Linking weather patterns to observed and modelled turbine hub-height winds offshore U.S. West Coast

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Abstract

The U.S. West Coast holds great potential for wind power generation, although its potential varies due to complex coastal processes. Characterizing and modelling turbine hub-height winds under different weather conditions are vital for wind resources assessment and management. This study uses a two-stage machine learning algorithm to identify five large-scale meteorological patterns (LSMPs): post-trough, post-ridge, pre-ridge, pre-trough, and California-high. The LSMPs are linked to offshore wind patterns, specifically at lidar buoy locations within lease areas for future wind farm development off Humboldt and Morro Bay. Distinct wind speed, wind direction, diurnal variation, and jet feature responses are observed for each LSMP and at both lidar locations. The wind speed at Humboldt is higher during the post-trough, pre-ridge, and California-high LSMPs and lower during the remaining LSMPs. Morro Bay has smaller responses in mean speeds, showing increased wind speed during the post-trough and California-high LSMPs. Besides the LSMPs, local factors, including the land-sea thermal contrast and topography, also modify mean winds and diurnal variation. The High-Resolution Rapid Refresh model analysis does a good job of capturing the mean and variation at Humboldt but produces large biases at Morro Bay, particularly during the pre-ridge and California-high LSMPs. The findings are anticipated to guide the selection of cases for studying the influence of specific large-scale and local factors on California offshore winds and to contribute to refining numerical weather prediction models, thereby enhancing the efficiency and reliability of offshore wind energy production.

1 Introduction

With over 6000 megawatts (MW) of potential offshore wind-generating capacity in the development and operational pipeline, the U.S. West Coast is next in line following the successes of offshore wind deployment along the U.S. Atlantic Coast (Musial et al., 2023). The growth of wind power generation increases the dependence of the power system on variable weather and climate (Meenal et al., 2022). As wind energy sources are highly intermittent and variable, accurate weather forecasts are

30   essential to mitigate the related uncertainties (Frías-Paredes et al., 2017), improving decision-making and reducing cost (Turner et al., 2022; Jeon et al., 2022).

    The wind energy sector has greatly benefited from the use of numerical weather prediction. The High-Resolution Rapid Refresh (HRRR) model, operational at NOAA/NCEP since 2014, is a convection-permitting implementation of the Advanced Research version of the Weather Research and Forecasting (WRF-ARW) model with hourly data assimilation (Dowell et al.,

35   2022). The HRRR model has been proven to provide skillful forecasts of near-surface winds, leading to potential cost savings of \$14.3–\$46.6 million yearly when more advanced model configurations were applied (Fovell and Gallagher, 2022; Turner et al., 2022; Jeon et al., 2022). Despite its overall promise, the HRRR model's capabilities vary across locations and under different weather conditions (e.g., Pichugina et al., 2020). Liu et al (2024) found that HRRR tends to overestimate the turbine hub-height wind speed over complex terrain in the southeastern U.S., while Pichugina et al. (2019) reported that the HRRR

40   model underestimated strong winds speeds (>12 m s^{-1}). Most of the evaluations focused on onshore wind energy applications due to general lack of offshore high-quality wind profile measurements (Banta et al., 2013; Shaw et al., 2019; Wilczak et al., 2019). However, Banta et al. (2017) evaluated HRRR model wind forecasts against offshore Doppler lidar measurements along the U.S. Atlantic Coast. They found an average of 1.5 to 2 m s^{-1} model errors at 100 to 500 m above mean sea level (MSL). Off the California coast, there were five wind energy lease areas as of October 2023 – two off Humboldt County and three off

45   Morro Bay (BOEM, 2023). The U.S. Department of Energy funded the installation of two research buoys in these areas, equipped with lidar and other instruments to collect wind measurements for resource assessment and model evaluation (Krishnamurthy et al., 2023).

    Both onshore and offshore evaluation suggest that mean wind speed and the model bias are sensitive to weather conditions (James et al., 2018, 2017). The wind speed and/or model biases tend to be larger in winter than summer in the contiguous U.S.,

50   Pacific Northwest, Great Plains, southeastern U.S., and the U.S. Atlantic offshore (Berg et al., 2021; James et al., 2018, 2017); however, an opposite trend is observed along the California offshore coast (Liu et al., 2024; Krishnamurthy et al., 2023). In addition to large-scale weather patterns, offshore wind profiles and energy production are influenced by local factors such as frontal passages, sea breezes, and low-level jets (LLJs) (Kalverla et al., 2019; Liu et al., 2024; Sheridan et al., 2024; Gaudet et al., 2022). Specifically, sea-breeze circulation entails diurnal variations in wind speed due to thermal contrast between land

55   and sea (e.g., Gilliam et al., 2004; Burk and Thompson, 1996). During summer, this thermal contrast can cause diurnal variations in wind speed via thermal wind effect, without significant changes in wind direction (e.g., Liu et al., 2024). The presence of the North Pacific High (NPH) system and the interaction with thermal wind forcing, shallow marine boundary layer (MBL), and local topography often leads to a maximum wind speed at the top of the MBL (near turbine height), resulting in the formation of LLJs (Burk and Thompson, 1996).

60   While local factors can have a pronounced impact on near-surface winds, model biases during a period characterized by multiple weather conditions can mask local factors and ultimately lead to overlooking their impacts (Ohba et al., 2016; Liu et al., 2022; Spassiani and Mason, 2021). For instance, an approaching large-scale trough and ridge induces southerly and northerly winds, respectively. Averaging over these two periods cancels individual effects. In this study, first, we use a two-

stage clustering method to identify the predominant large-scale meteorological patterns (LSMPs) influencing the California
65   offshore. Then, we characterize the wind resources under each LSMP before evaluating the HRRR model's simulated winds
under these LSMPs.

2 Data and method

2.1 Lidar buoy data

The U.S. Department of Energy, in collaboration with the Bureau of Ocean Energy Management, placed two buoys equipped
70   with Doppler lidars along the California coastline to directly observe the offshore wind resource. These buoys were positioned
in the wind energy lease areas off the coasts of Humboldt and Morro Bay (Krishnamurthy et al., 2023). Over a full year, the
buoys gathered data on wind patterns and turbulence from 40 to 240 m above MSL, surface meteorology, sea surface
temperature, solar radiation, two-dimensional wave spectra, and ocean current profiles including speed and direction.

The lidar on the Humboldt buoy temporarily required servicing due to a power system failure, and as a result, its
75   observations are only available from October to December 2020 and from June to December 2021. In contrast, the Morro Bay
buoys operated consistently throughout these periods (i.e., from October 2020 to December 2021). Any impact on the accuracy
of lidar measurements caused by precipitation events and foggy conditions remains a subject of ongoing research. Upon
detailed examination of the carrier to noise ratio and horizontal wind speed depicted in the time-height plots during the analyzed
periods, no consistent issues with the observations were identified.

2.2 HRRR analysis

The HRRR dataset, which is accessible through Amazon Web Services, is available in both hybrid and pressure vertical
coordinates with a 3-km horizontal grid spacing (https://registry.opendata.aws/noaa-hrrr-pds/). We obtain the wind speed and
direction from HRRR at 80 m above MSL from October 2020 to December 2021. The hourly wind components are interpolated
horizontally, aligning with the locations of observation sites.

2.3 Meteorological variables describing weather patterns

Atmospheric patterns over the North Pacific influence coastal winds, but the HRRR model's western boundary is too close to
the coastline to effectively capture these circulations. To address this, the European Centre for Medium-Range Weather
Forecasting 5th Generation Reanalysis (ERA5, Hersbach et al., 2020), with its global coverage, is used to perform the weather
pattern classification. The hourly variables for the period of 2019-2022 at horizontal resolution of 0.25°×0.25° are obtained
90   from https://www.ecmwf.int/en/forecasts/dataset/ecmwf-reanalysis-v5 (last accessed January 2024). The 500-hPa
geopotential height (Z_{500}) and surface pressure (P_{sfc}) are obtained to describe the large-scale pressure gradient, and the 2-m
temperature (T_2) is used for land-sea thermal contrast.

2.4 Large-scale meteorological weather pattern clustering

The two-stage clustering method consists of a self-organizing maps (SOMs, Vesanto and Alhoniemi, 2000; Kohonen, 1982)
analysis to reduce the dimension of the input vectors (Z_{500}, P_{sfc}, and T_2), and a K-means cluster to further group the SOM
prototypes into fewer LSMPs (Liu et al., 2023). In the first stage, we train a SOM to generate a low-dimensional discretized
representation of the data in the original feature space while preserving the topological properties (relative position between
the SOM nodes) of the data. In the second stage, we use the SOM prototypes as input to train the K-means method for final
clustering. The SOM is a widely used clustering analysis tool (Ohba et al., 2016; Huang et al., 2022; Liu et al., 2022) that
performs a topology-preserving mapping. However, when the number of nodes is small, SOM tends to cluster the input vectors
into symmetrically paired nodes, whereas those from K-means are independent of each other (Liu et al., 2023). Directly using
K-means for clustering is not recommended, as K-means is highly sensitive to the positions of the initial nodes and outliers
and is not suitable for high-dimensional datasets (Mingoti and Lima, 2006; Misra et al., 2020). As a result, this two-stage
procedure approach combines the strengths of both SOMs and K-means while addressing their individual shortcomings.

The numbers of prototypes and LSMPs are prescribed depending on the scale of meteorological patterns. We choose a
large map size, 10×10 SOM prototypes, which is sufficient to represent all possible mesoscale variations (on the order of 100
km) such as sea breezes, squall lines, and mesoscale convective complexes. Then, the silhouette score (SS) is used to determine
the number of LSMPs (on the order of 1000 km) in the K-means analysis. The SS measures the separation distance between
the resulting clusters: A larger SS indicates larger distinctions among the clusters (Shutaywi and Kachouie, 2021). We test 3
to 14 clusters and find that 5-cluster classification produces the largest SS (figure not shown), which therefore is chosen to
perform the K-means clustering analysis. Before performing the two-stage procedure, we calculate the anomalies of input
vectors by subtracting the climatological hourly mean from the timeseries at each grid point so that the annual and monthly
variations are excluded. The SOM analysis is performed over 30°N–45°N, 130°W–118°W, which is chosen to include the
Pacific jet exit (Athanasiadis et al., 2010).

2.5 Low-level jet identification

An LLJ is typically a local maximum in wind speed at altitude from near the surface up to about 2 km, yet there is no standard
method for quantifying LLJs. The "fall-off" method, commonly used to determine the height of peak wind speeds (the height
of the jet core), involves identifying where wind speed decreases after reaching its maximum (the jet core speed). An LLJ is
recognized if the difference between the jet core speed and the next local minimum above it exceeds a certain threshold, which
varies among studies (Carroll et al., 2019; Kalverla et al., 2019; Hallgren et al., 2020). Following Sheridan et al. (2024), a
study that comprehensively evaluated West Coast LLJs using the same observational dataset, this study uses a 2 m s⁻¹ fall-off
threshold to define LLJs, without specifying the vertical distance between the jet core and the threshold height as long as it is
within the observational limit of 240 m above MSL. We only calculate LLJ from observations since HRRR provides too few
near-surface points in the vertical direction for LLJ detection.

3 Results

Before exploring the LSMPs, we review the meteorological systems influencing near-surface winds offshore of California. The strength and location of the NPH system is the major contributor to the LSMP (Burk and Thompson, 1996). In summer, the NPH and a thermal low over the southwest U.S. create an enhanced cross-coastline pressure gradient, which primarily drives the prevailing northerly winds offshore (Brewer et al., 2012). The subsidence within the NPH induces a pronounced air temperature inversion that is most intense and lowest near the coast, capping the MBL and limiting its vertical extent. Along the coastline, a northerly to northwesterly LLJ is typically observed at the top of the MBL, resulting from the thermal wind due to significant coastal baroclinity superimposed on the generally northerly flow (Burk and Thompson, 1996). During the day, the land is typically warmer than the ocean. Above the MBL, the thermal wind is southerly aloft, and so the geostrophic winds become more northerly closer to the surface, increasing northerly wind speed until surface friction slows it down within the MBL, at around 500 m (e.g., Liu et al., 2024). In contrast, during winter, the NPH weakens, leading to distinct synoptic-scale weather conditions including storms and fronts originating from the Gulf of Alaska propagating southward through the California region. The propagation of these systems can drive strong winds and a wind direction shift from northwest to southeast. The MBL also deepens during this season, influenced by the changing dynamics of the NPH.

3.1 SOM prototypes

In the first stage clustering, 10×10 SOM prototypes resemble the LSMPs modified by mesoscale perturbations (Fig. 1). Viewing Fig. 1 from left to right, a 500-hPa high moves from west to east, while from top to bottom, a 500-hPa high moves from north to south. As these systems move eastward, the high and low centres rotate clockwise. Changes in the mid-level atmosphere correspond to surface alterations. The P_{sfc} generally resembles the similar patterns of Z_{500}, leading to a range of P_{sfc} gradients that drive variable surface winds (figure not shown). Mid-level high pressure, often associated with subsidence, inhibits cloud formation, thereby increasing surface solar radiation and temperature (Dadashazar et al., 2020). Conversely, mid-level low pressure can induce opposite changes. Meanwhile, land exhibits larger temperature variations than ocean due to its smaller heat capacity. As a result, positive and negative T_2 anomalies are found over the land area underlying respective positive and negative Z_{500} anomalies.
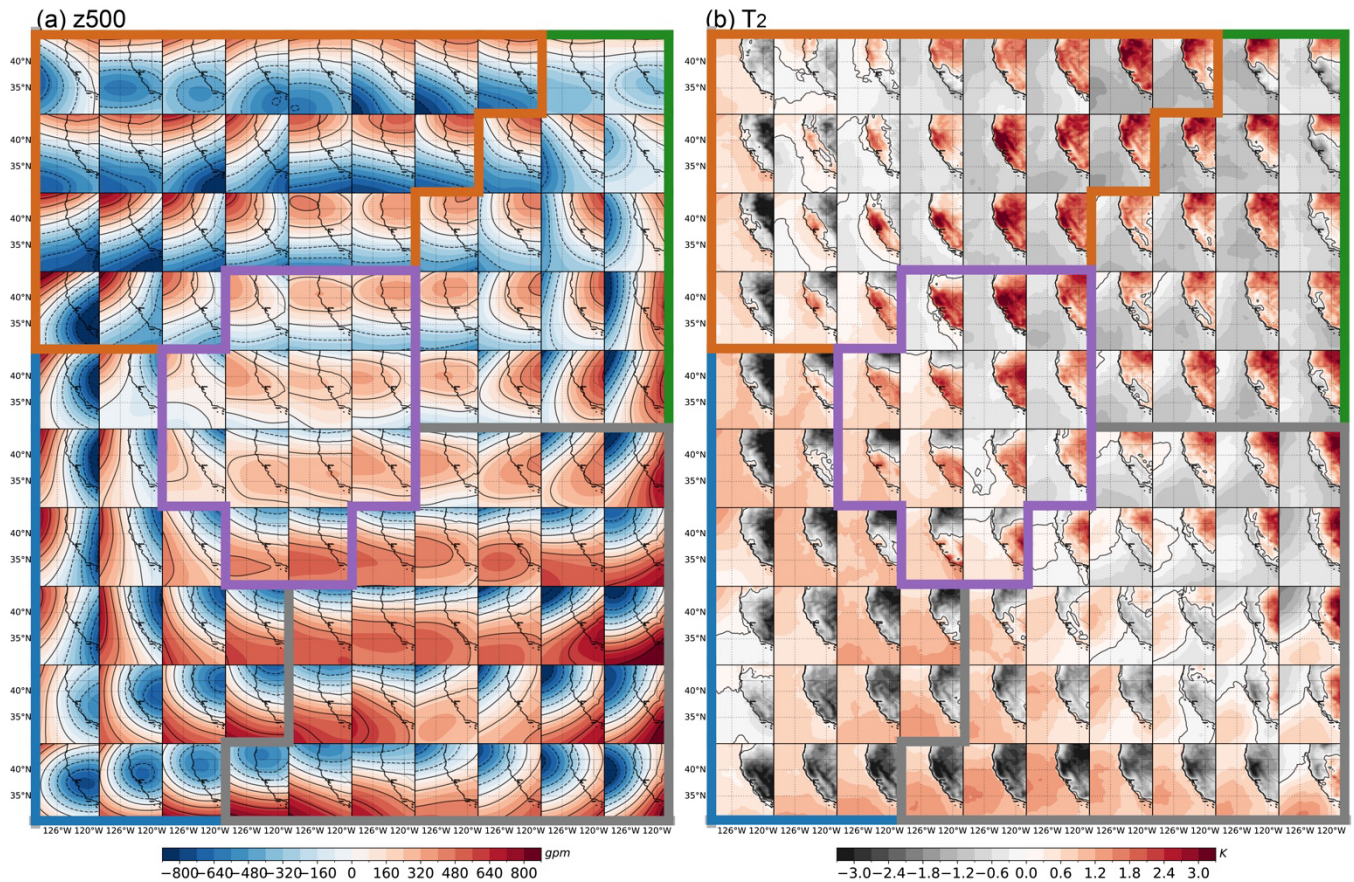
Figure 1: Composite anomalies of Z_{500} and T_2 for each SOM prototype. The red shading and solid contours indicate positive anomalies, and blue and dashed contours indicate negative anomalies. The coloured lines outline the five LSMPs: post-trough (blue), post-ridge (green), pre-ridge (orange), pre-trough (grey), and California-high (purple).

3.2 Dominant large-scale meteorological weather patterns and associated wind patterns

In the second stage clustering, the K-mean analysis produces five LSMPs. The first two LSMPs, each accounting for 17% of the total hours during the study period, resemble large-scale ridges and troughs centred over the western U.S., with the buoys located behind (west) of those systems (Fig. 2a, b). For the post-trough weather pattern, the passing of the mid-level trough and the following ridge of high pressure intensifies the cross-coastline pressure gradient, which enhances the northerly winds offshore. Although the cold northerly winds cool the land surface, the land-sea thermal contrast still exists and further accelerates the northerly winds through the thermal wind effect. Consequently, the post-trough LSMP drives strong offshore winds along the coast, with enhanced expansion fans (an area of high wind speed) downwind of Cape Mendocino and Point Conception (Fig. 3a).

The second LSMP, namely post-ridge, is associated with blocking systems over the western U.S.; this is an elongated area of relatively high atmospheric pressure. The high-pressure systems are associated with subsiding air, which discourages cloud

formation and leads to stable weather conditions. The inland surface temperature can be higher due to the increased solar
165   radiation and the general downward motion of air, which warms adiabatically as it descends. Under post-ridge conditions, the
weakened cross-coastline pressure gradient decreases the prevailing surface northerly wind (Fig. 3b). In contrast, the enhanced
land-sea thermal contrast along the Oregon coast tends to increase offshore winds through the thermal wind theory (Liu et al.,
2024), which mitigates the overall decrease in offshore wind speed.

The third and fourth LSMPs resemble large-scale ridges and troughs centred over the Pacific Ocean, and the buoys are
170   located ahead (east) of those systems (Fig. 2c, d), accounting for 26% and 28% of the total hours, respectively. Like the post-
ridge pattern, the pre-ridge pattern is associated with high-pressure systems favourable to subsidence, warm land surface, and
decreased offshore winds (Fig. 3c). In contrast, the pre-ridge pattern features an anomalous north-to-south pressure gradient at
500 hPa and a strong surface pressure gradient centred offshore of northern California and Oregon, accelerating wind speed.
Meanwhile, the warmer land surface over the northern California mountains further increases the offshore winds. As a result,
175   the pre-ridge increases (decreases) wind speed off northern (southern) California.

For the fourth LSMP, namely pre-trough, the buoys are located ahead (east) of a mid-level trough, where the dynamics can
create conditions favourable for a cold front and strong convection. The enhanced NPH over California and the Aleutian Low
intensify the pressure gradient at 500 hPa and the surface, which often is associated with a cold front approaching California
from the northeast Pacific Ocean. As a cold front approaches, the tightened pressure gradient can lead to an increase in wind
180   speeds and change the wind direction to south or southwest (Fig. 3d). This occurs more frequently offshore of northern
California and thus has less impact on the wind speed in the south.

The fifth LSMP, namely California-high, accounts for 11.4% of the hours and exhibits an anomalous high at 500 hPa
centred offshore of California (Fig. 2e). The mid-level and surface pressure and temperature patterns are similar to the pre-
ridge LSMP, except they have larger magnitudes. Therefore, the area with accelerated winds extends from northern California
185   to the south (Fig. 3e).
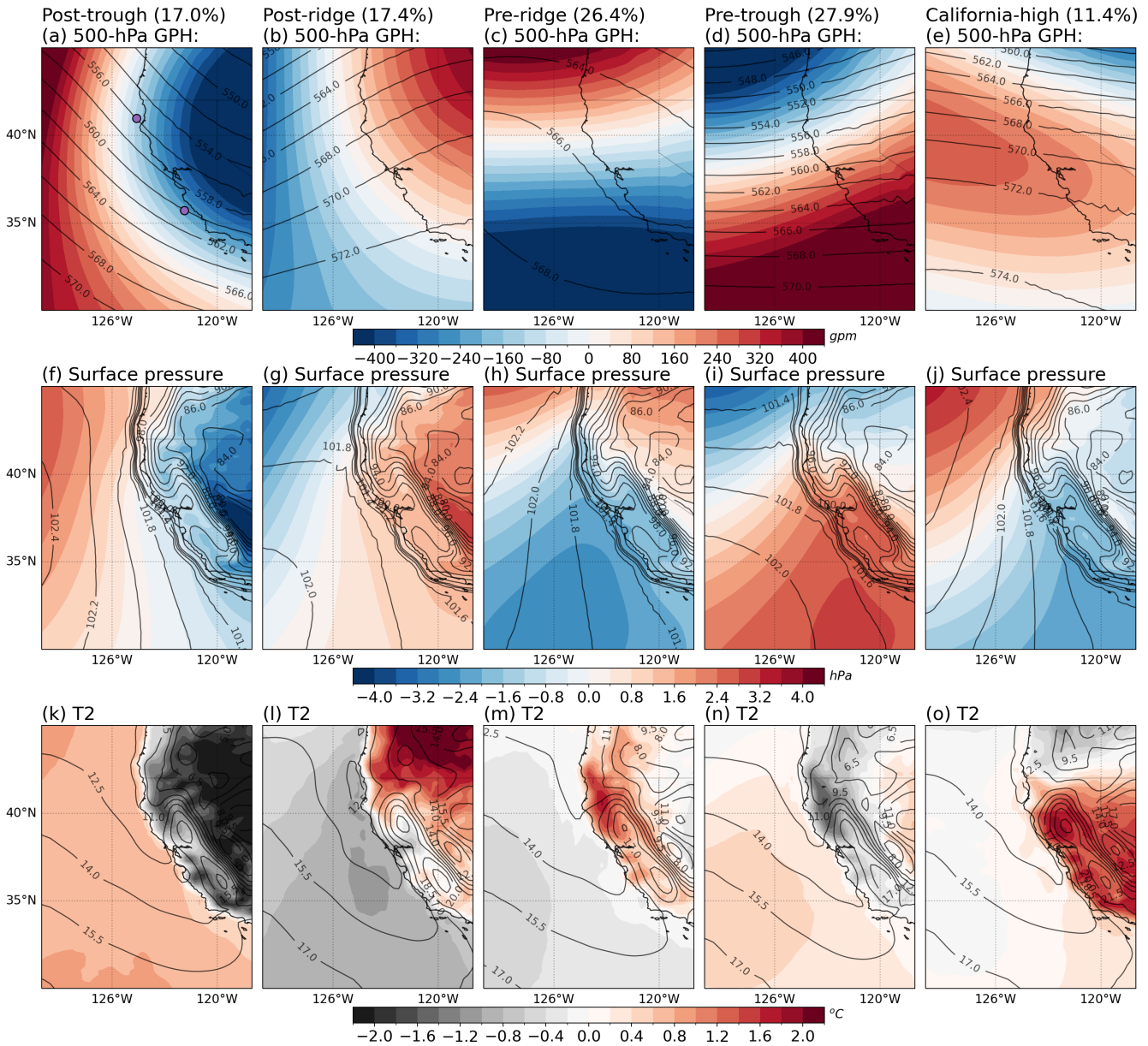
Figure 2: Spatial distributions of mean (contour) and anomalies (shading) of 500 hPa GPH (a-e), surface pressure (f-j), and T2m (k-o). The value at the top of each column indicates the frequency of each LSMP. The purple dots on panel (a) indicate the lidar buoy locations at Humboldt in the north and Morro Bay in the south.
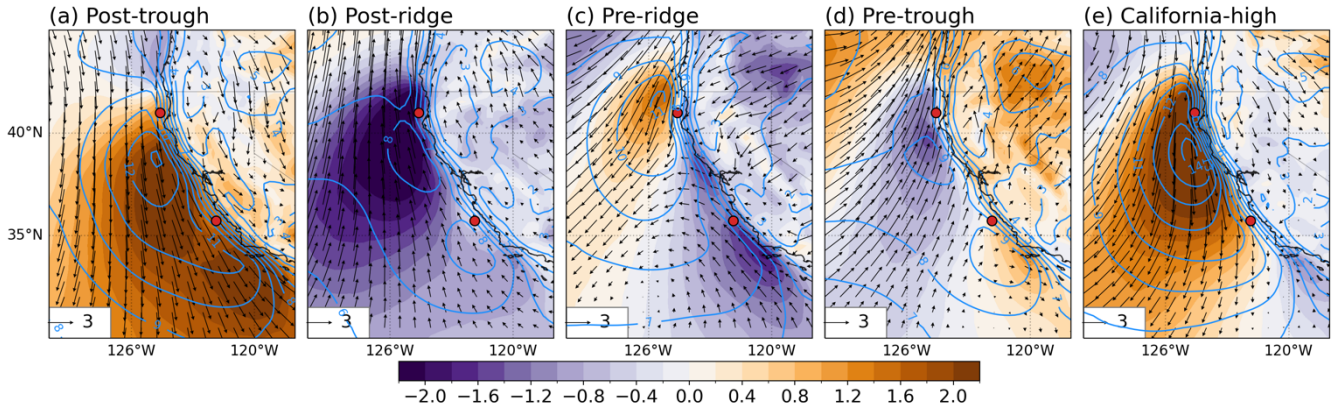
Figure 3: Spatial distributions of mean (contour) and anomalies (shading and vector) of 80-m wind speed and direction from HRRR. The red dots on panel (a) indicate the lidar buoy locations at Humboldt in the north and Morro Bay in the south.

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Table 1: Summary of LSMPs and the associated wind patterns

LSMP	Frequency (%)	Weather pattern	Wind pattern
Post-trough	17.0	500-hPa trough centred over western U.S land. Intensified cross-coastline surface pressure gradient. Weakened land-sea thermal contrast.	Strong northerly to northwesterly winds with expansion fans downstream the capes.
Post-ridge	17.4	500-hPa ridge centred over western U.S. land. Weakened cross-coastline surface pressure gradient. Enhanced land-sea thermal contrast.	Decreased overall wind speeds offshore of north California.
Pre-ridge	26.4	500-hPa ridge centred over north Pacific Ocean. Anomalous north-to-south surface pressure gradient, with enhanced cross-coastline surface pressure gradient offshore of north California, and decreased pressure gradient in the south. Enhanced land-sea thermal contrast.	Increased northerlies offshore of north California, and decreased wind speed in the south.
Pre-trough	27.9	500-hPa trough centred over north Pacific Ocean. Anomalous south-to-north surface pressure gradient, with decreased cross-coastline surface pressure gradient offshore of north California, and decreased pressure gradient in the north. Decreased land-sea thermal contrast.	Reduced northerlies offshore of north California, and increased wind speed in the south.
California-high	11.4	500-hPa high centred over California. Enhanced cross-coastline surface pressure gradient. Increased land-sea thermal contrast.	Strong northerly winds offshore of central and north California.

3.3 80-m wind regimes at buoy locations

Wind patterns at Humboldt and Morro Bay exhibit distinct responses to LSMPs and smaller-scale atmospheric disturbances as recognized by the SOM prototypes. Figure 4 presents the average 80-m wind speed for each SOM prototype. Post-trough

200   conditions reveal stronger northerly to northwesterly winds along the coast. The Humboldt buoy, situated at the boundary of the northern expansion fan, experiences a pronounced horizontal wind speed gradient (Fig. 3a), leading to a wide range of mean wind speeds from 1 to 14 m s^{-1}. In contrast, the Morro Bay buoy consistently records high wind speeds.

Under post-ridge conditions, anomalous southerlies are most intense offshore of northern California, significantly reducing wind speeds at Humboldt and potentially causing wind direction change when the southerly anomaly is greater than the

205   prevailing north wind. When the mid-level ridge intensifies, these offshore winds can become southerlies with speeds exceeding 15 m s^{-1} in extreme cases. At Morro Bay, wind speeds generally weaken to 6–11 m s^{-1}.

The pre-ridge LSMP results in increased mean wind speeds off northern California and decreased speeds to the south. This wind pattern is manifested by anomalies in surface temperature. The dominance of NPH-induced clear skies leads to warmer land temperatures that enhance the land-sea thermal contrast, thereby accelerating offshore winds. Across the SOM prototypes,

210   the wind speed at Humboldt range of 6–15 m s^{-1}, while the Morro Bay buoy records a range of 4–10 m s^{-1}.

Pre-trough conditions, marked by an approaching mid-level trough, generate moderate to high winds (7–15 m s^{-1}) at Humboldt, with wind directions shifting from northwest to southwest during surface frontal passages. The impact of the mid-level trough on the Morro Bay buoy is minimal, with prevailing northwesterlies under the influence of the NPH. However, if the trough deepens, wind speeds at Morro Bay may transition from weaker northwesterlies to stronger southwesterlies.

215   During California-high conditions, both Humboldt and Morro Bay record increased wind speeds due to an enhanced NPH and the amplified land-sea thermal contrast. The mean wind speed exceeds 12 m s^{-1} at Humboldt and 10 m s^{-1} at Morro Bay.
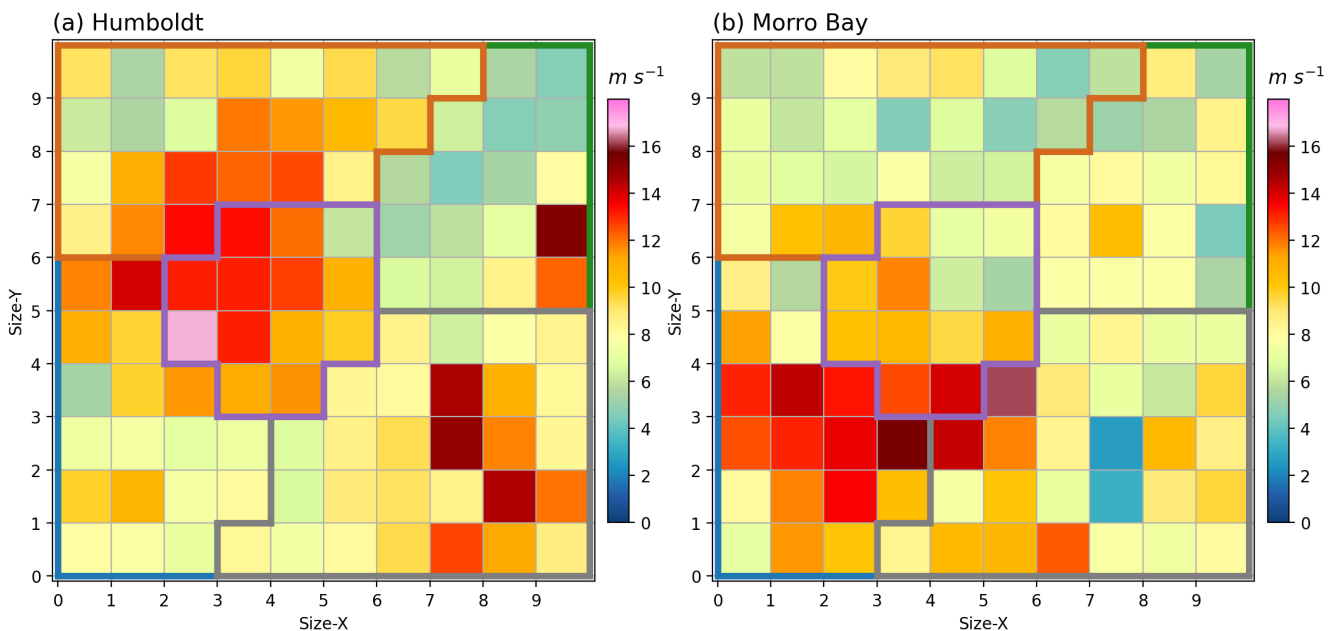
Figure 4: Mean 80-m wind speed of each SOM prototype at Humboldt (a) and Morro Bay (b). The coloured lines outlined the five LSMPs: post-trough (blue), post-ridge (green), pre-ridge (orange), pre-trough (grey), and California-high (purple).

3.4 Observed and modelled diurnal variation

The diurnal variation in surface heating alters the wind speed and direction via the thermal wind effect. The thermal wind effect results in northerly winds relative to the free troposphere during the day and southerly winds at night. The maximum speed occurs a few hours after the peak of baroclinity in the mid-afternoon as the inertial turning of the land-sea breeze circulation to the free troposphere wind direction through the thermal wind effect (Burk and Thompson, 1996).

At Humboldt, the prevailing free tropospheric wind direction varies across the LSMPs, showing a northerly trend during the post-trough, pre-ridge, and California-high conditions and a southwesterly shift during the post-ridge conditions (Fig. 5). The maximum speeds are observed near midnight (09–12 UTC) under northerly free troposphere wind conditions and in the afternoon (00 UTC) during southwesterly conditions. Large-scale cold fronts primarily influence the winds during pre-trough conditions, resulting in small diurnal variations. In contrast, at Morro Bay, the maximum speed is observed from evening to midnight (00–06 UTC) due to the prevailing northerly to northwesterly winds throughout all LSMPs.

The diurnal variations in 80-m wind speed modelled by HRRR are compared with observations (Fig. 5). HRRR does a good job of capturing the mean wind speed at Humboldt, with biases ranging from -0.5 to 0.1 m s^{-1}. The HRRR effectively captures diurnal variations during post-trough and post-ridge LSMPs but underestimates the variation during pre-ridge conditions. At Morro Bay, the HRRR produces small biases during post-trough, post-ridge, and pre-trough conditions, while it largely overestimates the daily mean during pre-ridge LSMP by 2.2 m s^{-1} and California-high LSMP by 2.6 m s^{-1}. The overestimation may be connected to frequent LLJs occurring during these two LSMPs, which results in a large vertical wind speed gradient. Liu et al. (2024) also reported that an overpredicted land-sea thermal contrast can lead to an overestimation of wind speed. Generally, HRRR tends to reproduce the diurnal variations with a slight delay during most LSMPs except for pre-trough, where the model shows no diurnal changes.
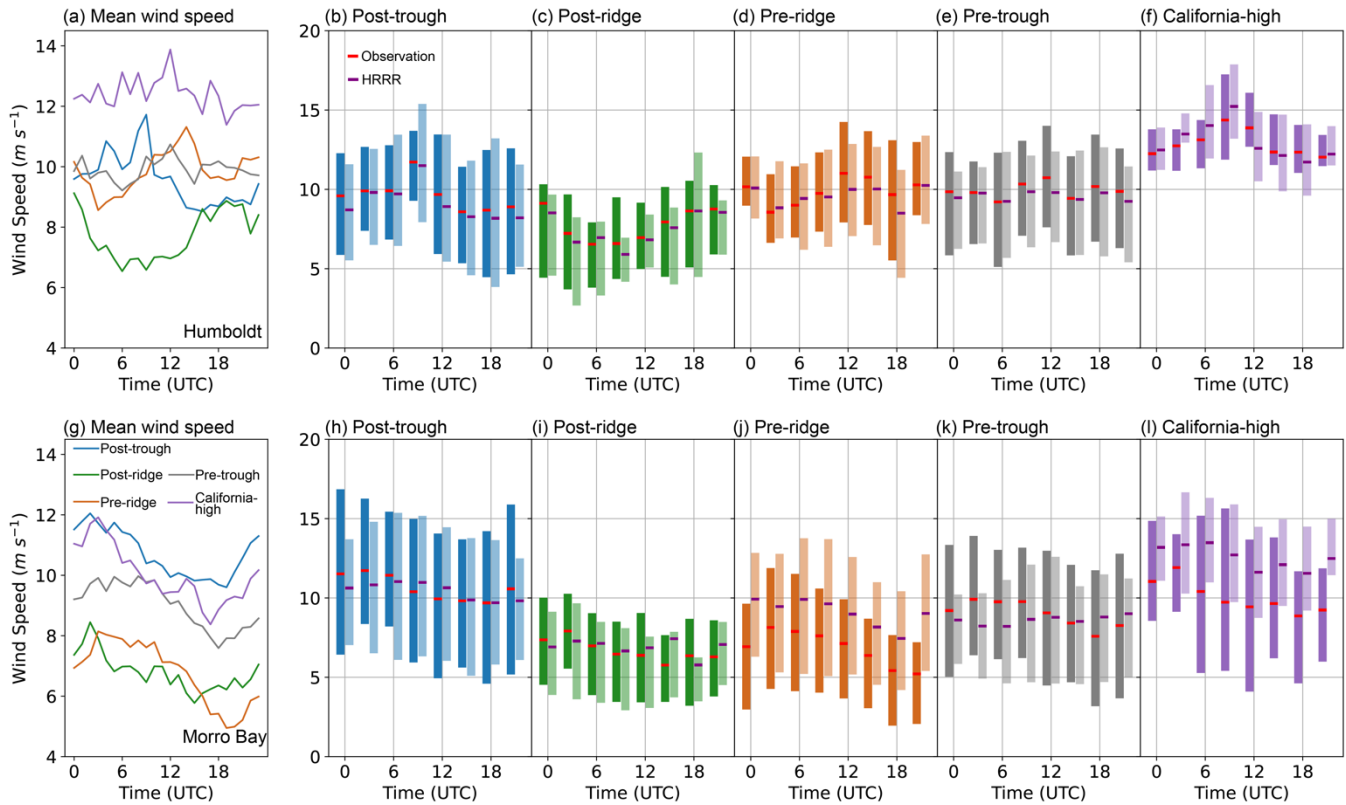
Figure 5: Diurnal variations of wind speed at Humboldt (a-f) and Morro Bay (g-l). Observed mean wind speed associated with each LSMP (a, g) and box plot comparison for individual LSMPs (b-f) and (h-l). The line in the centre of each box indicates the mean value and the extends of the box indicate the first quartile (Q1) to the third quartile (Q3) of the data.

3.5 Observed low-level jet

245   The LLJ offshore of California is often characterized by a strong vertical wind speed gradient, which can introduce significant biases in modelled wind speed at 80 m (Liu et al., 2024). We examine the occurrence, jet-core height, and jet-core wind speed across various LSMPs (Fig. 6). Out of all the lidar observations, 1107 (4%) and 1911 (4%) record LLJ occurrences at Humboldt and Morro Bay, respectively, consistent with the findings by Sheridan et al. (2024) using the same lidar dataset. At Humboldt, the pre-ridge LSMP has the highest LLJ occurrence (406 LLJs), with the occurrence of jet-core height peaking at 160 m. The

250   pre-ridge and pre-trough LSMPs record 309 and 238 LLJs, respectively, with a bimodal distribution of occurrence of jet-core height peaking near 140 and 220 m for both LSMPs. Interestingly, the post-trough and California-high LSMPs record few LLJs despite being associated with high-speed winds. This is likely because high-speed winds consistently appear throughout the lidar measuring range of 40–240 m without forming an identifiable jet structure within this height range.

          The mean jet-core wind speed at Humboldt varies between 9.7 and 11.4 m s^{-1} across all LSMPs and generally increases

255   to its height with a maximum at 200 m (figure not shown). Larger variations appear associated with individual LSMPs. The

California-high LSMP has the largest mean jet-core wind speed of 18.3 m s^{-1} at 220 m. For other LSMPs, the maximum mean jet-core speed ranges from 13.9 to 16.5 m s^{-1} at various heights.

Similar jet characteristics are observed at Morro Bay during most LSMPs, except for the California-high, which records the most LLJ occurrences (663 LLJs) at Morro Bay. The discrepancy between the two locations during the California-high
260 LSMP may be connected to the typically lower MBL at Morro Bay compared to Humboldt (Zhou et al., 2020), resulting in lower jet cores peaking at 160 m, within the lidar range. The maximum mean jet-core wind speed varies from 12.5 to 17.4 m s^{-1}. Despite the opposite sign anomalous 100-m wind speed at two locations during the pre-ridge and pre-trough LSMPs (Fig. 3c, d), minor differences are observed in jet characteristics. This supports previous literature reporting that the LLJ is a mesoscale phenomenon modified by local meteorology and topography.

265 It is important to note that the lidar's maximum measurement height limitation (220 m) likely results in an underrepresentation of LLJ occurrence at heights above 220 m. The consistent increase in mean wind speed with height suggests potential jet cores above the highest measurement. To the best of our knowledge, long-term LLJ measurements do not exist in this region. Therefore, it is difficult to know the true frequency of LLJ conditions. Nonetheless, we anticipate that the limited range of the lidar contributes to an underestimation of LLJ frequency, which reanalysis and global climate models
270 estimate to be ~20-30% annually off the California coastline (e.g., Lima et al., 2022, Juliano et al., 2024).
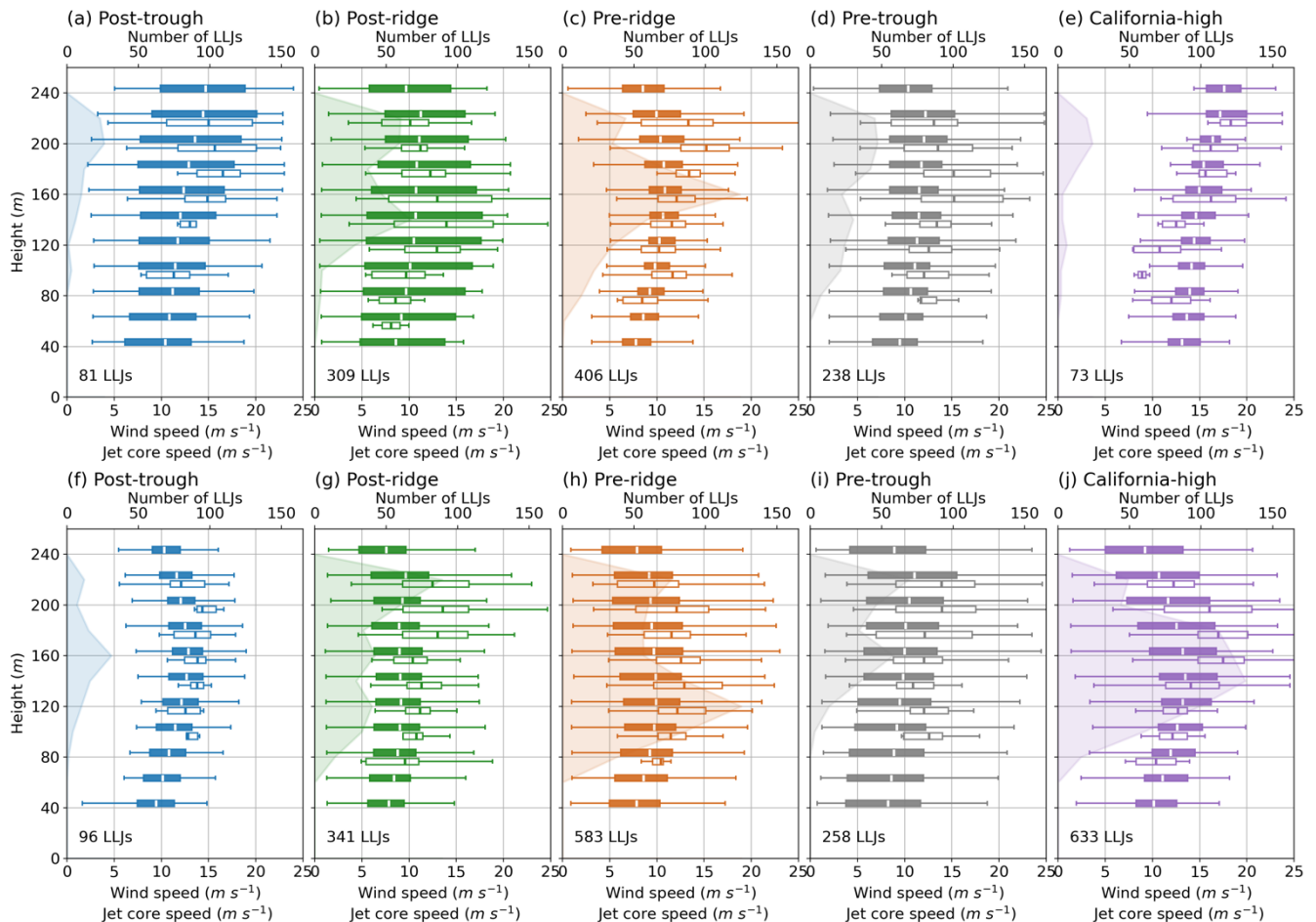
Figure 6: Observed wind speed profile (filled box), occurrence of jet-core height (shading), and jet-core wind speed (open box) at Humboldt (a-e) and Morro Bay (f-j) for each LSMP. The value in the lower-left indicates the number of LLJ events observed during each LSMP. The box extends from the first quartile (Q1) to the third quartile (Q3) of the data, with a dash line at the mean. The whiskers extend from Q1-1.5× (Q3-Q1) to Q3+1.5×(Q3-Q1).

4 Conclusion

In this study, we use a two-stage clustering method to identify the LSMPs influencing near-surface winds off the California coast. The 10×10 SOM prototypes resemble high-pressure systems moving eastward and southward (Fig. 1), driving variations in wind speed and direction at the Humboldt and Morro Bay buoys. The SOM prototypes are aggregated using the K-means method into five LSMPs: post-trough, post-ridge, pre-ridge, pre-trough, and California-high (Fig. 2). The post-trough and post-ridge LSMPs resemble west-to-east mid-level and surface anomalies, enhancing and reducing the cross-coastline pressure gradients, respectively, leading to accelerated and deaccelerated offshore winds (Fig. 3). The pre-ridge and pre-trough LSMPs resemble north-to-south mid-level and surface anomalies with intensified anomalous surface pressure gradient offshore of north California and Oregon. The pre-ridge LSMP increases wind speed at Humboldt and slightly decreases wind speed at

Morro Bay. Opposite changes are observed in mean wind speed at the two locations associated with pre-trough LSMP. However, a strong surface front can drive relatively strong southwesterly winds that are larger in magnitude than the prevailing northerly winds.

290   The offshore near-surface winds are modified by the diurnally varying land-sea thermal contrast (Fig. 5). The maximum wind speeds occur a few hours after peak baroclinity in the afternoon and vary with the prevailing free tropospheric wind direction, which is influenced by different LSMPs. The HRRR model generally captures these variations well at Humboldt, with minor biases, but it overestimates mean wind speeds at Morro Bay during pre-ridge and California-high LSMPs, possibly due to frequent LLJ occurrences and an overpredicted land-sea thermal contrast. HRRR also tends to show a slight delay in reproducing diurnal variations and does not reflect changes during pre-trough conditions.

295   At Humboldt, the highest LLJ occurrence is during the pre-ridge LSMP, with jet-core heights peaking at 160 m, while at Morro Bay, the California-high LSMP records the highest number of LLJs (Fig. 6). The mean jet-core wind speed at Humboldt ranges from 9.7 to 11.4 m s^{-1}, increasing with height, and the California-high LSMP shows the highest mean speeds up to 18.3 m s^{-1}. Despite some differences in wind speeds and LLJ characteristics between Humboldt and Morro Bay, the general consistency between the two locations (800 km apart) suggests the LLJ is a meso-alpha scale phenomenon modified by local conditions.

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This study introduces a new approach to characterizing offshore winds and associated model biases, linking them to LSMPs. The results are anticipated to guide the selection of cases for studying the influence of specific large-scale and local factors on winds off the California coast. Furthermore, this research aids in refining numerical weather prediction models, thereby enhancing the efficiency and reliability of offshore wind energy production.

305   **Code and Data availability**

The lidar buoy data utilized in this study are publicly available from the U.S. Department of Energy for Humboldt (https://doi.org/10.21947/1783809) and Morro Bay (https://doi.org/10.21947/1959721). ERA5 is available through the Copernicus Climate Change Service Climate Data Store at http://cds.climate.copernicus.eu. The HRRR dataset is publicly available at https://registry.opendata.aws/noaa-hrrr-pds/.

310   **Author contributions**

YL and RK conceived the idea. YL performed the research, analysed the data, and prepared the manuscript with contributions from all co-authors. All authors contributed to the paper edits and technical review.

Competing interests

The contact author has declared that none of the authors has any competing interests.

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