Over the 2019–2020 winter in the NOW-23 dataset, eight total events are flagged as candidates for FSS because
291 the longest event in January 2020 (Figure 4b) is split among two separate events; all eight events have a
292 corresponding CAO (Figure 4c). Over the 20-year period, *all* CAO events occur in conjunction with an FSS event
293 (positive PPR). However, many FSS events occur without CAO present meaning that CAO is only one of the
294 drivers, and large interannual variability can exist. For instance, while 97 % of CAO timestamps concur with FSS
295 during the 2011–2012 season, only 9 % do during the 2013–2014 season.

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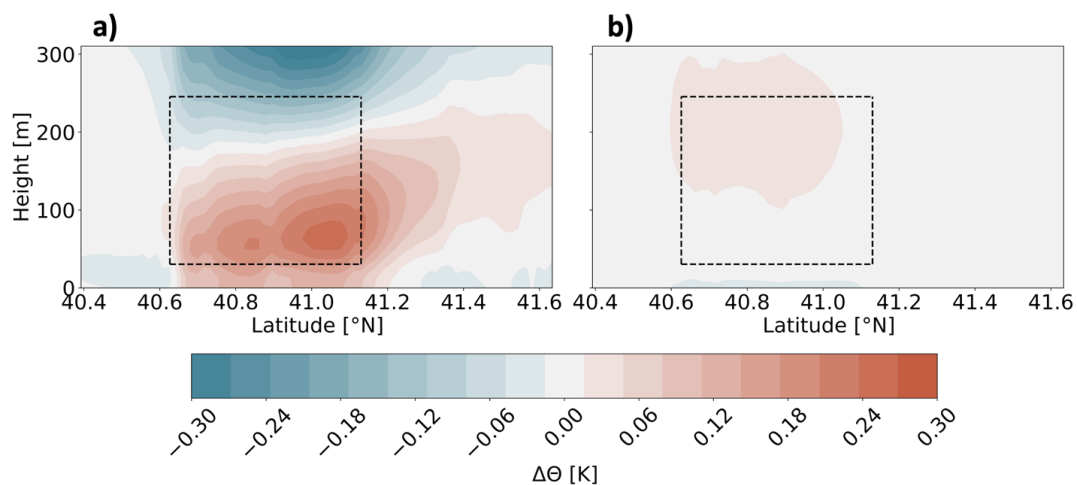
297 The 2019–2020 winter ice accumulation rate is similar to other winters. The average PPR during freezing
298 events from 2019 to 2020 is 4.3, which corresponds to a light ice accumulation rate of less than 0.7 cm h^{-1} (Table 2).
299 Over the 20-year period, the average PPR among events is 8.1, which corresponds to the same accumulation rate.
300 The 2003–2004 winter period features the greatest mean PPR of 15.7, which also corresponds to a light ice
301 accumulation rate. During this period, a moderate risk for icing occurred 18 % of the time, and a heavy risk occurred
302 3 % of the time, corresponding with icing rates between $0.7\text{--}2.0\text{ cm h}^{-1}$ and $2.0\text{--}4.0\text{ cm h}^{-1}$, respectively, and
303 possibly triggering heavy freezing spray watches in the NWS advisory.

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305 **3.3 Modifications by wind plants**

306 Icing is more probable in cold temperatures. In unstable conditions, which occur 64 % of the time from
307 November 2019 through March 2020 in NWF assessed at the POI, wind turbines introduce near-surface cooling,
308 which could increase the likelihood of icing. For instance, mean cooling and warming during unstable conditions
309 reach magnitudes up to -0.041 K at the surface and 0.022 K within the rotor-swept region, respectively, along a
310 cross section extending through the RIMA block (Figure 1, Figure 5b). During stable conditions, which occur 25 %
311 of the time from November through March, cooling aloft reaches up to -0.34 K , and near-surface warming reaches
312 0.26 K (Figure 5a). Near-surface cooling exists adjacent to the wind plant cluster.

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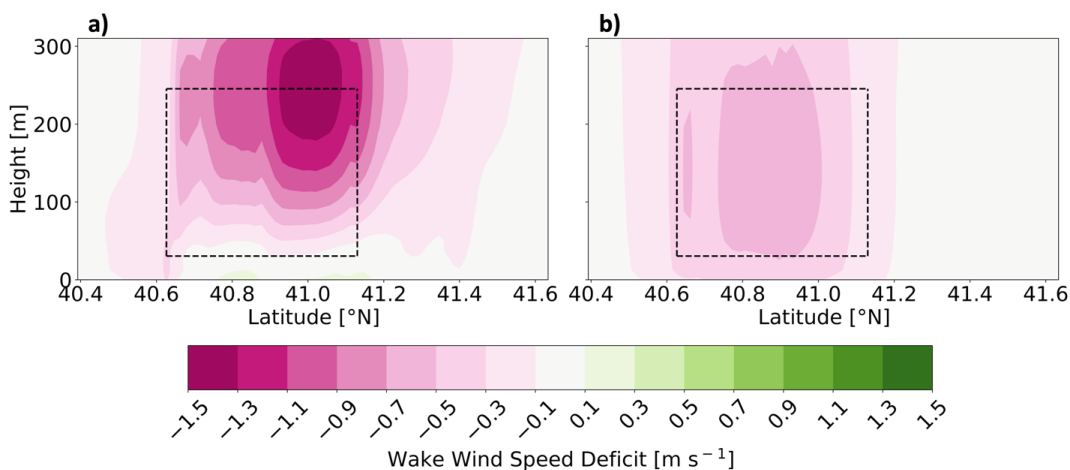
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Figure 5. The mean (WFP-NWF) potential temperature difference during (a) stable stratification and (b) unstable stratification, from November 2019 to March 2020. The cross section spans the RIMA block of lease areas (Figure 1). Red contouring indicates warming, and blue indicates cooling. Dashed lines outline the wind plant area and rotor-swept region.

Conversely, the reduction of wind speeds in the wake modifies the chance for icing. In stable conditions, the wake wind speed deficit is largest, reaching -1.4 m s^{-1} near the top of the rotor-swept plane, reducing the chance for freezing. Because vertical motion is suppressed in stable stratification, winds increase and flow around and under the wind plant area (Figure 6a), reaching a subtle enhancement near the surface of 0.17 m s^{-1} . In unstable stratification, available buoyant turbulence promotes mixing which transports momentum from above the rotor-swept region down to within the wake. The injection of momentum allows wake wind speeds to recover, leaving a smaller maximum averaged wake deficit of -0.57 m s^{-1} (Figure 6b). There is no enhancement of wind speeds adjacent to the RIMA block along the cross section in unstable conditions.



329

330 **Figure 6. The mean (WFP-NWF) wind speed difference during (a) stable and (b) unstable stratification, from November**
 331 **2019 to March 2020. The cross section spans the RIMA block of lease areas (Figure 1). Pink contouring indicates a wind**
 332 **speed reduction, and green indicates wind speed enhancement. Dashed lines outline the wind plant area and rotor-swept**
 333 **region. Note the very small enhancement of wind speeds near the surface in stable conditions.**

334

335 Despite near-surface cooling, net freezing conditions occur less often when diagnosed using wind speed, air
 336 temperature, and SST criteria because of the wake wind speed reduction. At 10 m, the turbine–atmosphere
 337 interaction alters possible icing conditions the most in January and February, with a maximum reduction by 13 hours
 338 (Table 3). At 20 m, wind plants cause a reduction by up to 17 hours in January. In each case, the reduction in
 339 possible freezing conditions is spatially coincident with the wind plant areas (Figure 7). At the 138 m hub height, the
 340 change to the number of freezing conditions maximizes in December, with a reduction by 6 hours.

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Table 3. The turbine-induced change in freezing hours by month and height.

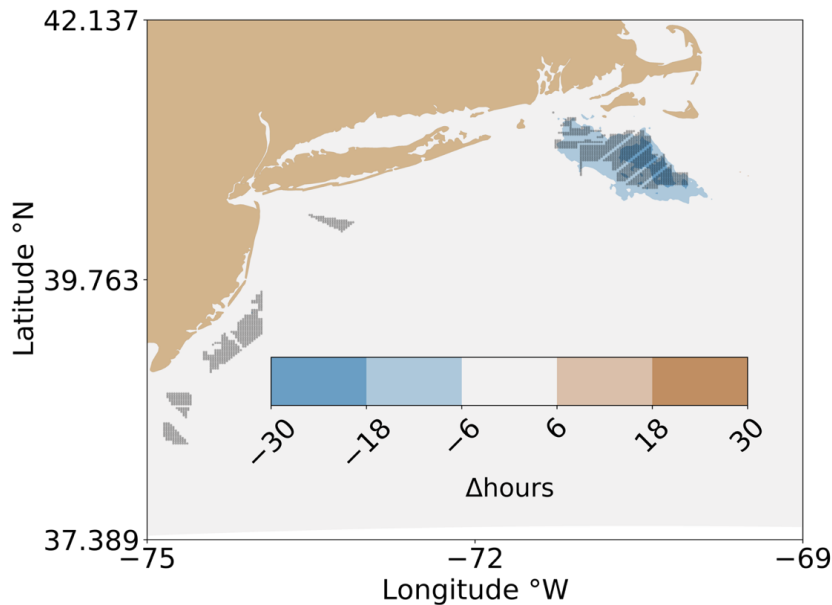
Change in Number of Freezing Hours Throughout Domain						
	November	December	January	February	March	April
10 m	0	−2	−13	−13	−10	0
20 m	0	−4	−17	−16	−12	0
138 m	0	−6	−3	−3	−3	0

343

344 The introduction of wind turbines also increases the chance for freezing surrounding the wind plants, reaching
 345 maxima in March of 5 and 6 hours at 10 m and 23 m, respectively (Figure A4, Figure A5). Flow acceleration is
 346 present adjacent to the wake as winds deflect around the clusters (Stoelinga et al., 2022; Golbazi et al., 2022).
 347 However, freezing enhancement is isolated to a speckled pattern (Figure 7), does not coincide with the wind speed
 348 enhancement, and thus may result from numerical noise introduced by the WFP (Ansell et al., 2018; Lauridsen and
 349 Ansell, 2018). As modifications to the percentage of freezing conditions at the hub height are not spatially
 350 coincident with the lease areas or prevailing wind, these changes may also result from numerical noise (Figure A6).



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Figure 7. The (WFP-NWF) change in percentage of freezing hours at 10 m November 2019 to March 2020. Blue contours indicate a percentage reduction, and orange contours indicate a percentage increase.

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Overall, the presence of wind turbines has a minimal impact to the number of hours freezing conditions occur by means of icing PPR at the POI. The duration of nonzero PPR over the November through March winter period increases by 3 hours, or from 253 to 256 hours total, at a point centered on the RIMA block. The total duration of CAO decreases by 1 hour, or from 200 hours to 199 hours, after the installation of wind plants. The total number of events (seven) does not change in the presence of wind turbines, and all flagged timestamps still cause light icing of less than 0.7 cm h^{-1} .

4 Conclusions

Here, we assess the threat of freezing conditions due to sea-spray icing and hub-height riming on wind turbines. The simulation study encompasses the mid-Atlantic Outer Continental Shelf based on a 20-year WRF dataset from 01 January 2000 to 31 December 2020 and another WRF dataset using year-long simulations from 01 September 2019 to 31 August 2020. In each case, we focus on the wintertime period from November through March. We consider the present icing risk from simulations with no wind farms (NOW-23, NWF) and assess the post-production adjustments by incorporating the effects of turbines (WFP) in a full buildout of the wind plant lease areas.

We detect seven events flagged for freezing sea spray conditions in NWF with a total duration of 253 hours during the November 2019 to March 2020 period. All times during the period with nonzero icing predictability



374 (PPR) contain light ice accumulation of less than 0.7 cm h^{-1} , which is typical for the region as assessed from 2000 to
375 2020. Centered at the RIMA block of lease areas, six of the seven events have an associated CAO in NWF during
376 the 2019–2020 winter, and all seven events have corresponding CAO at the northeast end of the block. In the NOW-
377 23 dataset over the same period, eight total events are flagged, and all eight correspond with CAO. From 2000 to
378 2020, every CAO event has a corresponding freezing sea spray event, although not all freezing events have attendant
379 CAO. Thus, offshore freezing may be forecast with reasonable fidelity through accompanying CAO, although other
380 drivers exist. There is strong teleconnection between anomalous arctic sea level pressure sea level pressure and
381 CAO, as 93 % of CAO events in the eastern U.S. contained an antecedent positive arctic sea level pressure anomaly
382 a week in advance (Vavrus et al., 2006).

383

384 Freezing conditions exhibit spatial variability. The hazards intensify toward higher latitudes, near the land
385 surface where cold air advects offshore, and by Nantucket and the Long Island Sound where SSTs are colder.
386 Freezing conditions at the hub height from supercooled liquid water are less frequent. The icing hazard is greatest
387 during January when wind speeds are fast and temperatures are cold. At 10 m in January, favorable conditions for
388 icing occur up to 66 hours. At 20 m in January, the duration of icing conditions increases to 70 hours. Finally, at the
389 hub height, freezing conditions occur for up to 29 hours in the Gulf of Maine and pose no risk to the lease areas.
390 Overall, the 2019–2020 winter period is the mildest winter when considering the 20-year period. Although the
391 2019–2020 winter season has the fewest number of freezing hours, all winters contain light ice accumulation rates of
392 0.7 cm h^{-1} .

393

394 The introduction of large wind plants makes a small impact on the icing risk within the wind plant clusters. In
395 wintertime unstable conditions, which occur 64 % of the time from November 2019 through March 2020, wind
396 turbines introduce a mean near-surface cooling effect. Despite the enhanced freezing risk from supplementary
397 cooling, slower wind speeds in the wake mitigate the hazard. Mean reductions in wind speed within the wakes reach
398 up to -0.57 m s^{-1} in unstable conditions with an introduction of cooler air up to -0.041 K . As assessed using wind
399 speed, air temperature, and SST criteria, the change in icing risk over the 2019–2020 wintertime period is a net
400 reduction, by up to 13 hours at 10 m. At 20 m, mitigation reaches up to 17 hours. The alleviation by slower wind
401 speeds is largest within the RIMA block of wind plants which contains the greatest number of turbines. When
402 assessed using icing PPR centered on the RIMA block, the number of hours freezing conditions and CAO occur
403 change by 3 and -1 hours, respectively. However, the 2019 through 2020 winter period is the mildest winter, so the
404 introduction of wind plants may make more significant changes during harsher winters.

405

406 Future OCS winter storm frequency may differ due to climate change. For instance, warming Arctic
407 temperatures, which reduce the meridional geopotential height gradient between the Arctic and midlatitudes, can
408 weaken the jet stream. Slower zonal winds and more pronounced Rossby waves amplify the transport of extreme
409 winter weather to the midlatitudes (Cohen et al., 2020). Future East Coast storm activity and temperature may
410 experience modulations based on large-scale teleconnections such as El Niño and the North Atlantic Oscillation



411 (Hall and Booth, 2017). Further, Arctic amplification may increase the strength of teleconnection found between
412 positive Arctic sea level pressure anomalies and CAO (Vavrus et al., 2006).

413

414 Finally, we assume that sea spray provides a consistent moisture flux at 10 and 20 m during fast wind
415 conditions, that the droplet size of spray is homogeneous, and that the number distribution by height is constant.
416 Future studies may benefit from coupling WRF with wave models, such as Wave Watch III (Tolman et al., 2019)
417 and Simulating Waves Nearshore (SWAN Team, 2020) for precise modeling of wave induced sea spray and for
418 current dynamics, such as stratified cold pooling around Cape Cod. New satellite methods are being developed to
419 quantify occurrences of freezing sea spray (Line et al., 2022), and future developments should compare the FSS
420 criteria to satellite observations of FSS.

421

422 **5 Code and data availability**

423 The dataset and files that support this work are publicly available. The ERA5 initial and boundary conditions can be
424 downloaded from the ECMWF Climate Data Store at [https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-
425 era5-pressure-levels?tab=form](https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-pressure-levels?tab=form). Shapefiles including the bounds for wind energy lease areas are at
426 <https://www.boem.gov/renewable-energy/mapping-and-data/renewable-energy-gis-data>. Wind turbine coordinates
427 and their power and thrust curves are provided at <https://zenodo.org/record/7374283#.Y4YZxC-B1KM>. WRF
428 namelists for NWF and WFP simulations may be acquired from [https://zenodo.org/record/7374239#.Y4YaOy-
429 B1KM](https://zenodo.org/record/7374239#.Y4YaOy-B1KM). The NOW-23 simulation output data are available in HDF5 format at <https://doi.org/10.25984/1821404>.

430

431 **6 Author contributions**

432 Conceptualization: JKL, MO. Resources: MO, NB. Methodology: DR, JKL. Software: DR. Formal analysis and
433 visualization: DR. Investigation: DR and JKL. Writing – original draft: DR and JKL. Writing – review and editing:
434 all co-authors. Supervision: JKL.

435

436 **7 Competing interests**

437 At least one of the (co-)authors is a member of the editorial board of Wind Energy Science. Furthermore, Mike
438 Optis co-authored the submitted manuscript while an employee of the National Renewable Energy Laboratory. He
439 has since founded Veer Renewables, which recently released a wind modeling product, WakeMap, which is based
440 on a similar numerical weather prediction modeling framework as the one described in this manuscript. Data from
441 WakeMap is sold to wind energy stakeholders for profit. Public content on WakeMap include a website
442 (<https://veer.eco/wakemap/>), a white paper ([https://veer.eco/wp-
443 content/uploads/2023/02/WakeMap_White_Paper_Veer_Renewables.pdf](https://veer.eco/wp-content/uploads/2023/02/WakeMap_White_Paper_Veer_Renewables.pdf)) and several LinkedIn posts promoting
444 WakeMap. Mike Optis is the founder and president of Veer Renewables, a for-profit consulting company. Mike
445 Optis is a shareholder of Veer Renewables and owns 92 % of its stock.

446

447 **8 Acknowledgements**

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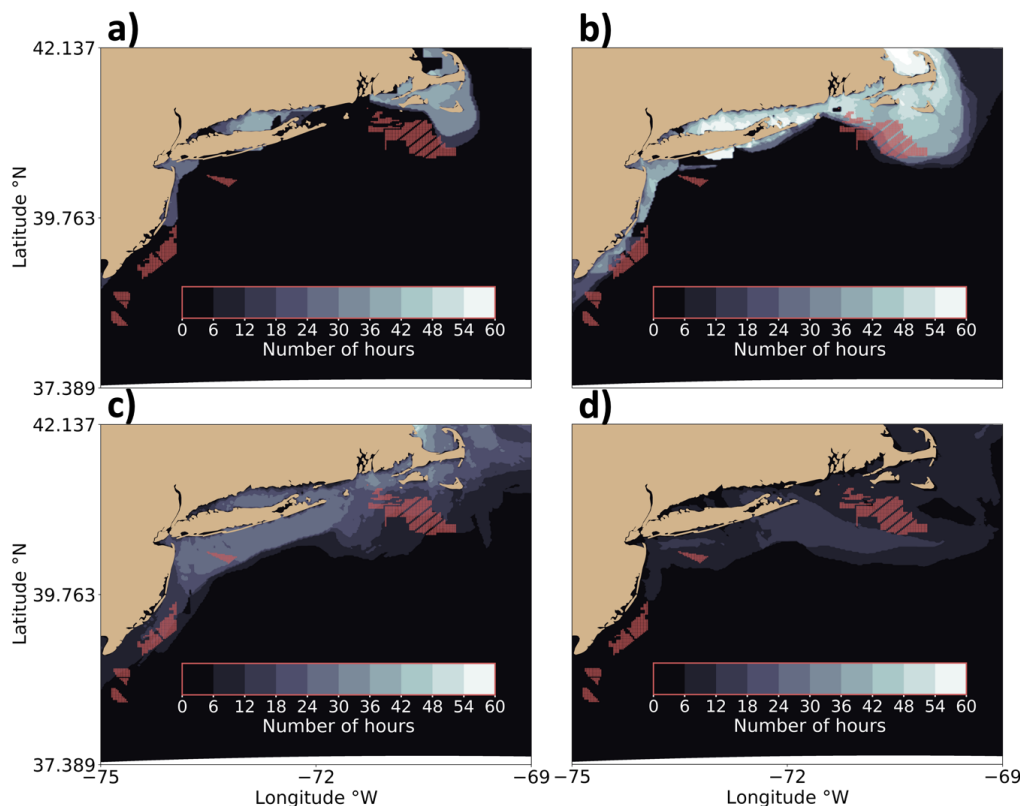
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462

463 9 Appendices

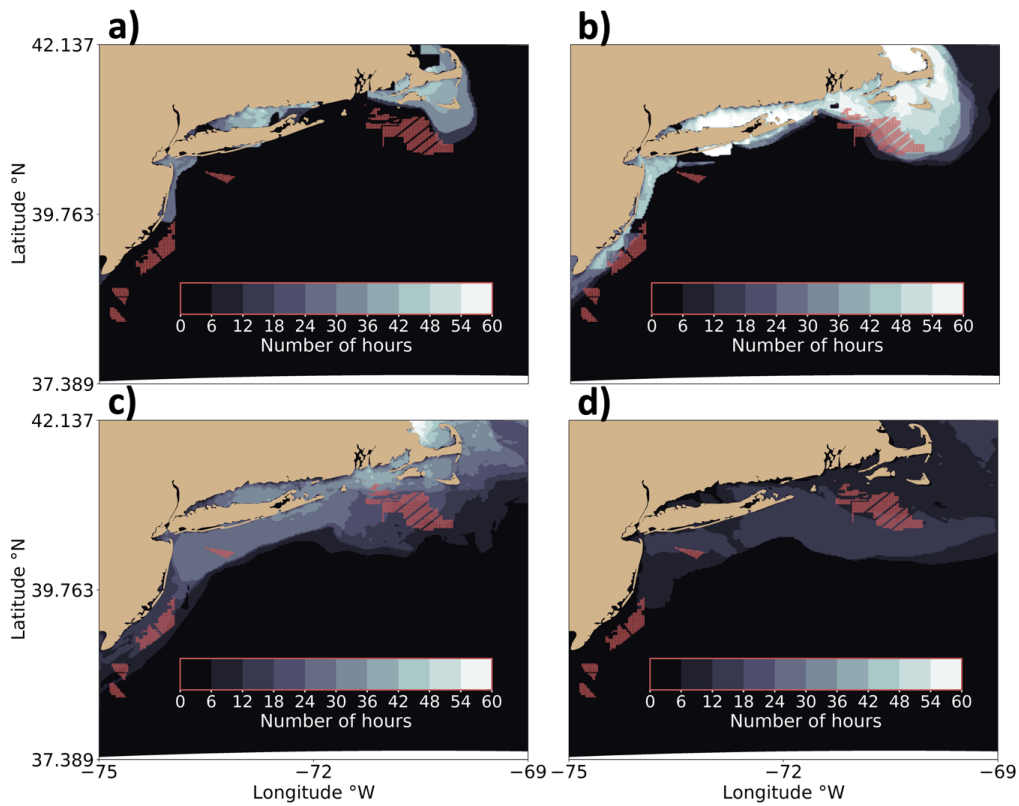
464 Appendix A



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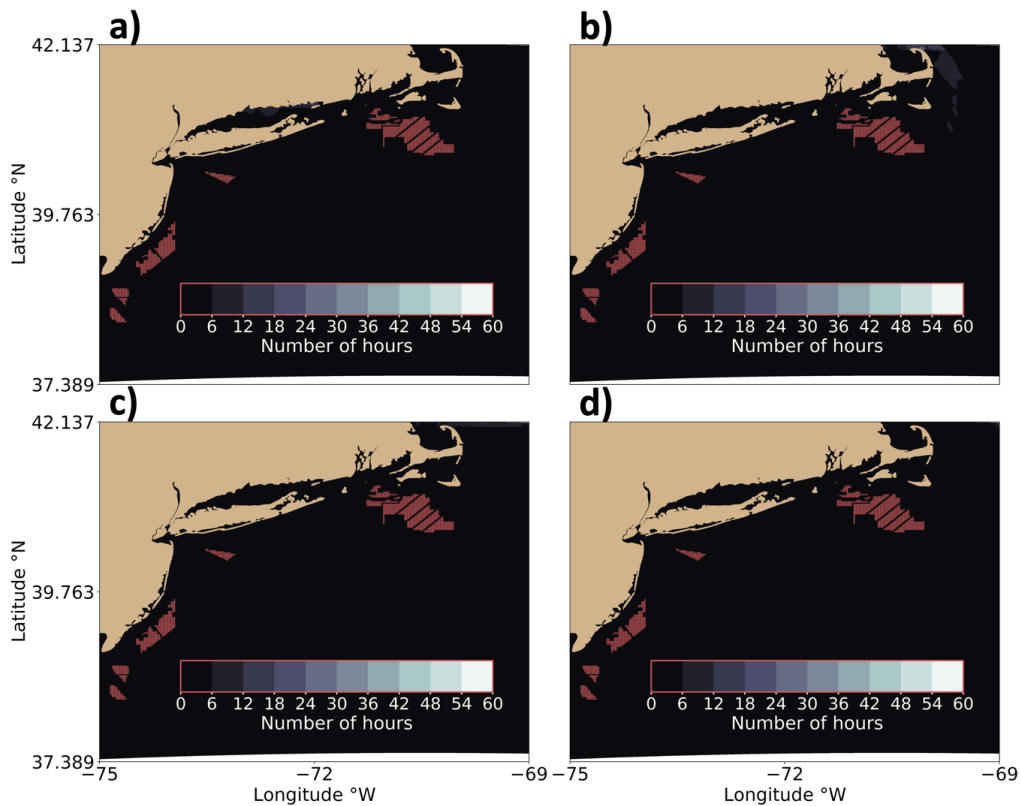
466 **Figure. A1.** The number of freezing hours at 10 m during (a) December 2019, (b) January 2020, (c) February 2020, and
467 (d) March 2020. Lighter contouring indicates higher percentages. Red dots indicate turbine locations.

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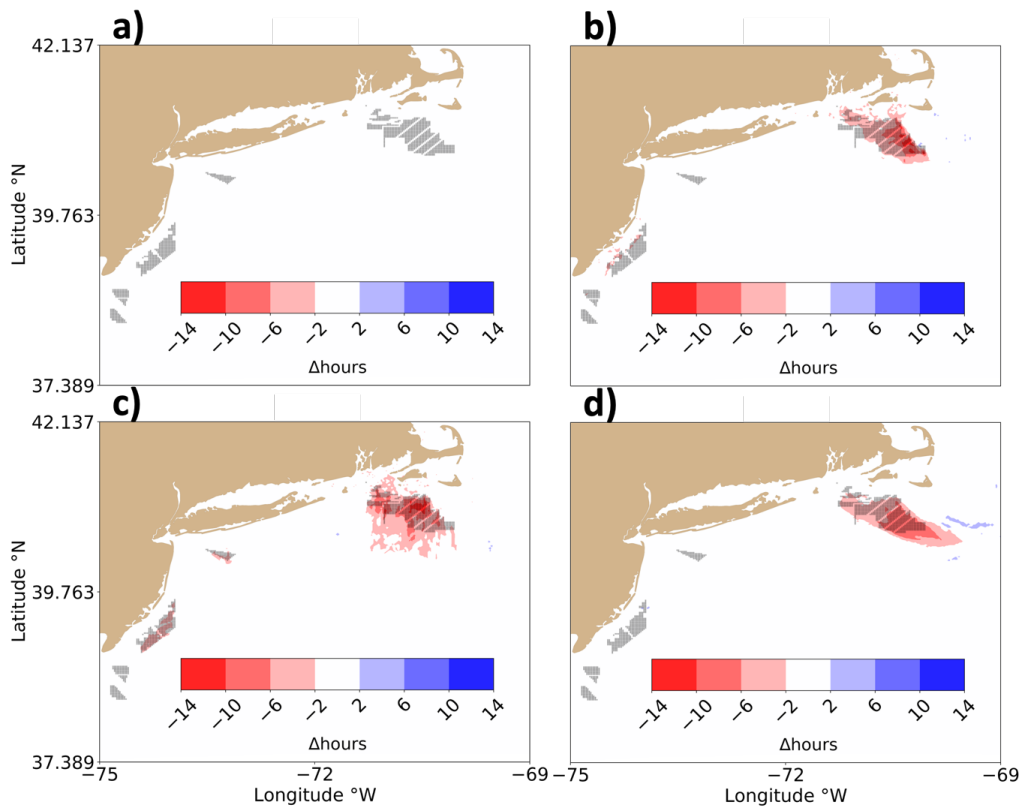
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Figure A2. The number of freezing hours at 20 m during (a) December 2019, (b) January 2020, (c) February 2020, and (d) March 2020. Lighter contouring indicates higher percentages. Red dots indicate turbine locations.



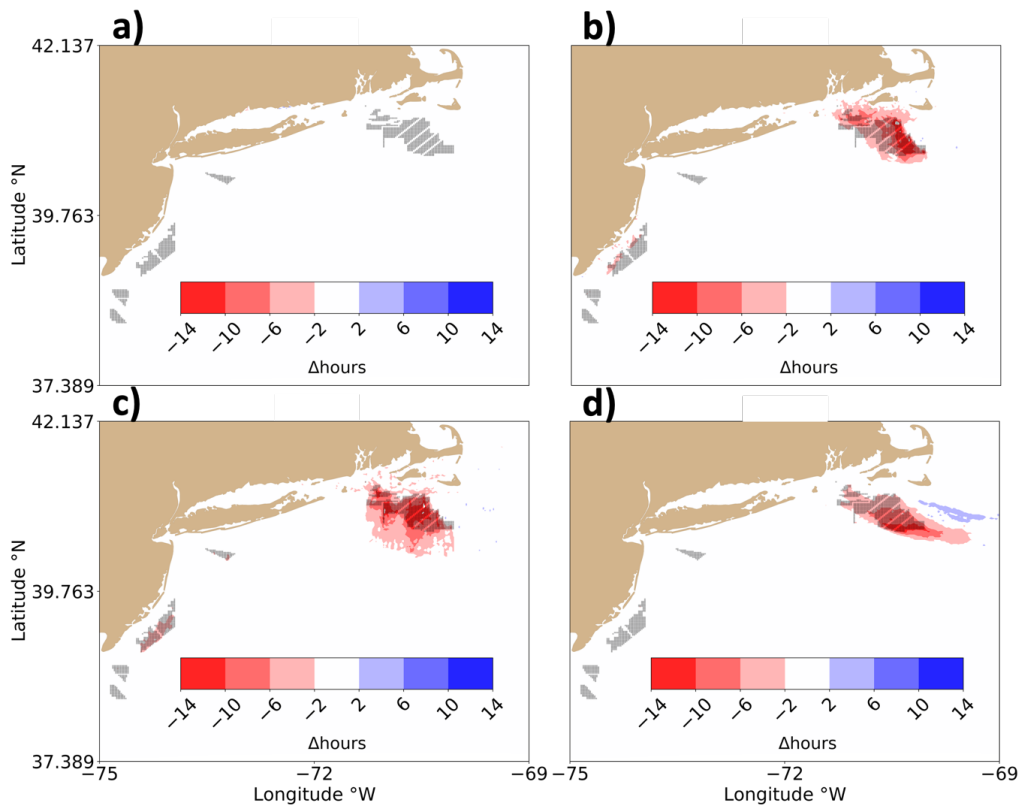
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Figure A3. The number of freezing hours at hub height during (a) December 2019, (b) January 2020, (c) February 2020, and (d) March 2020. Lighter contouring indicates higher percentages. Note the color scheme is different from Supplementary Figs. 1 and 2. Red dots indicate turbine locations.



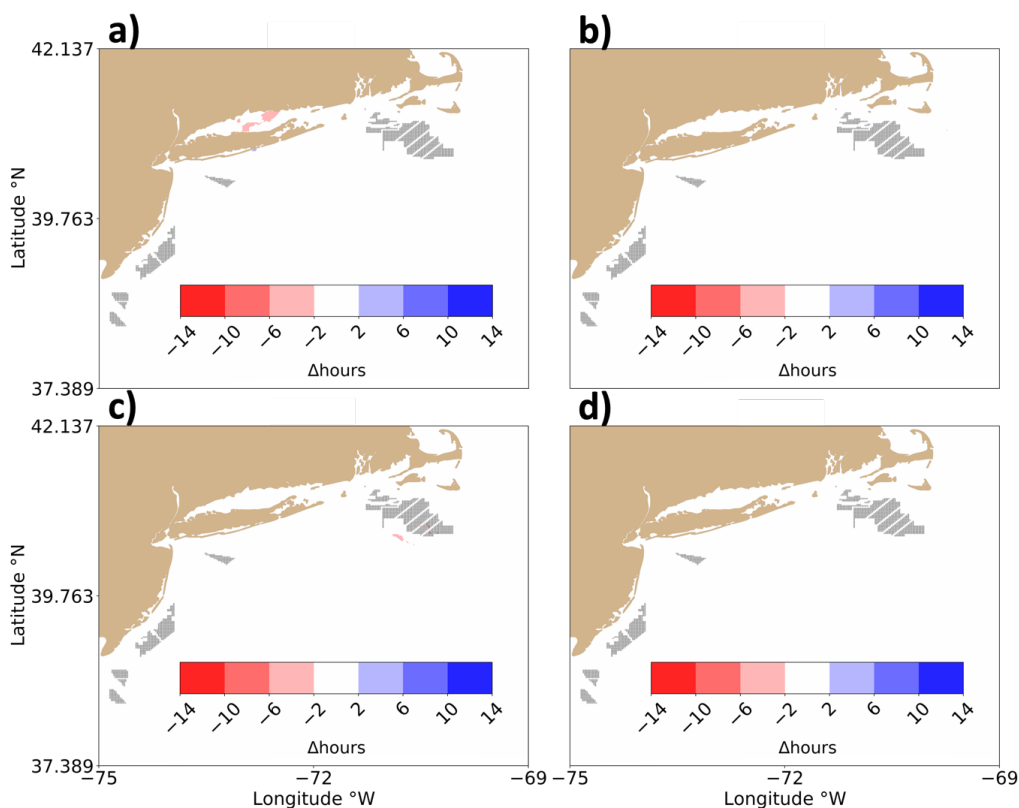
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Figure A4. The (WFP-NWF) difference in freezing hours at 10 m during (a) December 2019, (b) January 2020, (c) February 2020, and (d) March 2020. Blue (red) contouring indicates higher (lower) percentages. Gray dots indicate turbine locations.



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Figure A5. The (WFP_0-NWF) difference in freezing hours at 20 m during (a) December 2019, (b) January 2020, (c) February 2020, and (d) March 2020. Blue (red) contouring indicates higher (lower) percentages. Gray dots indicate turbine locations.



488

489 **Figure A6.** The (WFP-NWF) difference in freezing hours at the hub height during (a) December 2019, (b) January 2020,
490 (c) February 2020, and (d) March 2020. Blue (red) contouring indicates higher (lower) percentages. Gray dots indicate
491 turbine locations.

492

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