Hurricane impacts in the United States East Coast offshore wind energy lease areas

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Abstract. Wind turbines deployed in offshore wind energy lease areas along the U.S. East Coast could significantly contribute to the national electricity supply. This region is also impacted by powerful tropical and extra-tropical cyclones that may lead to high structural loading on wind turbines and support structures and, in the event of above cut-out wind speedsspeed, low power production (capacity factors < 0.2). Four sets of high-resolution simulations are performed for two category 3 tropical cyclones that tracked close to current offshore wind energy lease areas to assess the possible impacts on, and from, wind turbines. Simulations of Hurricanes Irene and Sandy are performed at convective permitting resolution (grid spacing in inner domain of 1.33 km) with both the Weather Research and Forecasting Modelmodel (WRF, v4.2.2) and the Coupled Ocean-Atmosphere-Wave-Sediment Transport (COAWST) Model, v3.7) model to characterize geophysical conditions of relevance to offshore wind turbines. These simulations are performed without and with a wind farm parameterization (WFP) active with the latter using the assumption that existing lease areas are fully populated with 15 MW wind turbines at a 1.85 km spacing. Many aspects (e.g., track, near-surface wind speed, sea level pressure, precipitation volumes) are well reproduced in control simulations (no WFP) with both WRF and COAWST particularly for Hurricane Sandy. COAWST simulations lead to more intense cyclones with a slightly larger area of storm-force wind speeds, a higher likelihood of hub-height wind speeds > 25 m s⁻¹, plus higher precipitation volumes, possibly indicating under-estimation of hurricane risk in uncoupled simulations. All eight simulations indicate maximum hub-height wind speeds (HHWS) within the existing lease areas below 50 m s⁻¹. However, COAWST simulations indicate frequent wind-wave misalignment of > 30° and the joint occurrence of significant wave height, hub-height wind speed, and wave period in some lease areas reach levels that are likely to be associated with large structural loads. This work re-emphasizes the utility of coupled simulations in describing geophysical conditions of relevance to offshore wind turbine operating conditions.

1 Introduction

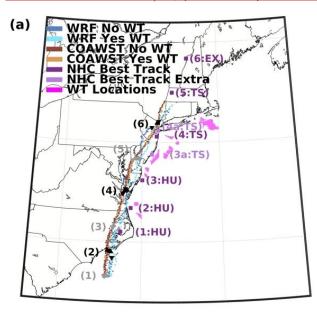
1.1 Motivation

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At the end of 2023, the global offshore wind energy installed capacity (IC) was approximately 75.2 GW (GWEC, 2024a) due in part to a 24% increase in installed capacity during 2023 (GWEC, 2024b). The plentiful offshore wind resource At the end of 2023, the global offshore wind energy installed capacity (IC) was approximately 75.2 GW (GWEC, 2024b) due in part to a 24% increase in IC during 2023 (GWEC, 2024a). The plentiful offshore wind resource (Marvel et al., 2013; Bodini et al., 2024; Pryor and Barthelmie, 2024b2024a) and recent reductions in the Levelized Cost of Energy for offshore deployments (Jansen et al., 2020; Wiser et al., 2021)—mean that the number of wind turbines (WT) deployed in (coastal) offshore regions is projected to rapidly increase (GWEC, 2024b; Pryor and Barthelmie, 2024b).

mean that the number of wind turbines (WTs) deployed in (coastal) offshore regions is projected to rapidly increase (GWEC, 2024a; Pryor and Barthelmie, 2024a). For example, substantial offshore wind energy developments are planned or in progress

along the U.S. East Coast (Fig. 1) in regions with high wind resource (power generation potential), close proximity to major demand centers, and shallow water depths (Pryor et al., 2021; Pryor and Barthelmie, 2024a, b).



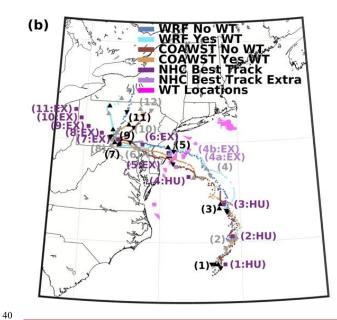


Figure 1: Tracks of Hurricane (a) Irene and (b) Sandy from the WRF and COAWST simulations derived from the minimum sea level pressure (SLP) every 10 min, along with the corresponding National Hurricane Center (NHC) best track locations every 6 h. Two letter abbreviations indicate variations in cyclone intensity/type: HU indicates hurricane stage, TS indicates tropical storm stage, and EX indicates extratropical stage. Also shown (in magenta) are the locations of the wind turbines (WTs) in the offshore lease areas (LAs) considered herein. Numbers in brackets represent the location at the 6 h NHC increments, where "1" represents 27 August 2011 18:00 UTC with Irene and 29 October 2012 00:00 UTC with Sandy; 3a (28 August 2011 09:35 UTC) and 4a (28

August 2011 13:00 UTC) represent landfall with Irene, while 4a (29 October 2012 21:00 UTC) represents a downgrade to extratropical cyclone intensity and 4b (29 October 2012 23:30 UTC) represents when Hurricane Sandy made landfall. Additional details are provided in Table S1.

The offshore environment presents significant challenges for making long-term, climatologically representative robust measurements of properties such as wind speed at wind turbine WT hub-height (HH WS) (Foody et al., 2024) that are critical for determining the wind resource and key aspects of operating conditions (IEC, 2019b, a; Mudd and Vickery, 2024). The relative paucity of measurements leads to financial uncertainty and thus potentially jeopardizes realizing U.S. national goals for achieving the energy transition (Hansen et al., 2024). It also means that numerical modeling is playing a critical role in projecting wind resource and operating conditions in offshore wind energy development areas (Kresning et al., 2020; Pryor and Barthelmie, 2021; Bodini et al., 2024; Pryor and Barthelmie, 2024a; Wang et al., 2024).

Substantial offshore Limited over-ocean observations also limit our ability to characterize the characteristics of high intensity hurricanes, including those of relevance to the wind energy developments are planned orindustry, particularly in progress alongenvironments such as the U.S. East Coast (Fig. 1) in regions with high wind resource (power generation potential), close proximity to major demand centers, and shallow water depths (Pryor et al., 2021; Pryor and Barthelmie, 2024b, a). This region alsowhich has the potential to be impacted by tropical cyclones and/or transitioning tropical-extratropical cyclones (Xie et al., 2005; Baldini et al., 2016; Barthelmie et al., 2021; Wang et al., 2024); but experiences only relatively few such storms each century (Schreck III et al., 2021).

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Wind speeds (WSs) within tropical cyclones frequently exceed the threshold at which wind turbines WTs cease power production (25-to 30 m s⁻¹) to avoid high operational loads (Petrović and Bottasso, 2014). There are reports of individual wind turbineWT failures during hurricanes (Chen and Xu, 2016), and six wind turbinesWTs in a wind farm without hurricaneresistant wind turbines WTs were damaged by 65 m s⁻¹ wind speeds WSs during Typhoon Yagi in September 2024 (Yihe, 2024). Accordingly, hurricane-induced extreme wind conditions represent an important component of wind turbineWT design standards (IEC, 2019b, a; Ju et al., 2021; Martín del Campo et al., 2021). For offshore wind farm lease areas (LALAs) in coastal waters along the U.S. East Coast (Fig. 1), past research has suggested areas north of Maryland have the lowest risk of hurricane damage (< 5% probability that in a 20 yearyr period more than 10% of the wind turbinesWTs would be destroyed) (Rose et al., 2012b, a). Offshore LALAs near North Carolina experienced fewer than 40 instances of hurricane force winds (10-m wind speeds WS > 33 m s⁻¹) between 1900 and 2013, while those located near Maryland and farther northward experienced ~ 20 instances (Hallowell et al., 2018). Based on output from the 30 km resolution ERA5 gridded dataset, the highest 50-year yr return period (RP) wind speed (WS) at 100 m above sea level (a.s.l.) and significant wave heights (Hs, i.e., the mean height - crest to trough of the largest one-third of waves) for the U.S. East Coast offshore wind energy LALAs are ~ 39.7 m s⁻¹ and ~ 11 m, respectively (Barthelmie et al., 2021). Equivalent estimates from buoy measurements are; 32.6 m s⁻¹ and 9.5 m (Kresning et al., 2024). A further model-based study indicated 50-year yr RP Hs of 9-to 11 m for the northernmost LA considered here (McElman et al., 2024).

Wind-wave coupling plays a key role in both near-surface atmospheric processes and wind turbineWT loading (Valamanesh et al., 2013; Valamanesh et al., 2015; Koukoura et al., 2016; Hallowell et al., 2018; Hashemi et al., 2021; Li et al., 2022; Müller et al., 2024). Analyses based on buoy-based measurements of wind and waves along the U.S. East Coast indicated the mean failure probability during a 20-year wind turbine yr WT lifetime is 9.6 × 10-6 for a functional wind turbineWT yaw control system and 2.9 × 10-4 for a non-functional yaw control system (Hallowell et al., 2018). Wind-wave directional offset is also considered in offshore wind-turbineWT design codes (IEC61400-3) for loading on the support structure (IEC, 2019a). A 90° wind-wave misalignment is projected to increase the mud-line bending moment for a monopile foundation by up to a factor of five, and even more modest misalignment of 30° approximately doubles this bending moment (Fischer et al., 2011). This amplification of bending moment with wind-wave misalignment is greatly enhanced under high HH WS (Stewart and Lackner, 2014).

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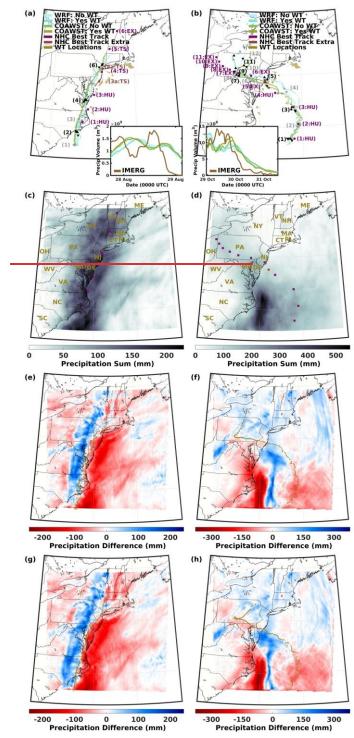


Figure 1: Hurricane Irene (left) and Hurricane Sandy (right). (a, b) Hurricane tracks from model simulations presented herein (WRF without the action of wind turbines, WRF with the action of wind turbines, COAWST without the action of wind turbines, and COAWST without the action of wind turbines, derived from minimum SLP every 10 min and the corresponding NHC "best track" locations. HU indicates hurricane stage, TS indicates tropical storm stage, and EX indicates extratropical stage. Numbers in brackets represent the location at the times specified in Table 1. The time series plots are the 1-h precipitation volume within 375 km from the track location for each simulation and IMERG V07. (c, d) Accumulated IMERG V07 precipitation during the entire simulation period (1200 UTC on 24 August 2011 through 1200 UTC on 29 August 2011 and 1200 UTC on 25 October 2012 through 1200 UTC on 1. November 2012, respectively) and NHC best track locations (from (a) and (b)). Also shown are the U.S. state abbreviations; Maine (MF), New Hampshire (NH), Vermont (VT), Massachusetts (MA), Rhode Island (RI), Connecticut (CT), New York (NY), New Jersey (NJ), Pennsylvania (PA), Ohio (OH), West Virginia (WV), Maryland (MD), Delaware (DE), Virginia (VA), North Carolina (NC), and South Carolina (SC), (c, f) Difference in accumulated precipitation (WRF without the action of wind turbines minus IMERG) and the hurricane track from the WRF no WT simulation (every 10 min), (g, h) Difference in accumulated precipitation (CV) on the COAWST without the action of wind turbines minus IMERG) and the hurricane track from the WRF no WT simulation (every 10 min).

Table 1: The time (UTC) and date that correspond to the numbered locations for Hurricanes Irene and Sandy shown in Fig. 1. All dates for Irene are in 2011 and all dates for Sandy are in 2012.

location	Irene time and date	Sandy time and date
1	1800 27 Aug	0000 29 Oct
2	0000 28 Aug	0600 29 Oct
3	0600 28 Aug	1200 29 Oct
3a	0935-28 Aug (landfall at	-
	Brigantine Island, NJ)	
4	1200 28 Aug	1800 29 Oct
4 a	1300 28 Aug (landfall at	2100 29 Oct (downgraded
	Coney Island, NY)	to extratropical)
4 b	-	2330 29 Oct (landfall at
		Brigantine Island, NJ)
5	1800 28 Aug	0000 30 Oct
6	0000 29 Aug	0600 30 Oct
7	-	1200 30 Oct
8	-	1800 30 Oct
9	-	0000 31 Oct
10	-	0600 31 Oct
11	-	1200 31 Oct
12	-	1800 31 Oct

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Estimation of design criteria extreme wind speedsWSs from numerical modeling is critically dependent on the grid spacing at which the model is applied (Larsén et al., 2012) and momentum dissipation at the ocean surface which in turn is determined by wind-wave coupling and the parameterization used to dictate the surface roughness length (z₀) (Larsén et al., 2019; Wang et al., 2024). The Coupled Ocean-Atmosphere-Wave-Sediment Transport (COAWST) modeling system (Warner et al., 2010) comprises a series of linked model components. In thethis research presented herein these modelsmodel components are; the; Weather Research and Forecasting Model-(WRF) model (Skamarock et al., 2019), the Regional Ocean Modeling System (ROMS) (Shchepetkin and McWilliams, 2005; Haidvogel et al., 2008; Shchepetkin and McWilliams, 2009), and Simulating Waves Nearshore (SWAN) model (Booij et al., 1999) and they interact through use of the Model Coupling Toolkit (MCT) (Jacob et al., 2005; Larson et al., 2005; Warner et al., 2008).

Only limited previous research has sought to quantify the degree to which wind-wave coupling improves simulation fidelity and/or intensity for wind speedsWSs at heights of relevance to offshore wind turbinesWTs. One such study focused on 23 intense cyclones in the North Sea and found that when WRF is coupled to SWAN through a wave boundary layer model with an innermost domain with grid spacing of 2 km, the inferred 50-year yr RP wind speeds WSs were systematically higher than those from WRF alone and the degree of agreement in extreme wind speedsWSs at five offshore and/or coastal masts was improved (Larsén et al., 2019). A further study found that for Tropical Storm Ana in the mid-Atlantic Bight, two-way coupled WRF and WaveWatch III (WW3) simulated peak 90-m wind speedsWS were a closer match to observations than the corresponding values from either a standalone WRF or one-way coupled WRF simulation (Gaudet et al., 2022). Simulation of Hurricane Sandy with WRF with FVCOM (coupled to the unstructured-grid, Finite-Volume Community Ocean Model) 135 coupled (FVCOM) through the Earth System Model Framework also found improved agreement with observations for the central pressure location and intensity plus 10- m wind speedsWSs relative to simulation solely with WRF (Li and Chen, 2022). COAWST (configured with WRF 3.2, ROMS 3.3, and SWAN 40.81) coupled with the MCT showed "modest improvement in track but significant improvement in intensity versus uncoupled (e.g., standalone atmosphere, ocean, or 140 wave) model simulations" for Hurricane Ivan (Zambon et al., 2014a). Thus, there is provisional evidence that, in accord with expectations, detailed coupling of the (Zambon et al., 2014b). Thus, there is provisional evidence that, in accord with expectations, detailed coupling of atmosphere-wave and ocean models improves simulation of atmospheric parameters within these extreme events relative to simulations with WRF alone.

In addition to the importance of intense cyclones (e.g., hurricanes) to wind turbineWT design standards, there have also been suggestions that very wide-spread deployment of offshore wind turbinesWTs in the U.S. coastal zone could aid in reducing the intensity of tropical cyclones and thus reduce damage onshore (Jacobson et al., 2014). Simulations of three hurricanes using the GATOR-GCMOM model indicated that offshore wind turbineWT arrays comprising 110,000 and 420,000 wind turbines (installed capacities WTs (IC > 300 GW) at installed capacity densities (ICD) of 8 -to 17 MW km⁻² might reduce 15-m wind speedsWS by over 25 m s⁻¹ and reduce storm surge by up to 79% (Jacobson et al., 2014). Simulations with 22,000 to 74,000 wind turbines WTs also suggested that offshore wind turbines WTs could reduce the amount of precipitation over land, downstream of the wind farms (Pan et al., 2018).

1.2 Objectives

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Research presented herein uses storyline simulations of two of the most powerful hurricanes that have occurred within the U.S. East Coasteastern coastal waters in which offshore wind energy LALAs have been auctioned (Fig. 1, see further details in Table S1 and Figs. S1-S2). Four sets of simulations are performed for each of these hurricanes; (a) WRF, (b) WRF with the action of wind turbines WTs included in offshore wind energy lease areas LAs purchased prior to mid-2023, (c) COAWST, and (d) COAWST with the action of wind turbines WTs included. Our specific research questions are as follows:

- 1) Are the characteristics of these hurricanes well captured using either the WRF or COAWST models? A subcomponent of this question is does the more explicit coupling in COAWST improve simulation fidelity?
- 2) Do these simulations suggest that either of these hurricanes would have been characterized by either (a) widespread loss of power production across these lease areasLAs due to cut-out at high wind speedsWSs and for how long and/or (b) exceedance of wind turbineWT design wind speedsWSs and/or very high wind-wave structural loading? Again, a sub-component of this question is does use of COAWST versus WRF change wind speedWS intensity and/or the duration of time with low power production?

3) If wind turbine WT rotor extraction of momentum is simulated using a wind farm parameterization WFP in WRF and COAWST, is there evidence of weakening of the hurricanes for wind turbineWT numbers and installed capacity densities [CD] that are likely to be achieved using the offshore wind energy lease areas [As] considered here?

2 Data and Methods

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2.1 Characteristics of the hurricanes considered herein

Research presented herein focuses on two recent hurricanes:

- 1) Hurricane Irene became a category 3 hurricane, with 54 m s⁻¹ wind speedsWSs at 10- m height in the Bahamas at 1200 UTC on 24 August 2011 12:00 UTC (Avila and Cangialosi, 2011). It made landfall at Cape Lookout, North Carolina at 1200 UTC on 27 August 12:00 UTC with 39 m s⁻¹ 10-m wind speeds WSs. After moving out over the water, it again made landfall, this time as a tropical storm, with 31 m s⁻¹ wind speeds WSs reported at Brigantine, New Jersey at 0935 UTCon 28 August, 2011 09:35 UTC (Fig. 1a). The cyclone then moved over Coney Island, New York with 28 m s⁻¹ wind speedsWSs reported at 130013:00 UTC. Simulations presented herein are initialized at 1200 UTC on 24 August 2011 12:00 UTC and run through 1200 UTC on 29 August 2011 12:00 UTC.
- 2) Hurricane Sandy became a category 3 hurricane, with 51 m s⁻¹ wind speeds WSs at 10-m height in eastern Cuba at 0525 UTC on 25 October 2012 05:25 UTC (Blake et al., 2013; Lackmann, 2015). It grew to have a roughly 1611 km diameter of tropical-storm-force wind speedsWSs, before making landfall near Brigantine, New Jersey as a posttropical cyclone with 36 m s⁻¹ 10-m wind speeds WSs and a minimum pressure of 945 hPa at 2330 UTCon 29 October-2012 23:30 UTC (Fig. 1b). Simulations presented herein run from 1200 UTC on 25 October 2012 12:00 UTC through 1200 UTC on 1 November 2012 12:00 UTC.

190 2.2 Modeling

The source of initial and lateral boundary conditions (Khaira and Astitha, 2023) and specific model configurations employed within WRF and COAWST (including the coupling system) have a critical impact on simulated flow conditions (Mooney et al., 2019). In this research, both WRF (v4.2.2) and COAWST (v3.7 and MCT v2.6.0) simulations use two domains (Fig. 2a) and the coupling interval in COAWST is 10 min (Fig. 2b). The source of boundary and initial conditions and key physics options (Tables 3 and 4). At this coupling interval, a number of variables that are critical to air-sea coupling and lower atmosphere structure and/or WT design standards are exchanged between the model components (Fig. 2b, Fig. S3, and Table S2). The selection of these variables is based on previous research (Warner et al., 2010; Zambon et al., 2014b) and include sea surface temperature (SST) that is passed from ROMS to WRF, 10 m u- and v-wind components which are passed from WRF to SWAN, plus Hs and Tp (period of peak energy in the wave spectrum) that are passed from SWAN to WRF and ROMS. The source of boundary and initial conditions (Table 1) and key physics options are informed by previous simulations of Hurricanes Sandy (Zambon et al., 2014b)(Zambon et al., 2014a) and Irene (Mooney et al., 2016). The MYNN2 planetary boundary layer scheme is used due to the compatibility with the Fitch wind farm parameterization (WFP)Physics settings include the WRF single-moment 6-class (WSM6; (Hong and Lim, 2006)) microphysics scheme, the Rapid Radiative Transfer Model (RRTM; (Mlawer et al., 1997)) for longwave radiation, the Dudhia scheme (MM5; (Dudhia, 1989)) for shortwave 205 radiation, and the Unified Noah land surface model (Chen and Dudhia, 2001b, a; Ek et al., 2003; Tewari et al., 2004). The Kain-Fritsch (Kain, 2004) cumulus parameterization is used in the outer domain and no cumulus parameterization is used in the inner domain. The Mellor-Yamada Nakanishi and Niino Level 2.5 (MYNN2; (Nakanishi and Niino, 2006)) planetary boundary layer scheme is used due to the compatibility with the Fitch windfarm parameterization (WFP: (Fitch et al., 2012))

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that is used here in both domains to compute power production, momentum extraction, and turbulent kinetic energy (TKE) induced by the action of wind turbines. WTs. Following previous research (Pryor and Barthelmie, 2024b, a), we assume that all auctioned offshore lease areas along the U.S. East Coast shown in Fig. 2a are populated with IEA reference 15 MW wind turbines (Fig. 2e) with a spacing of 1.85 km for an average ICD of 4.3 MW km⁻². This results in a total of 2642 wind turbines (Fig. 2a), each of which has a hub height of 150 m, and a rotor diameter of 240 m. The 1.85 km wind turbine (Pryor and Barthelmie, 2024a, b), we assume that all auctioned offshore LAs along the U.S. East Coast (Fig. 2a) are populated with 2642 215 IEA reference 15 MW WTs, each of which has a hub height of 150 m, and a rotor diameter of 240 m (see power and thrust curves in Fig. 2c), at a spacing of 1.85 km for an average ICD of 4.3 MW km⁻². This spacing and 1.33 km WRF domain 02 (d02) grid spacing (dx) results in 2641 grid cells with at least one wind turbineWT; one grid cell has two wind turbinesWTs₄ Of the 71 unstaggered WRF vertical levels, level 15 has a mean height of 155 m in grid cells with wind turbines WTs and is therefore used for HH WS. Note the wind speeds it is not an expectation that spatially averaged model output from d02 are for a nominal model will perfectly match time-step of 2 s but are representative of a spatial average of 1.33 km by 1.33 km, whileaveraged point observations and further, the design standards are articulated for a sustained wind speedWS at a point (Larsén and Ott, 2022). The WSs presented here are output from d02, represent a nominal model time step of 2 s, and are from a grid cell with an area of 1.33 km by 1.33 km, but the effective model resolution is ~ 7 times the grid spacing (Skamarock, 2004) thus any spatial gradients will be under-estimated.

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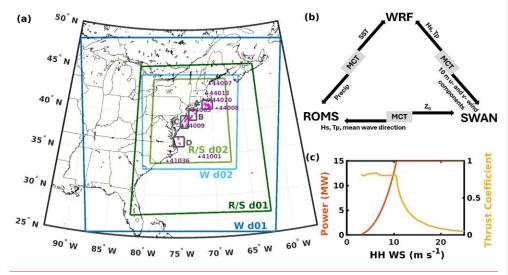


Figure 2: (a) Simulation domains (d01 and d02) for WRF (W), ROMS (R), and SWAN (S) and locations of the wind turbines (WTs) in the offshore wind energy lease areas (LAs) considered herein (magenta). In the simulations with WTs, these LAs contain 2642 WTs. Also shown are the locations of the eight National Data Buoy Center (NDBC) buoys used in the simulation evaluation (purple). (b) Schematic of information flow between COAWST model components for variables discussed herein: significant wave height (Hs), peak wave period (Tp), 10 m u- and v- wind components, surface roughness length (z0), mean wave direction, precipitation (Precip), and sea surface temperature (SST). For the complete information flow, see Fig. S3 and Table S2. (c) Power and thrust coefficients as a function of hub-height wind speed (HH WS) for the 15 MW IEA reference WT (Gaertner et al., 2020).

Variation of wave state and surface roughness length with wind speed is an important determinant of extreme, near surface wind speeds and turbulence intensity (Zambon et al., 2014a; Porchetta et al., 2019; Porchetta et al., 2020; Porchetta et al., 2021; Wang et al., 2024) Variation of wave state and z₀ with WS is an important determinant of extreme, near surface WSs and Formatted: English (United Kingdom)

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Table

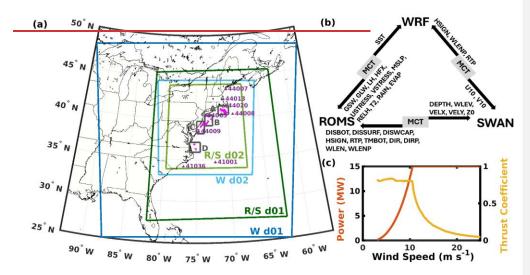


Figure 2: (a) Domains 01 and 02 (d01 and d02) for WRF (W), ROMS (R), and SWAN (S) and locations of offshore wind energy LA considered herein (in the simulations with wind turbines, these LA contain 2642 wind turbines which are shown in magenta). Also shown are the eight NDBC buoys used in the simulation evaluation. (b) Schematic of information flow among the COAWST model components. See Table 2 for additional details. (c) Power and thrust coefficient for the 15 MW IEA reference wind turbine!: Model configuration for WRF and COAWST simulations. (Gaertner et al., 2020).

275 Table 2: Additional details about the information flow among the COAWST model components in Fig 2b.

name	abbreviation	units
depth below mean sea level	DEPTH	m
mean wave direction	DIR	rad
peak wave direction	DIRP	rad
energy dissipation due to bottom friction	DISBOT	W m -2
energy dissipation due to surf breaking	DISSURF	W m -2
energy dissipation due to white-capping	-DISWCAP	W m ⁻²
evaporation rate	EVAP	kg m⁻² s⁻¹
downward long wave flux at ground surface	GLW	W m ⁻²
net short wave flux at ground surface	GSW	W m ⁻²
upward heat flux at the surface	HFX	W-m ⁻²
significant wave height	HSIGN	m
latent heat flux at the surface	LH	-W m ⁻²
mean sea level pressure	-MSLP	mb
rainfall rate	RAIN	kg m ⁻² s ⁻¹
relative humidity	RELH	-
relative peak period	RTP	s
sea surface temperature	SST	K
temperature at 2 m	T2	<u>°</u> €
near bottom wave period	TMBOT	<u>-\$</u>
x-wind component at 10 m	U10	m s ⁻¹
surface u-stress	USTRESS	N-m ⁻²
y-wind component at 10 m	V10	m·s ⁻¹
current velocity component in x direction	VELX	m s ⁻¹
current velocity component in y direction	VELY	m-s ⁻¹
surface v-stress	VSTRESS	N-m ⁻²
mean wave length	WLEN	m
peak wave length	WLENP	m
water level	WLEV	m
roughness length	Z0	m

Table 3: Sources of initial and boundary conditions for WRF and COAWST.

WRF (version 4.2.2)	
Atmosphere:	North American Mesoscale Forecast System (NAM; 12 km, 6 h)
Sea Surface Temperature:	Group for High Resolution Sea Surface Temperature (GHRSST) Level 4 Operational Sea Surface Temperature and Sea Ice Analysis (OSTIA) (OSTIA-UKMO-L4-GLOB-v2.0; 0.05°, 24 h)
Horizontal resolution Resolution:	4 km for d01, 1.33 km for d02
Model top / # vertical levels Top /	50 hPa / 72
Vertical Levels:	
Time stepStep (dt):	6 s for d01, 2 s for d02
ROMS (version 3.9)	
Coastlines:	Global Self-consistent, Hierarchical, High-resolution Geography Database (GSHHG; full resolution)
Bathymetry (also for SWAN):	General Bathymetric Chart of the Oceans (GEBCO) 2022 (15 arcsecond)
3D boundary conditions, initial	HYbrid Coordinate Ocean Model (HYCOM GLBa0.08 expt 90.9)
conditionsBoundary Conditions,	
<u>Initial</u>	
Conditions, and	
elimatologyClimatology:	
2D boundary conditions	Advanced Three-Dimensional Circulation Model (ADCIRC 2001v2e)
(tidesBoundary Conditions (Tides):	101 (101 2 221 (102
Horizontal resolution Resolution:	10 km for d01, 3.33 km for d02
Time step (baroclinic / barotropic Step	1.5 s (d01), 0.5 s (d02) / 30 (d01 & d02)
(Baroclinic / Barotropic): # vertical levels Vertical Levels:	25
	23
SWAN (version 41.31)	
Wind forcingForcing:	Global Forecast System (GFS): 0.5°, 6 h
Boundary conditions Conditions:	parametric forcing file (TPAR): 30 min WaveWatch III (WW3) data
Horizontal resolution Resolution:	10 km for d01, 3.33 km for d02
Time stepStep:	12 s for d01, 4 s for d02
Frequency range Range:	0.04 to 1.0

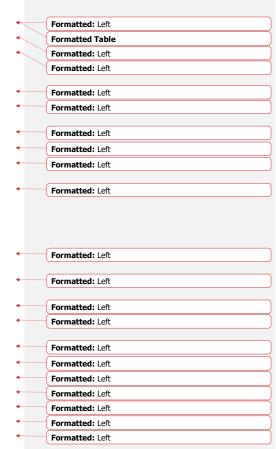


Table 4: Physics settings used with WRF and COAWST simulations.

Model physics	Key reference(s)
microphysics: WRF single-moment 6-class scheme (WSM6)	(Hong and Lim, 2006)
longwave radiation: Rapid Radiative Transfer Model (RRTM)	(Mlawer et al., 1997)

shortwave radiation: Dudhia scheme (MM5)	(Dudhia, 1989)
surface layer: MYNN	(Olson et al., 2021)
land surface: Unified Noah land surface model	(Chen and Dudhia, 2001b, a; Ek et
	al., 2003; Tewari et al., 2004)
planetary boundary layer: Mellor-Yamada Nakanishi and Niino Level 2.5	(Nakanishi and Niino, 2006)
(MYNN2)	
cumulus param.: Kain-Fritsch (d01; none for d02)	(Kain, 2004)
wind farm param. (both domains)	(Fitch et al., 2012)

295 2.3 Evaluation data sets

 $Critical \ aspects \ of \ the \ WRF \ and \ COAWST \ simulations \ without \ the \ action \ of \ \frac{wind \ turbines}{wind \ turbines} \\ \underline{WTs} \ are \ evaluated \ using: \ using \ u$

- (i) National Hurricane Center (NHC) "best track" information and wind radii maximum extent data from Tropical Cyclone Reports (Avila and Cangialosi, 2011; Blake et al., 2013) and the Atlantic hurricane database (HURDAT2; (Landsea and Franklin, 2013)).
 - (ii) 30-minute min precipitation at 0.1-degree^o resolution from the Integrated Multi-satellitE Retrievals for the Global Precipitation Measurement (GPM) mission (IMERG) V07 final run data set (Huffman et al., 2024).
- National Data Buoy Center (NDBC) buoy-based measurements of <u>WS and</u> wind <u>speeds (WSgust (WG)</u>, sea level pressure (SLP), <u>sea surface temperature (SST)</u>, and <u>significant wave height (Hs) (NDBC, 2009)(NDBC, 2009) (https://www.ndbe.noaa.gov) (see Fig. (Fig. 2a)</u>. Note: The eight buoys from which data are presented are a mixture of 3-<u>meter m</u> foam and 3-<u>meter m</u> discus buoys. The anemometer and barometer heights vary between 3.8 and 4.1 m, and 2.4 and 3.4 m.

These data sets do have some inherent constraints, which include use of; subjective smoothing to produce representative 6 h best track data which does not necessarily equate to a precise storm history (Landsea and Franklin, 2013), spatial averaging on the gridded IMERG data which can underestimate high precipitation rates compared to point measurements (Hu and Franzke, 2020; Nie and Sun, 2020; Huffman et al., 2024), and the limited number and spatial coverage of buoys (NDBC, 2009).

2.4 Analysis approach

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Hurricane centroid locations are computed every 10 minutesmin as the minimum SLP after 3×3 smoothing is applied to the model output and are used(a mean value of SLP is computed for comparisoneach grid cell based on output for that grid cell and the eight adjacent grid cells) and compared with the NHC best track information. The initial tracking position is the first time step when the minimum SLP is fullylies within d02 and tracking continues until the implied translational speed/direction of motion between adjacent time steps is inconsistent with physical expectations in terms of direction or translational speed. Hurricane Irene is tracked for 33 h, from 1800 UTC 27 August 2011 18:00 UTC through 0300 UTC 29 August 2011 03:00 UTC and Hurricane Sandy is tracked for 67 h, from 2300 UTC 28 October 2012 23:00 UTC through 1800 UTC 31 October 2012 18:00 UTC (Fig. 1). Evaluation relative to SLP and wind speedWS data from the NDBC buoys is performed using a search area of 3×3 grid cells. Evaluation of simulated precipitation within d02 relative to IMERG over all of d02 is performed after regridding output to the IMERG grid. (Fig. 3). The volume of water exhausted as precipitation from the tropical cyclone

is computed using a search radius of 375 km (see examples in Fig. 34) around the cyclone centroid in the simulation output and from the best track locations applied to IMERG.

330 WhenBecause prior research has indicated the challenges in perfectly reproducing hurricane tracks, when evaluating hurricane impacts on wind turbines WTs within the U.S. East Coast LALAs, analyses are presented for both all 2641 grid cells containing wind turbines, and WTs, four LA clusters defined as in (Pryor and Barthelmie, 2024b) and listed from north to south; (Pryor and Barthelmie, 2024a); A (1073 WT), B (662 WT), C (624 WT), and D (283 WT) (Fig.), and 2a). Because prior research has indicated the challenges in perfectly reproducing hurricane tracks, we consider conditions in both grid cells containing wind turbines and/or all ocean-based grid cells within the respective LA cluster area. To facilitate comparison across the LA clusters, power production computed by the Fitch WFP is used to compute capacity (Fig. 2a). Capacity factors (CFCFs), which are the ratio of the power produced divided by that which would be produced if all wind turbines WTs were operating at rated capacity (15 MW)-), are used to facilitate comparison of power production computed by the Fitch WFP across the LA clusters. Systemwide CF < 0.2 are used here as an indicator of low power production.</p>

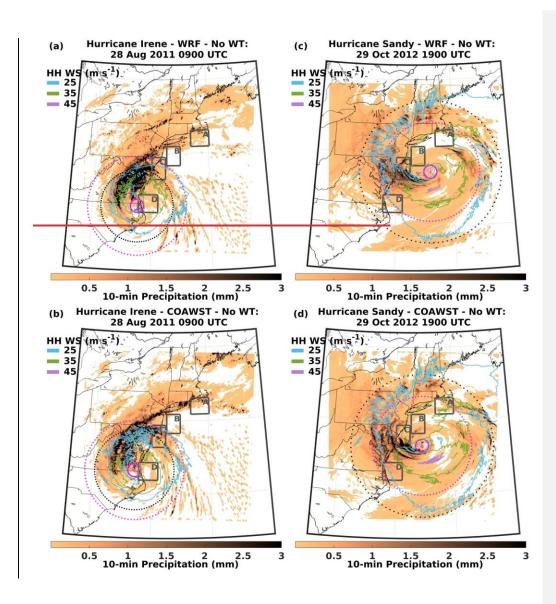
Three-dimensional and joint occurrences of HH WS, Hs, and peak period (Tp), in WT-containing grid cells from the COAWST simulations are presented along with histograms of estimated wind-wave misalignment at the LA cluster centers in HH WS classes of $3 - to < 10.6 \text{ m s}^{-1}$, $10.6 - to 25 \text{ m s}^{-1}$, and $> 25 \text{ m s}^{-1}$, to represent high thrust, moderate thrust, and above rated wind speeds WS (Fig. 2c)

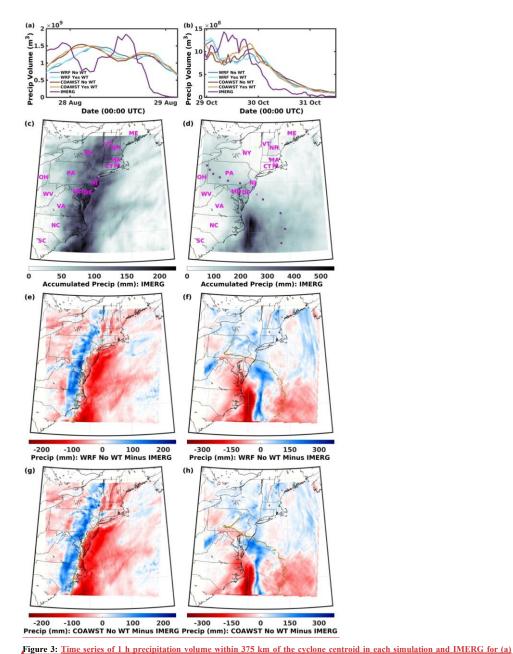
speedsWS (Fig. 2c).

Three metrics are used to analyze the impact of wind turbines WTs on hurricane intensity and are compared for simulations with WRF and COAWST without and with the WFP active. The cumulative volume of precipitation within 375 km of the minimum SLP and the mean wind speed WS at 500 hPa (approximately the level of non-divergence, e.g., (Riehl and Malkus, 1961)) computed for grid cells that lie 50-to 375 km from the centroid (i.e., beyond the likely eye radius) (Morin et al., 2024; Müller et al., 2024) are used as metrics of intensity. The mean outermost radius of tropical storm force WSs at 10 m (R₁8, 18 m s⁻¹, see Fig. 3) wind speeds at 10 m4) is computed using azimuth sectors of 10° (Powell and Reinhold, 2007) for all sectors where the distance from the cyclone centroid to the d02 boundary is ≥ 200 km isand used as a measure of cyclone size. Mood's test (Hettmansperger and Malin, 1975) is used to assess the statistical significance of differences in the median values of these metrics.

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Hurricane Irene and (b) Hurricane Sandy. Accumulated precipitation from IMERG for (c) Hurricane Irene: 24 August 2011 12:00 UTC through 29 August 2011 12:00 UTC and (d) Hurricane Sandy: 25 October 2012 12:00 UTC through 1 November 2012 12:00 UTC. National Hurricane Center (NHC) best track locations are shown in purple and the magenta letters denote, the U.S. state abbreviations; Maine (ME), New Hampshire (NH), Vermont (VT), Massachusetts (MA), Rhode Island (RI), Connecticut (CT), New York (NY), New Jersey (NJ), Pennsylvania (PA), Ohio (OH), West Virginia (WV), Maryland (MD), Delaware (DE), Virginia (VA), North Carolina (NC), and South Carolina (SC). Difference in modelled total precipitation minus IMERG for (e, g) WRF for (e) Hurricane Irene and (g) Hurricane Sandy and (f, h) COAWST for (f) Hurricane Irene and (h) Hurricane Sandy. Simulations exclude the actions of wind turbines (WTs). Brown dots denote the hurricane location (every 10 min) from the simulations.

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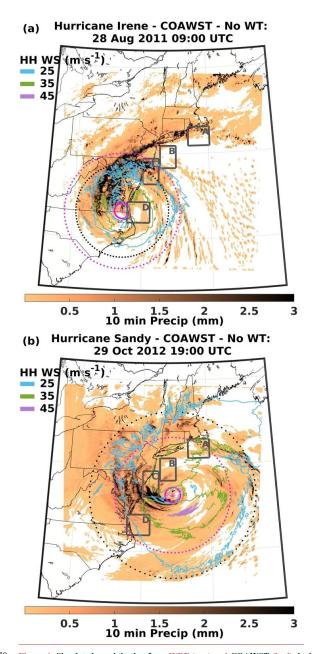


Figure 4: Simulated precipitation from WRF (a, c) and COAWST (b, d) during two example 10 minutemin periods (background color) and contours of hub-height wind speed (HH WS) at 25, 35, and 45 m s⁻¹ for 0900 UTC(a) 28 August 2011 (a, 09:00 UTC and (b) and 1900 UTC(2) October 2012 (e, d)19:00 UTC when the hurricanes are close to wind turbine LA-(WT) lease areas (LAs). Magenta rings mark 50 km and 375 km from the minimum sea level pressure (SLP-). The black rings mark R14the mean outermost radii of 18 m s⁻¹ WSs at 10 m (R18) of (a) 274 km, (b) 301 km, (e) 539 km, and (db) 541 km. For legibility, the colorbar is truncated. Maximum 10 min precipitation in any grid cell is (a) 14.0 mm, (b) 13.8 mm, (e) 18.4 mm, and (db) 16.8 mm. Similar output from WRF is shown in Fig. S4.

3 Results and discussion

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3.1 Evaluation of the no-wind turbine simulations without the action of wind turbines

Simulations3.1.1 Hurricane Irene

As shown in detail below, simulations of Hurricane Irene exhibit lower fidelity than those of Hurricane Sandy. The centroid of Hurricane Irene is consistently displaced west (further inland, Fig. 1a) than the NHC best track data, and the translational speed is also negatively biased in simulations with both WRF and COAWST simulations. This bias is consistent with previous COAWST simulations of this hurricane performed with 12 km grid spacing and using a range of initial and lateral boundary conditions (Mooney et al., 2019). Simulation bias is also evident in comparison with buoy observations (Table \$2, Fig. 45, and Figs. 83-8985-811) both in terms of the magnitudes and timing of the maximum wind speedsWSs and minimum SLP. The maximum Maximum near-surface wind speeds-WSs (at 2.6 m) differ (model minus buoy observations) by between -4.2 andto 0.7 m s⁻¹ (WRF at 2.6 m) and -3.2 andto 6.0 m s⁻¹ (COAWST at 2.6 m). The displacement of the with Hurricane Irene centroid in the simulations results in higher over land precipitation (by up to 209 mm in some IMERG grid cells) and negative bias offshore (Fig. 43). However, the volume of water vented from the hurricane is relatively well reproduced in the simulations. During the time of tracking, the mean (1-h) precipitation volume within 375 km of the centroid is 9.87×10^8 m³ based on IMERG combined with the NHC best track data, while the corresponding values (and percent error) are 1.14 × 109 m³ (15.1% overestimation) and 1.18 × 109 m3 (19.9% overestimation), for the WRF and COAWST simulations, respectively (Fig. 43). Mean R₁₈ is 279 km in the WRF simulation but is larger by an average of 23 km in 192 of the 199 tracked 10- min positions in the COAWST simulation. These mean R_{18} values are similar to, but smaller than, those reported at a 6-hrly h interval from HURDAT2 based on analyses in four quadrants (mean R₁₈ of 495 km in the SE quadrant and 157 km in the NW quadrant). The differences in R₁₈ and precipitation in thesesthese simulations versus HURDAT2 and IMERG are likely due to bias introduced by proximity challenges in accurately computing these metrics when the simulated position is close to the d02 boundary in the simulation and differences in the fraction of the systemcyclone over land (Chen and Yau, 2003) due to differences in storm tracks (Fig. 1).

3.1.2 Hurricane Sandy

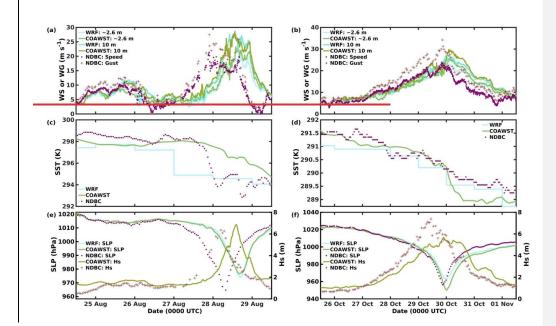
As in past research (Zambon et al., 2014b) (Zambon et al., 2014a), both the WRF and COAWST simulations of Hurricane Sandy exhibit good agreement with NHC best track data up to about 12 h after landfall in New Jersey (Fig. 1b). The mean distance between the NHC and simulated cyclone centroids for locations; 1, 2, 3, 4, 4a, 4b, and 5 in Fig. 1b is 66.7 km (WRF) and 51.4 km (COAWST). For those seven times the distances of separation The positional offsets range from 30.3 to 129.1 km (WRF) and 12.0 to 71.1 km (COAWST). These positional offsets) and are smaller than those presented in previous research on, for example, cyclonic storm Ockhi (Mukherjee and Ramakrishnan, 2022) and the mean values fromplus Tropical Storm Delta, Hurricane Ophelia, Hurricane Leslie, and Tropical Storm Theta (Calvo-Sancho et al., 2023). Consistent with 410 expectations, agreement tends to degrade once the cyclone has made landfall as the system becomes less organized and more asymmetric (Zambon et al., 2014b). Modeled time series of SLP, SST, Hs, and wind speeds (Zambon et al., 2014a). Modeled time series of SLP, SST, Hs, and WS exhibit some level of agreement with the NDBC buoy observations in terms of timevariability (Table 63, Fig. 45, and Figs. 83-8985-811). As expected, due to spatial averaging and the difference in height, the maximum wind speeds WSs from the lowest model level (~2.6 m) are generally lower than the point observations on the buoys at 3.8 or 4.1 m. Seven of the eight values are smaller in the WRF simulation than reported by the buoys, and six of eight comparisons indicate lower values in the COAWST simulation. The maximum Maximum near-surface wind speeds WSs

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(model minus buoy observations) differ by 3.1 to 9.0 m s⁻¹ (WRF) and 2.9 to 7.8 m s⁻¹ (COAWST). The mean absolute difference in minimum SLP is 3.1 hPa (WRF) and 2.6 hPa (COAWST). Significant wave heights Hs from COAWST are generally lower than those from the buoy observations, negatively biased with adifferences (model minus buoy observations) difference of -3.7 to 0.1 m (a-bias of up to 37%). The spatial pattern of precipitation from the simulations exhibits some similarity with IMERG, although the centroid of the region of maximum precipitation centroid is displaced westward (towards the coast of North Carolina) in IMERG relative to the WRF and COAWST simulations by approximately 125 km (Fig. 43). The volume of precipitation vented from the hurricane exhibits relatively good agreement between the simulations and IMERG. The mean (1-h) precipitation volume within 375 km from the centroid is 5.35 × 108 m³ based on IMERG combined with the NHC best track data, while the corresponding values (and percent error) are 5.82 × 108 m³ (+8.7%) and 5.83 × 108 m³ (+8.9%) for the WRF and COAWST simulations (Fig. 43). Mean R₁₈ prior to landfall is 506 km in the WRF simulation and is larger by an average of 14 km in 119 of the 151 tracked 10-min positions in the COAWST simulation. The R₁₈ estimates from 6-h HURDAT2 data before and after landfall are 705 km and 776 km and are larger than those from the WRF and COAWST simulation due to biases in the calculation when the hurricane extends beyond the d02 boundary.

Table 5: Comparison of WRF and COAWST output (the simulations without wind turbines) and buoy measurements. The magnitude2: Magnitude and time (in UTC) of thenear-surface maximum wind speed (Max WS), maximum significant wave height (Max Hs), and minimum sea level pressure (Min SLP are given) for each buoy and simulations with WRF and COAWST (simulated WS-without wind turbines (WTs) for Hurricane Irene (August 2011). Simulated WSs are shown at two heights, 10 m | 2.6 m). All the times with Irene- and all variables are in August 2011. With the simulations, magnitudes are provided computed every 10 min. With WS magnitudes from the buoys, magnitudes are provided available every 10 min for WS while Hs and provided SLP are reported at 50 min past the top of the hour for Hs and SLP.

Irene		maxMax	time maxTime Max WS	max	time maxTime Max	minMi	time minTime
110110		WS (m s ⁻¹)		Max	Hs	n SLP	Min SLP
		(")		Hs		(hPa)	
				(m)		, ,	
	buoy	19.7	1840 & 1850 27 Aug <u>18:40 &</u>	10.0	1550- 27 Aug <u>15:50</u>	997.5	1850 —27 Aug
41001	WRF	18.3 16.0	0150 -28 Aug <u>01:50</u>	-	-	1002.0	1020 —28 Aug
4	COAWST	20.2 17.9	0540 -28 Aug <u>05:40</u>	5.1	0500- 28 Aug <u>05:00</u>	1002.1	1040 —28 Aug
9	buoy	25.1	1620 -27 Aug <u>16:20</u>	8.6	0550 -27 Aug <u>05:50</u>	957.1	0950 _27 Aug
41036	WRF	27.1 22.9	2240 -27 Aug <u>22:40</u>	-	-	965.3	2340 —27 Aug
4	COAWST	34.4 31.1	2340 -27 Aug <u>23:40</u>	7.6	2040- 27 Aug <u>20:40</u>	958.2	2210 —27 Aug
_	buoy	16.5	1930 -28 Aug <u>19:30</u>	4.5	2150-28 Aug 21:50	983.2	0150 —29 Aug
44007	WRF	22.5 15.6	0700 -29 Aug <u>07:00</u>	-	-	977.5	0810 –29 Aug
4	COAWST	23.3 19.1	1140 1150-29 Aug 11:40	5.5	0700- 29 Aug <u>07:00</u>	975.6	0810 –29 Aug
	buoy	18.3	1750 -28 Aug <u>17:50</u>	8.2	1750- 28 Aug <u>17:50</u>	996.1	1950 —28 Aug
44008	WRF	20.0 14.1	0700 -29 Aug <u>07:00</u>	-	-	997.7	0420 —29 Aug
4	COAWST	22.0 15.1	0550 0250 -29 Aug <u>05:50 </u>	7.9	0650- 29 Aug <u>06:50</u>	997.1	0300 –29 Aug
	buoy	21.4	2140 -27 Aug <u>21:40</u>	6.4	0450 -28 Aug <u>04:50</u>	958.3	0650 —28 Aug
44009	WRF	27.9 22.1	1330 1710 28 Aug <u>13:30 </u>	-	-	974.0	1700 –28 Aug
4	COAWST	28.6 23.3	1330 1400 28 Aug <u>13:30 </u>	6.9	1410- 28 Aug <u>14:10</u>	976.5	1610 _28 Aug
	buoy	18.8	1830- 28 Aug <u>18:30</u>	3.8	1550 -28 Aug <u>15:50</u>	984.0	2050 —28 Aug
44013	WRF	24.7 17.1	0500 -29 Aug <u>05:00</u>	-	-	979.2	0540 —29 Aug
4	COAWST	22.7 15.7	0410- 29 Aug <u>04:10</u>	3.5	0420 & 0430 29 Aug	978.3	0610 –29 Aug
-0	buoy	21.2	2350-28 Aug 23:50	2.4	1550 -28 Aug <u>15:50</u>	989.2	2050—28 Aug
44020	WRF	26.7 18.5	0600 0840 29 Aug <u>06:00 </u>	-	-	987.2	0600 –29 Aug
4	COAWST	28.0 18.9	0520- 29 Aug <u> 05:20</u>	3.6	0410- 29 Aug <u>04:10</u>	985.9	0520 —29 Aug
35	buoy	21.1	1220 -28 Aug <u>12:20</u>	8.0	1250 -28 Aug <u>12:50</u>	968.0	1250 —28 Aug
44065	WRF	26.6 19.7	2030- 28 Aug 0310 20:30 29	-	-	974.9	2350 —28 Aug
4	COAWST	27.7 23.9	0350 -29 Aug <u>03:50</u>	5.5	2110-28 Aug 21:10	973.7	2240 —28 Aug



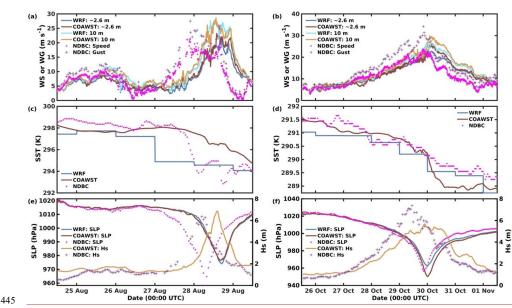


Figure 45: Time series of (a, b) wind speedsspeed (WS) and wind gusts (WS or gust (WG [m s⁻¹])₂), (c, d) sea surface temperatures (SST [K])₂), and (e, f) sea level pressure (SLP [hPa]) and significant wave height (Hs [m])) from observations and simulations (WRF and COAWST) without wind turbines [WTs]) for Hurricane Irene (a, c, and e) and Hurricane Sandy (b, d, and f) at buoy 44009 (see location in Fig. 2a). Simulations performed without the action of wind turbines See Figs. SS-S11 for other buovs.

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Table 6: Comparison of WRF and COAWST output (the simulations without wind turbines) and buoy measurements. The magnitude3: Magnitude and time (in UTC) of thenear-surface maximum wind speed (Max WS), maximum significant wave height (Max Hs), and minimum sea level pressure (Min SLP-are given) for each buoy and simulations with WRF and COAWST (simulated WS without wind turbines (WTs) for Hurricane Sandy (October 2012). Simulated WSs are shown at two heights, 10 m | 2.6 m). All the times with Sandy_and all variables are in October 2012. With the simulations, magnitudes are provided computed every 10 min. WithWS magnitudes from the buoys, magnitudes are provided available every 10 min for WSwhile Hs and provided SLP are reported at 50 min past the top of the hour-for Hs and SLP._.

Sandy		max Max	time max Time Max WS	max	time max Time Max	min <u>Mi</u>	time minTime	
		WS (m s ⁻¹)		Max	Hs	<u>n</u> SLP	Min SLP	
				Hs		(hPa)		
		20.4	0610 20 0 0610	(m)	1050 00 0 110 50	0.00.6	2250 20 0	
10	buoy	28.4	0610 -29 Oct <u>06:10</u>	10.1	1350 -29 Oct <u>13:50</u>	969.6	2350—28 Oct	
41001	WRF	37.4 32.2	0630 -29 Oct <u>06:30</u>	-	-	962.8	0250 —29 Oct	
	COAWST	36.2 31.8	0630- 29 Oct <u>06:30</u>	9.8	1520 -28 Oct <u>15:20</u>	964.3	0200 —29 Oct	
9	buoy	21.7	2250 27 Oct <u>22:50</u>	5.7	2050 & 2150 27 Oct	992.1	0950 —28 Oct	
41036	WRF	25.0 20.9	0720 -28 Oct <u>07:20</u>	-	-	995.3	1300 —28 Oct	
4	COAWST	24.6 21.5	0800 -28 Oct <u>08:00</u>	5.1	0630 -28 Oct <u>06:30</u>	995.0	1530 —28 Oct	
7	buoy	18.0	0050 -30 Oct <u>00:50</u>	7.1	0350 -30 Oct <u>03:50</u>	995.9	0050 —30 Oct	
44007	WRF	23.8 16.6	0150 -30 Oct <u>01:50</u>	-	-	995.6	0140 —30 Oct	
4	COAWST	23.6 16.6	0020 -30 Oct 2350 <u>00:20</u> 29	6.6	0230 -30 Oct <u>02:30</u>	995.6	0330 —30 Oct	
	buoy	22.4	1640- 29 Oct <u>16:40</u>	11.0	2050- 29 Oct <u>20:50</u>	981.2	1750 —29 Oct	
44008	WRF	27.4 18.6	1930 -29 Oct <u>19:30</u>	-	-	981.9	1830 —29 Oct	
4	COAWST	26.5 21.5	1700 1620 29 Oct 17:00	8.6	1750 & 1800 29 Oct	982.5	1810 —29 Oct	
6	buoy	23.7	2040- 29 Oct <u>20:40</u>	7.4	1050- 29 Oct <u>10:50</u>	956.4	2050 —29 Oct	
44009	WRF	26.8 22.8	2320-29 Oct 040023:20 30	-	-	963.0	0030 —30 Oct	
4	COAWST	29.8 25.6	0210 -30 Oct <u>02:10</u>	5.7	2140-29 Oct 21:40	950.3	0010 —30 Oct	
	buoy	20.4	1920 -29 Oct <u>19:20</u>	6.9	0150 -30 Oct <u>01:50</u>	988.2	0050 —30 Oct	
44013	WRF	25.0 17.3	2030 -29 Oct <u>20:30</u>	-	-	989.2	2300 —29 Oct	
4	COAWST	24.3 16.3	2140-29 Oct 21:40	7.0	2210-29 Oct 22:10	989.0	2230 —29 Oct	
	buoy	20.6	2000 -29 Oct <u>20:00</u>	3.1	1850 -29 Oct <u>18:50</u>	983.3	1950 —29 Oct	
44020	WRF	28.0 19.0	1940 -29 Oct <u>19:40</u>	-	-	984.4	2040 —29 Oct	
4	COAWST	25.9 17.4	1930 -29 Oct <u>19:30</u>	2.9	1530 -29 Oct <u>15:30</u>	984.7	2110 —29 Oct	
10	buoy	24.9	0010 -30 Oct <u>00:10</u>	9.9	0050 -30 Oct <u>00:50</u>	958.1	2150 —29 Oct	
44065	WRF	30.5 22.6	2310 -29 Oct 0310 23:10 30	-	-	952.5	2310 —29 Oct	
4	COAWST	30.9 22.7	2200- 29 Oct 0100 22:00 30	6.2	0100 -30 Oct <u>01:00</u>	960.9	2320 —29 Oct	

The evaluation 3.1.3 Synthesis

Evaluation of the WRF and COAWST simulations of Hurricane Sandy thus indicates relatively high fidelity. Nevertheless, the fidelity is lower for simulations of Hurricane Irene and biases relative to observations provide important context for the following analyses which focus on power production and extreme conditions at prospective offshore wind turbine WT locations. Due to the presence of errors in tropical cyclone tracking in the simulations, in the following discussion of geophysical conditions we consider not only grid cells with wind turbines WTs in the LALAs, but also ocean-based grid cells nearby. In terms of agreement with: observed precipitation, cyclone size (R₁₈), near-surface WS, and cyclone tracking, COAWST simulations exhibit higher skill than those of WRF.

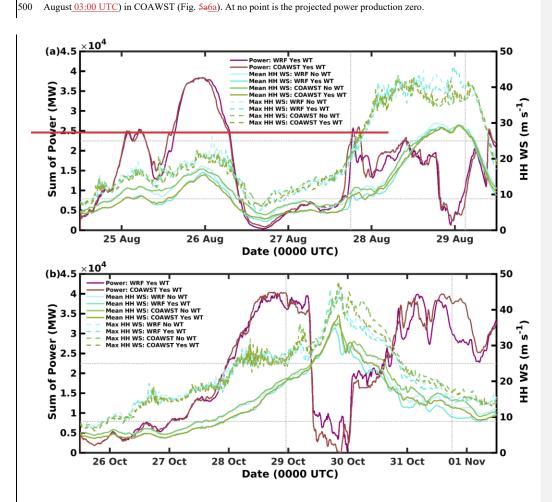
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490 3.2 Wind turbine power production and operating conditions:

3.2.1 Hurricane Irene

Mean power production and CF computed for the entire Hurricane Irene simulation period using WRF and COAWST are; 1.51×10^4 MW (0.38) and 1.56×10^4 MW (0.39), respectively. When Hurricane Irene is present in d02, equivalent CFCFs are 0.39 and 0.40, respectively. These CFvalues are slightly lower than the previously reported climatologically representative CF estimates of 0.45 presented previously (Pryor and Barthelmie, 2024b, a) (Pryor and Barthelmie, 2024a, b) due to relatively low wind speedsHH WS in the vicinity of the offshore LALAs early in the simulation and to an extended period of above cut-out wind speedsWSs during the hurricane passage from late on 27 August to the middle of 29 August (Fig. 5a6a and Fig. 6a, b7a). However, the system-wide CF only drops below 0.2 for continuous periods of 5 h 50 min (2000 UTC 28 August 20:00 UTC through 0150 UTC 29 August 01:50 UTC) in WRF and 7 h 10 min (1950 UTC 28 August 19:50 UTC through 0300 UTC 29 August 03:00 UTC) in COAWST (Fig. 5a6a). At no point is the projected power production zero.



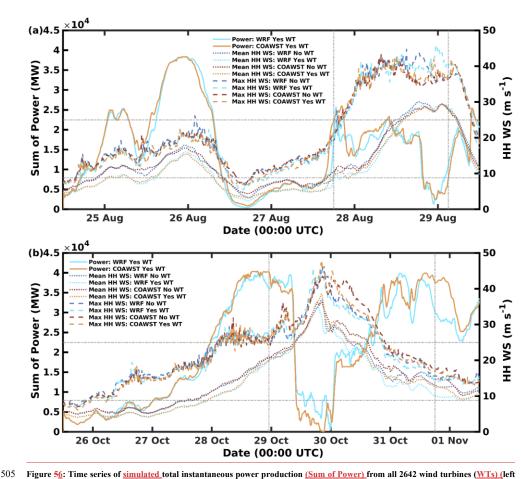
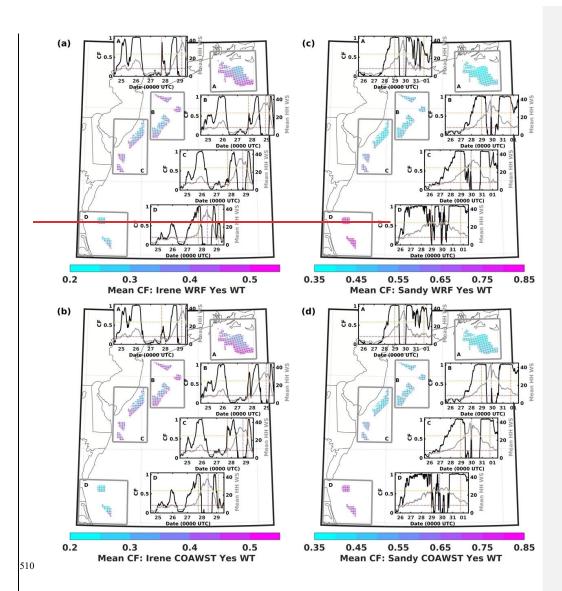


Figure 56: Time series of simulated total instantaneous power production (Sum of Power) from all 2642 wind turbines (WTs) (left axis) and mean and maximum hub-heighheight wind speed (HH WS,-) (right axis) in grid cells containing wind turbines WTs for (a) Hurricane Irene and (b) Hurricane Sandy-simulations. The dashed gray vertical lines mark the start and end time of storm tracking within d02. The lower dashed gray horizontal line marks power production equivalent to a capacity factor (CF) of 0.2, and the upper dashed gray horizontal line marks HH WS = 25 m s⁻¹.



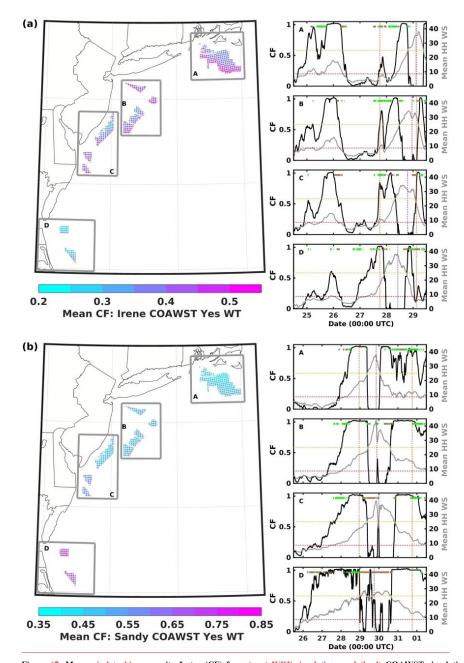


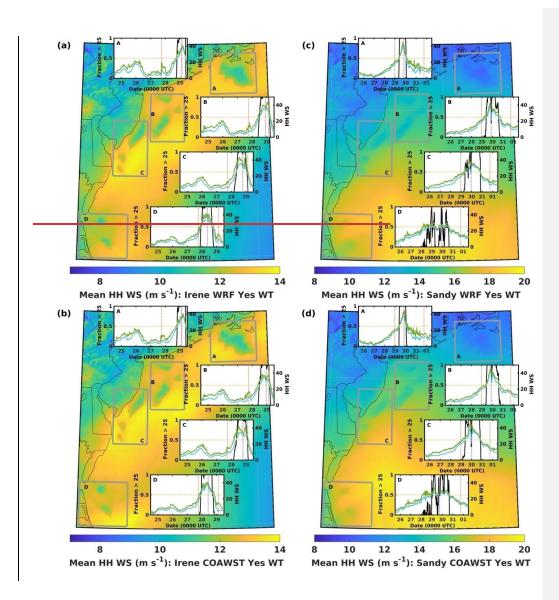
Figure 67: Mean wind turbine-capacity factor (CF) from (a, e) WRF simulations and (b, d)-COAWST simulations for (a, b) Hurricane Irene and (e, db) Hurricane Sandy. Also shown are time series of the CF and mean hub-height wind speed (HH WS) forover the four lease area (LA) clusters. Brown (green) symbols at the top of those time series indicate times when CFs from WRF exceed (are less than) those from COAWST by more than 0.05 (see also Fig. S12). Orange dashed lines indicate the start and end time of storm tracking within d02. The purple dashed line represents the time when the location of the minimum sea level pressure (SLP) is closest to the cluster center. The red dashed line indicates CF = 0.2, and the yellow line indicates HH WS of= 25 m s⁻¹.

520 Time series of power production from WRF and COAWST indicate a high degree of agreement (Fig. 5a) but there are times when the models deviate both in terms of power production and extreme wind speed. Late 6a). However, late on 27 August and early on 28 August when Hurricane Irene is south of the LAs, projected power production differs by a maximum of 1.09 \times 10⁴ MW (CF difference of 0.275). With Simulations with WRF, for the and the WFP active indicate a period of 5 h 50 min when thewith system-wide CF remains below < 0.2, the and mean HH WS in grid cells with WTs (WT grid cells for the 525 simulation with the WFP active ranges) ranging from 27.6 to 29.3 m s⁻¹. Within the LALAs the maximum HH WS is 45.4 m s⁻¹, which exceeds the 50-year yr RP WS at 100 m derived in earlier work using ERA5 output (Barthelmie et al., 2021), but remains below the 50 m s⁻¹ sustained wind speedWS threshold for class I wind turbinesWTs and the 57 m s⁻¹ threshold for tropical cyclone hardened wind turbines WTs (class T) (IEC, 2019a). Analyses including all ocean-based grid cells within the four LA clusters indicate the mean HH WS as simulated by WRF with WFP active ranges from 24.2 to 27.0 m s⁻¹ and the maximum HH WS reaches $46.4~\text{m s}^{-1}$. During the period when the system-wide CF from COAWST is < 0.2, the mean HH WS 530 in WT grid cells ranges from 27.2 to 29.3 m s⁻¹ and the maximum reaches 41.8 m s⁻¹. For all ocean-based grid cells within the four LA clusters, the mean HH WS as simulated by COAWST with the WFP active ranges from 23.1 to 27.0 m s⁻¹ and the maximum reaches 42.1 m s⁻¹.

Mean HH WS > 25 m s⁻¹ in WT grid cells and CF < 0.2 extend for 13.8, 13.3, and 7.2 h in the WRF simulation with the WFP active and 15.2, 12.8, and 7.3 h in the equivalent simulation with COAWST for LA clusters C, B, and A, respectively (Figs. 67 and 78). Due to slight differences in the hurricane tracking (Fig. 1a), the mean CF from COAWST exceeds that with WRF for offshore LA cluster B and in the northern part of LA cluster C, while the mean CF with WRF exceeds that with COAWST in the southern part of LA cluster C (Fig. 6a, b7a and Fig. SHS15). Hurricane Irene tracks very close to LA cluster D, which experiences mean HH WS in WT grid cells > 25 m s⁻¹ and CF < 0.2 for 15.5 and 17.3 h in the WRF and COAWST simulations, respectively. According to Mood's test, the median CF for this LA from the WRF and COAWST simulations differ at the 95% confidence level.

Nearly two-thirds of WT-containing grid cells and over three-quarters of ocean grid cells within the LA exhibit a higher frequency of HH WS > 25 m s⁻¹ in the COAWST simulations when Hurricane Irene is within d02. A larger R₁₈ value is also much more frequent (≥ 96% of time stamps) in each COAWST simulation. Thus, consistent with the analyses of precipitation volume from Hurricane Irene given above, there is evidence that the simulations with COAWST result in a more intense and larger tropical cyclone.

In the COAWST simulation with the WFP active, maximum Hs in WT grid cells within LA clusters A, B, C, and D is; 8.6, 8.5, 7.6, and 7.2 m, respectively (Fig. 89), and thus are all below the 50-year yr RP Hs of ~ 11 m estimated using the ERA5 dataset (Barthelmie et al., 2021). LA cluster A exhibits the highest frequency (~ 4%) of joint Hs, HH WS, and Tp values (Fig. 89) that fall in the classes centered at Hs ≥ 8.4 m, HH WS ≥ 35 m s⁻¹ (approximately equivalent to 5-m WS of 21.5 m s⁻¹), and Tp ≥ 11.2 s that were previously reported to be associated with high mudline bending moments based on simulations with 3D IFORM applied to the 5 MW NREL offshore reference wind-turbine WT (Valamanesh et al., 2015). The COAWST simulation also indicates frequent occurrence of wind-wave misalignment. In the HH WS class 10.6 -to 25 m s⁻¹, 47, 86, 74, and 32% of the time periods have wind-wave misalignment ≥ 30° at the center of LA clusters A, B, C, and D, respectively. For HH WS > 25 m s⁻¹, the corresponding values are 22, 41, 44, and 31%, respectively.



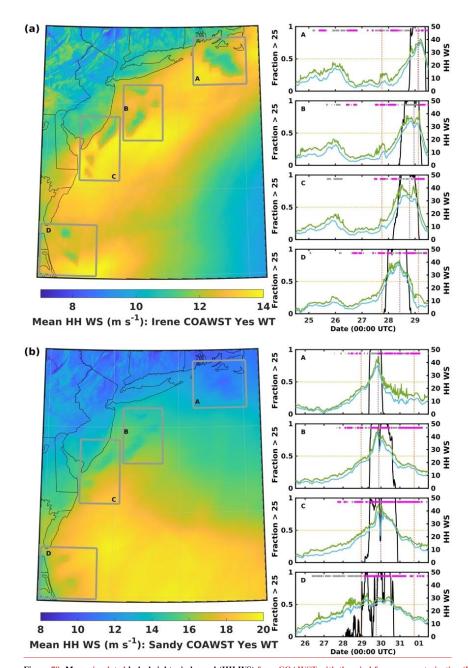


Figure 78: Mean simulated hub-height wind speed (HH WS) from COAWST with the wind farm parameterization (WFP) active for (a) Hurricane Irene and (b) Hurricane Sandy. The time series of show the fraction of wind turbine (WT) grid cells with HH WS > 25 m s⁻¹ (left axis, black) plus the mean (blue) and maximum (green) HH WS in those grid cells (right axis) in). Gray (magenta) symbols at the top of the time series indicate times when the mean HH WS from WRF and exceeds (is lower than) COAWST simulations with the WFP active by > 0.5 m s⁻¹. Orange dashed lines indicate the start and end time of storm tracking within d02. The purple dashed line represents the time when the location of the minimum sea level pressure (SLP) is closest to the lease area

570	(LA) cluster center. The yellow line indicates HH WS of= 25 m s ⁻¹ . For simulations without the WFP active, see Fig. S10. See also Figs. S13 and S14.	Formatted: Font:	Times New Roman,	9 pt, Ligatures: None
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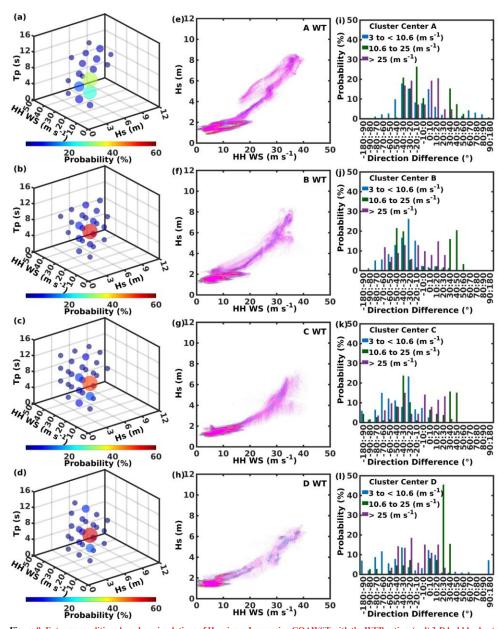


Figure 8: Extreme conditions based on simulations of Hurricane Irene using COAWST with the WFP active. (a-d) 3-D bubble charts of the joint 9: Joint occurrence of (a-d) hub-height wind speed (HH WS $_{7}$), significant wave height (Hs $_{7}$), and peak wave period (Tp) (5 classes for each variable for a total of 125 possible classes) for all wind in COAWST simulations of Hurricane Irene with wind farm parameterization (WFP) active for wind turbine (WT) grid cells in each lease area (LA) cluster (A-D).-(e-h) Joint probability distributions of HH WS and Hs where the magenta symbols denote 10-min output from all WT grid cells in each LA cluster and the contours denote probabilityprobabilities of 0.01 (blue), 0.02 (green), and 0.05 (yellow). (i-l) Histograms of the Directional misalignment (hub-height direction difference (HH WS-minus Hsmean wave direction) at the center of each LA cluster for the three HH WS classes: 3 to < 10.6 m s⁻¹, 10.6 to 25 m s⁻¹, and > 25 m s⁻¹.

3.3 Wind turbine power production and operating conditions:

3.2.2 Hurricane Sandy

Mean instantaneous power production and CF from WRF and COAWST for the entire-Hurricane Sandy simulation period are; 2.03×10^4 MW (0.51) and 2.07×10^4 MW (0.52), respectively. Considering only the time when When Hurricane Sandy is within d02, equivalent CFCFs are 0.62 and 0.61, respectively. The high CFCFs are reflective of high, but below cut-out, HH WS prior to the passage of the hurricane over the LALAs and the relatively short duration of HH WS > 25 m s⁻¹ within the LALAs (Fig. 5b6b and Fig. 6e, d7b). Simulated system-wide CF drops below 0.2 for 8 h (scattered during 1320 UTC-29 October 13:20 UTC through 0100 UTC-30 October 01:00 UTC) in the WRF simulation and for 15 h (1010 UTC-29 October 10:10 UTC through 0110 UTC-30 October 01:10 UTC) in the COAWST simulation (Fig. 5b6b). A single time-step (2100 UTC-29 October 21:00 UTC) has zero system-wide power production in the COAWST simulation.

During periods when When the system-wide CF < 0.2 (including landfall in New Jersey), the mean HH WS in WT grid cells is > 25 m s⁻¹ in both the WRF and COAWST simulations with the WFP active (Fig. 5b). In the WRF simulation with the WFP active, during6b). During the longest continuous time when the system-wide CF from WRF remains below 0.2 (471029 October 17:10 through 204020:40 UTC-29 October), the mean HH WS is 30.6 to 34.8 m s⁻¹ in WT grid cells and 29.7 to 31.9 m s⁻¹ in all ocean-based grid cells within the LA clusters. Equivalent values from COAWST (also with WFP active), are 26.0 to 38.5 m s⁻¹ and 25.8 to 35.0 m s⁻¹.

HH WS > 50 m s⁻¹ is simulated in tens of thousands of space-time sample combinations in both the WRF and COAWST simulations with the WFP active. However, none occurred within 170 km of any LA centroid. Maximum HH WS in WT grid cells and the frequency of HH WS > 25 m s⁻¹ in WT grid cells is higher (59% and 65% of time stamps when Hurricane Sandy is within d02) in the COAWST simulation than in the WRF simulation with and without the WFP active (Fig. 78). In all ocean-based grid cells within the LA clusters, 63% and 69% of the time stamps exhibit more grid cells with HH WS > 25 m s⁻¹ in the COAWST simulations. Maximum HH WS in ocean-based grid cells within the LA clusters is 45.1 m s⁻¹ in WRF and 48.9 m s⁻¹ in COAWST (Fig. 78). In the COAWST simulation with the WFP active, the maximum HH WS in WT grid cells in LA clusters A, B, and C are; 44.6, 47.7, and 45.4 m s⁻¹, respectively (Fig. 910). They thus exceed the highest 50-year yr RP wind speedWS at 100 m a.s.l. of 39.7 m s⁻¹ computed using ERA5 output (Barthelmie et al., 2021), but are below the 50 m s⁻¹ and 57 m s⁻¹ thresholds for class I and class T wind turbines WTs (IEC, 2019a). Larger R₁₈ values prior to landfall are also more frequent in the COAWST simulations (> 70% of time stamps in both the simulations without and with the WFP). Thus, consistent with analyses of the simulations of Hurricane Irene, there is evidence that use of COAWST (for the configuration used herein) results in a larger and more intense hurricane.

Minor differences in the tracking (Fig. 1b) and intensity of the hurricane-induced wind-speedsWS (Fig. 78), causes higher mean CF from LA cluster A and parts of B and C in the simulation with COAWST than the simulation with WRF (Fig. 67 and Fig. S11S15). The simulations track the centroid of Hurricane Sandy close to LA clusters B and C and accordingly, periods with CF < 0.2 are of the greatest duration for these clusters (20.0 and 23.5 h in WRF and 23.3 and 31.0 h in COAWST, respectively) (Fig. 6). Largest 7). The largest differences in CF are found for LA cluster D. The duration of time with CF < 0.2 is substantially longer in the COAWST simulation due to the prevalence of HH WS > 25 m s⁻¹ and the median CF for this cluster between the two simulations differs at the 99% confidence level according to Mood's test.

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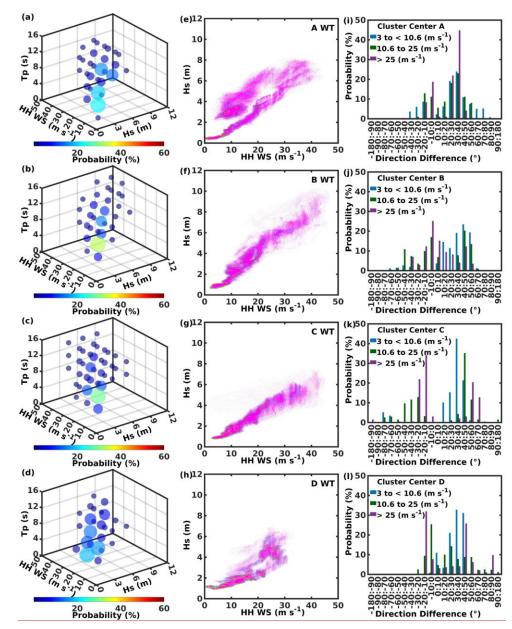


Figure 10: Joint occurrence of (a-d) hub-height wind speed (HH WS), significant wave height (Hs), and peak wave period (Tp) (5 classes for each variable for a total of 125 possible classes) in COAWST simulations of Hurricane Sandy with wind farm parameterization (WFP) active for wind turbine (WT) grid cells in each lease area (LA) cluster (A-D). (e-h) Joint probability distributions of HH WS and Hs where the magenta symbols denote 10 min output and the contours denote probabilities of 0.01 (blue), 0.02 (green), and 0.05 (yellow). (i-l) Directional misalignment (hub-height direction minus mean wave direction) at the center of each LA cluster for three HH WS classes: 3 to < 10.6 m s⁻¹, 10.6 to 25 m s⁻¹, and > 25 m s⁻¹.

Maximum simulated Hs of 8.3, 10.4, 7.5, and 6.9 m in LA clusters A, B, C, and D (Fig. 10) are higher than those for Hurricane

Irene but are also below the 50-year yr RP values of ~ 11 m derived from ERA5 (Barthelmie et al., 2021) (Fig. 9). LA cluster

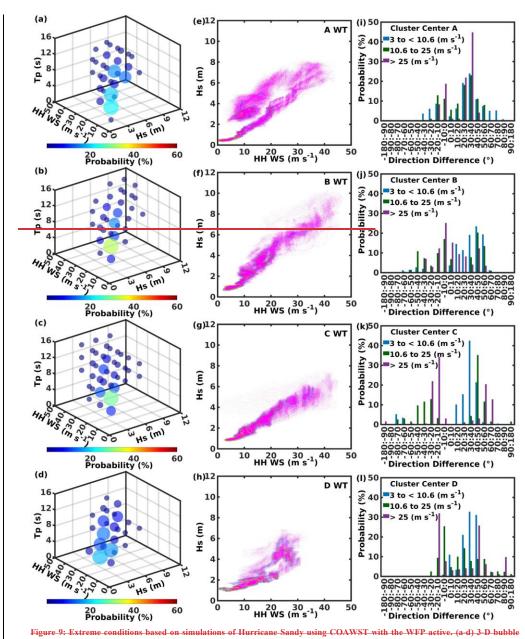
B exhibits the highest frequency (~ 4%) of joint Hs, HH WS, and Tp values (classes centered at Hs ≥ 8.4 m, HH WS ≥ 35 m

s⁻¹ [5-m WS ~ 21.5 m s⁻¹], and Tp ≥ 11.2 s) close to those associated with a peak mudline moment (of ~ 120 MN-m)

(Valamanesh et al., 2015). MisalignmentWind-wave misalignment of wind and waves by ≥ 30° is common in the COAWST

simulation for both the HH WS elassclasses of 10.6 -to 25 m s⁻¹ and > 25 m s⁻¹. Based on COAWST output from the centroids

of LA A, B, C, and D, wind-wave misalignment ≥ 30° is found for; 43 (49), 63 (27), 83 (41) and 34 (49) % of time steps (value in brackets for WS HH > 25 m s⁻¹).



that's of the joint occurrence of HH WS, Hs, and Tp (5 classes for each variable for a total of 125 possible classes) for all wind turbine grid cells in each LA cluster (A-D). (e-h) Joint probability distributions of HH WS and Hs where the magenta symbols denote 10-min output from all WT grid cells in each LA cluster and the contours denote probability of 0.01 (blue), 0.02 (green), and 0.05 (yellow). (i-l) Histograms of the direction difference (HH WS minus Hs) at the center of each LA cluster for the three HH WS classes: 3-to <10.6 m s⁻¹, 10.6 to 25 m s⁻¹, and > 25 m s⁻¹.

655 3.43.3 Wind turbine impacts on hurricane properties

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Consistent with expectations, removal of kinetic energy by wind turbines WTs means that maximum HH WS in WRF and COAWST simulations with the WFP active tend to be lower than those when the wind turbines WTs are excluded (see Fig. 56). For example, output from the no wind turbine WT COAWST simulation of Hurricane Irene indicates HH WS in oceanbased grid cells beyond the LALAs > 50 m s⁻¹ 837 times while the corresponding number for the simulation with wind turbinesWTs is 333. However, simulations using both WRF and COAWST with full deployment of wind turbinesWTs in existing lease areas (approximately 2600 at an ICD of 4.3 MW km⁻²); LAs, indicate that for this scale of offshore wind turbineWT deployment, the net impact is small expectexcept for hub-height wind speedsHH WS near the lease areasLAs.

For all other metrics, the simulations with WRF or COAWST differ more than the simulations with or without the action of wind turbines WTs included (Fig. 1011). The 10-min mean precipitation volume within 375 km of the Hurricane Irene centroid differs (without WT versus with WFP active) by 4.7% and 3.9% in simulations with WRF and COAWST, respectively, but differs (WRF versus COAWST) by 11.8% and 9.2% in the no WT and WT simulations, respectively. The equivalent values for Hurricane Sandy are 8.7% for WRF and 12.9% for COAWST (without WFP active versus with WFP active) and 13.0% for the no WT and 11.6% for the WT simulations (WRF versus COAWST). Similarly, the mean 500 hPa wind speedWS close to the hurricane centroids differ by < 2 m s⁻¹ in simulations of Hurricanes Irene and Sandy with and without WTWTs with no consistent signal in terms of which simulation is higher (Fig. $\frac{1011}{1000}$). The mean R_{18} for Hurricane Irene based on WRF simulations with and without the WFP active differ by < 2 km (mean of ~ 280 km). Mean R₁₈ from COAWST simulations with and without the WFP active also differ by < 2 km (mean of ~ 300 km). Thus, while median R_{18} from WRF versus COAWST are statistically different (at p < 0.01) for simulations with and without the WFP active, the use of the WFP does not yield significantly different R₁₈ values in simulations with a given model. Simulations with WFP active produce equal or slightly more total precipitation. For Hurricane Irene, the 10-min mean (median) precipitation volume from WRF without and with the WFP active are 1.88×10^8 m³ (1.89×10^8 m³) and 1.88×10^8 m³ (1.93×10^8 m³), while equivalent values from COAWST are 1.93×10^8 m³ (1.93×10^8 m³) and 1.98×10^8 m³ (2.02×10^8 m³), respectively. For Hurricane Sandy, the corresponding values are $8.64 \times 10^7 \text{ m}^3 \ (8.60 \times 10^7 \text{ m}^3) \text{ versus } 8.68 \times 10^7 \text{ m}^3 \ (8.10 \times 10^7 \text{ m}^3) \text{ from WRF and } 8.69 \times 10^7 \text{ m}^3 \ (8.10 \times 10^7 \text{ m}^3) \text{ from WRF}$ 680 $(9.19 \times 10^7 \text{ m}^3)$ versus $8.87 \times 10^7 \text{ m}^3$ $(7.79 \times 10^7 \text{ m}^3)$ from COAWST.

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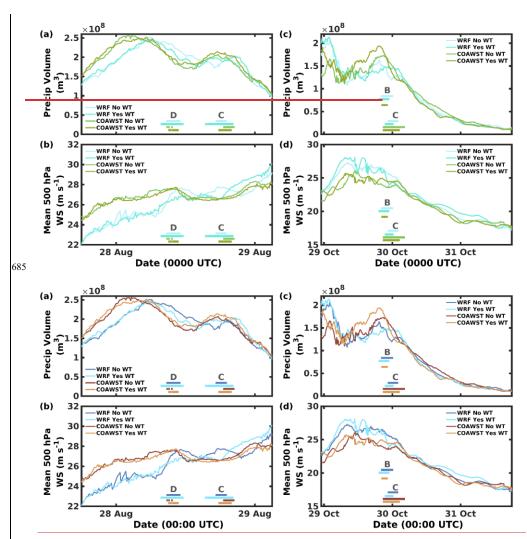


Figure 1011: Time series of (a, c) 10-min precipitation volume within a 375 km radius from the minimum sea level pressure (SLP location) and (b, d) mean 500 hPa wind speed of grid-cells from(WS) 50 to 375 km from the minimum SLP location in d02 for from simulations with WRF and COAWST of (a, b) Hurricane Irene and (c, d) Hurricane Sandy. The colored horizontal lines near "P" and "C" in panels a and b and near "B" and "C" for panels c and d, mark thelabels denote times when the minimum SLP is within 100 km of the center of the specified offshore wind energy lease area (LA) cluster centroid. See also Fig. 812 for comparison of wind speeds at 10-m a.s.1516_a

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4 Concluding remarks

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Results of analyses Analyses of simulations with WRF and COAWST of two of the most powerful hurricanes that tracked within 100 km of offshore wind energy lease areas LAs along the U.S. East Coast during the last 25 years can be summarized as:

1) Many aspects of Hurricane Sandy are well reproduced in WRF and COAWST control simulations that exclude the action of wind turbines. WTs. Consistent with past research, simulations of Hurricane Irene exhibit lower fidelity relative to a range of observations in part due to the negative bias in translational speed. In contrast to similar simulations of Typhoon Muifa (Liu et al., 2015) but consistent with past research on intense cyclones in the North Sea (Larsén et al., 2019), COAWST simulations of both hurricanes indicate generally better agreement with $\underline{observations,} \ a \ slightly \ larger \ area \ of \ storm-force \ \underline{wind \ speedsWS} \ (R_{18}) \ and \ \underline{hub \ height \ wind \ speedsHH \ WS} > 25 \ m$ s⁻¹ plus higher precipitation volumes than are indicated by the WRF-simulations. This coupled with the ability of COAWST to quantify additional geophysical parameters of importance to offshore structures strongly indicates the need for increasing investment in coupled simulations for the offshore wind energy industry.

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- 2) Despite the intensity and size of these hurricanes and their proximity to the offshore wind energy lease areas LAs, simulations presented herein, that assume a 15 MW reference wind turbine WT deployed with a spacing of 1.85 km, indicate only fairly-brief periods with low power production (system-wide CF < 0.2). System-wide capacity factors below CF < 0.2 due to wide-spread occurrence of hub height wind speeds above HH WS > 25 m s⁻¹ extend for only 6to 7 h in the simulations of Hurricane Irene and 8-to 15 h for Hurricane Sandy (the longer period is based on the COAWST simulation with COAWST). Further, neither hurricane is simulated to produce hub height wind speedsHH WS > 50 m s⁻¹ in the current offshore lease areasLAs. Thus, based on these simulations of these intense tropical eycloneshurricanes there is no evidence of a need for hurricane hardening of wind turbinesWTs deployed in the current offshore lease areas these LAs. Also, these simulations even suggest even such that the projected fleet of offshore wind turbines WTs will continue to supply substantial amounts of electricity to the grid even during these extreme events. However, simulations of both hurricanes with COAWST result in wave-wind conditions that have previously been identified as being associated with high mudline bending moments on monopile foundations.-The COAWST simulations of both hurricanes also indicate and a relativerelatively high frequency of HH WSwind-wave directional misalignment (> 30°) in these lease areasLAs.
- 3) There is no evidence that deployment of 2642 wind turbines WTs at an ICD of 4.3 MW km⁻² within existing offshore wind energy lease areas LAs along the U.S. East Coast would have substantially weakened either of the hurricanes considered herein. Although much denser and larger deployments might have an influence on hurricanes, even for Hurricane Irene that tracked closest to the offshore wind energy lease areasthese LAs, simulations with either WRF or COAWST differ more than simulations with either WFP inactive or active with respect to the volume of precipitation near the hurricane center, storm intensity, and/or extent.

Mesoscale simulations performed at convection permitting resolution such as those presented herein allow simulation of the hurricane lifespans and associated power production over large domains and can be used as here to assess whether improved treatment of atmosphere-ocean dynamical coupling alters extreme conditions of relevance to offshore wind turbines. WTs. However, it is important to acknowledge that the highest structural loading may occur in the cyclone eye-wall (Han et al., 2014) that is of a scale (Marks et al., 2008) that is not fully represented in the simulations presented here. Nevertheless, analyses of the simulations suggest the structure of the hurricanes is reasonably represented in our modeling (Fig. 34 and Fig. S4) and simulations performed at the same grid spacing were shown to represent some aspects of flow in the eve wall (Müller et al., 2024). Future work employing mesoscale-microscale coupling (Wang et al., 2024) could be used to evolve further details of geophysical properties of relevance to structural loading. Further, the hurricanes simulated herein were extremely powerful 740 and both tracked within 100 km of offshore wind energy lease areaLA cluster centers (C and D for Hurricane Irene, B and C for Hurricane Sandy). However, they do not represent a comprehensive climatology of historical or possible intense

tropical/extratropical cyclones (Barthelmie et al., 2021). Undertaking comparable simulations of additional extreme cyclones and simulations with different configurations including alternative z₀ parameterizations (Porchetta et al., 2019; Fu et al., 2023) and a wave boundary layer model within SWAN (Du et al., 2017) would also be useful in determining if findings presented herein are generalizable and to quantify the degree to which the meteorological and oceanic extreme conditions vary according to the precise model formulation.

Code and data availability

COAWST software can be downloaded from: https://github.com/DOI-USGS/COAWST. NAM data can be downloaded from: https://rda.ucar.edu/datasets/d609000/ and https://www.ncei.noaa.gov/products/weather-climate-models/north-americanmesoscale. OSTIA-UKMO-L4-GLOB-v2.0 SST data can be downloaded from: https://podaac.jpl.nasa.gov/dataset/OSTIA-<u>UKMO-L4-GLOB-v2.0</u>. GHSHHG data can be downloaded from: https://www.ngdc.noaa.gov/mgg/shorelines/. GEBCO data can be downloaded from: https://download.gebco.net/. HYCOM GLBa0.08 expt 90.9 data can be downloaded from: $\underline{https://tds.hycom.org/thredds/catalog.html}.$ ADCIRC 2001v2d data downloaded from: https://adcirc.org/products/adcirc-tidal-databases/. GFS wind forcing data be downloaded can 755 https://www.ncei.noaa.gov/thredds/catalog/model/gfs.html. WW3 data for boundary conditions can be downloaded from: https://www.ncci.noaa.gov/thredds-ocean/catalog/ncep/nww3/catalog.html. NHC "best track" data can be downloaded from: HURDAT2 https://www.nhc.noaa.gov/data/tcr/index.php. data can downloaded from: https://www.nhc.noaa.gov/data/#hurdat. **IMERG** data be downloaded from: can https://disc.gsfc.nasa.gov/datasets/GPM_3IMERGHH_07/summary?keywords="IMERG_final". NDBC buoy data can be downloaded from: https://www.ndbc.noaa.gov/. Scientific color downloaded maps can https://www.fabiocrameri.ch/colourmaps/ (Crameri et al., 2020). Namelist information can be found at https://zenodo.org/records/14895862.

Supplemental materials

See the attached document.

765 Author contributions

SCP and RJB obtained funding and computing support for this work. KBT performed the simulations, SCP and KBT processed data used for model evaluation, conducted the analyses with assistance from RJB, produced the figures, and drafted the initial manuscript. All authors contributed to the design of the analyses, interpretation of results, and manuscript revisions.

Competing interests

The authors declare that they have no conflict of interest.

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780 References

Avila, L. A. and Cangialosi, J.: Tropical Cyclone Report: Hurricane Irene, 45, https://www.nhc.noaa.gov/data/tcr/AL092011 Irene.pdf, 2011.

Baldini, L. M., Baldini, J. U. L., McElwaine, J. N., Frappier, A. B., Asmerom, Y., Liu, K.-b., Prufer, K. M., Ridley, H. E., Polyak, V., Kennett, D. J., Macpherson, C. G., Aquino, V. V., Awe, J., and Breitenbach, S. F. M.: Persistent northward North Atlantic tropical cyclone track migration over the past five centuries, Scientific Reports, 6, 37522, https://www.nature.com/articles/srep37522, 2016.

- Barthelmie, R. J., Dantuono, K. E., Renner, E. J., Letson, F. L., and Pryor, S. C.: Extreme Wind and Waves in U.S. East Coast Offshore Wind Energy Lease Areas, Energies, 14, 1053, https://www.mdpi.com/1996-1073/14/4/1053, 2021.
- Blake, E. S., Kimberlain, T. B., Berg, R. J., Cangialosi, J. P., and Beven II, J. L.: Tropical Cyclone Report: Hurricane Sandy, National Hurricane Center, 157, https://www.nhc.noaa.gov/data/tcr/AL182012 Sandy.pdf, 2013.
- Bodini, N., Optis, M., Redfern, S., Rosencrans, D., Rybchuck, A., Lundquist, J., Pronk, V., Castagneri, S., Purkayastha, A., Draxl, C., Krishnamurthy, R., Young, E., Roberts, B., Rosenlieb, E., and Musial, W.: The 2023 National Offshore Wind Data Set (NOW-23), Earth System Science Data, 16, 1965-2006, https://research-hub.nrel.gov/en/publications/the-2023-national-offshore-wind-data-set-now-23, 2024.
- Pooij, N., Ris, R. C., and Holthuijsen, L. H.: A third-generation wave model for coastal regions: 1. Model description and validation, Journal of Geophysical Research: Oceans, 104, 7649-7666, https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/98JC02622, 1999.
- Calvo-Sancho, C., Quitián-Hernández, L., González-Alemán, J. J., Bolgiani, P., Santos-Muñoz, D., and Martín, M. L.: Assessing the performance of the HARMONIE-AROME and WRF-ARW numerical models in North Atlantic Tropical
 Transitions, Atmospheric Research, 291, 106801, https://www.sciencedirect.com/science/article/pii/S0169809523001989, 2023
 - Chen, F. and Dudhia, J.: Coupling an Advanced Land Surface—Hydrology Model with the Penn State—NCAR MM5 Modeling System. Part I: Model Implementation and Sensitivity, Monthly Weather Review, 129, 569-585, https://journals.ametsoc.org/view/journals/mwre/129/4/1520-0493 2001 129 0569 caalsh 2.0.co 2.xml, 2001a.
- 805 Chen, F. and Dudhia, J.: Coupling an Advanced Land Surface–Hydrology Model with the Penn State–NCAR MM5 Modeling System. Part II: Preliminary Model Validation, Monthly Weather Review, 129, 587-604, https://journals.ametsoc.org/view/journals/mwre/129/4/1520-0493 2001 129 0587 caalsh 2.0.co 2.xml, 2001b.
- Chen, X. and Xu, J. Z.: Structural failure analysis of wind turbines impacted by super typhoon Usagi, Engineering Failure Analysis, 60, 391-404, https://www.sciencedirect.com/science/article/pii/S1350630715301552?via%3Dihub, 2016.
- [810] Chen, Y. and Yau, M. K.: Asymmetric Structures in a Simulated Landfalling Hrricane, Journal of the Atmospheric Sciences, 60, 2294-2312, https://journals.ametsoc.org/view/journals/atsc/60/18/1520-0469_2003_060_2294_asiasl_2.0.co_2.xml, 2003.
- Crameri, F., Shephard, G. E., and Heron, P. J.: The misuse of colour in science communication, Nature Communications, 11, 5444, https://www.nature.com/articles/s41467-020-19160-7, 2020.
- Davis, C., Wang, W., Chen, S. S., Chen, Y., Corbosiero, K., DeMaria, M., Dudhia, J., Holland, G., Klemp, J., Michalakes, J., Reeves, H., Rotunno, R., Snyder, C., and Xiao, Q.: Prediction of Landfalling Hurricanes with the Advanced Hurricane WRF Model, Monthly Weather Review, 136, 1990-2005, https://journals.ametsoc.org/view/journals/mwre/136/6/2007mwr2085.1.xml, 2008.

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Drennan, W. M., Taylor, P. K., and Yelland, M. J.: Parameterizing the Sea Surface Roughness, Journal of Physical Oceanography, 35, 835-848, https://journals.ametsoc.org/view/journals/phoc/35/5/jpo2704.1.xml, 2005.

Du, J., Bolaños, R., and Guo Larsén, X.: The use of a wave boundary layer model in SWAN, Journal of Geophysical Research: Oceans, 122, 42-62, https://agupubs.onlinelibrary.wiley.com/doi/full/10.1002/2016JC012104, 2017.

Dudhia, J.: Numerical Study of Convection Observed during the Winter Monsoon Experiment Using a Mesoscale TwoDimensional Model, Journal of the Atmospheric Sciences, 46, 3077-3107,
https://journals.ametsoc.org/view/journals/atsc/46/20/1520-0469 1989 046 3077 nsocod 2 0 co 2.xml, 1989.

825

855

Ek, M. B., Mitchell, K. E., Lin, Y., Rogers, E., Grunmann, P., Koren, V., Gayno, G., and Tarpley, J. D.: Implementation of Noah land surface model advances in the National Centers for Environmental Prediction operational mesoscale Eta model, Journal of Geophysical Research: Atmospheres, 108, 8851, https://agupubs.onlinelibrary.wiley.com/doi/full/10.1029/2002JD003296, 2003.

Fischer, T., Rainey, P., Bossanyi, E., and Kühn, M.: Study on Control Concepts suitable for Mitigation of Loads from Misaligned Wind and Waves on Offshore Wind Turbines supported on Monopiles, Wind Engineering, 35, 561-573, https://journals.sagepub.com/doi/10.1260/0309-524X.35.5.561, 2011.

Fitch, A. C., Olson, J. B., Lundquist, J. K., Dudhia, J., Gupta, A. K., Michalakes, J., and Barstad, I.: Local and Mesoscale Impacts of Wind Farms as Parameterized in a Mesoscale NWP Model, Monthly Weather Review, 140, 3017-3038, https://journals.ametsoc.org/view/journals/mwre/140/9/mwr-d-11-00352.1.xml, 2012.

Foody, R., Coburn, J., Aird, J. A., Barthelmie, R. J., and Pryor, S. C.: Quantitative comparison of power production and power quality onshore and offshore: a case study from the eastern United States, Wind Energy Science, 9, 263-280, https://wes.copernicus.org/articles/9/263/2024/wes-9-263-2024.pdf, 2024.

Fu, S., Huang, W., Luo, J., Yang, Z., Fu, H., Luo, Y., and Wang, B.: Deep Learning-Based Sea Surface Roughness

840 Parameterization Scheme Improves Sea Surface Wind Forecast, Geophysical Research Letters, 50, e2023GL106580, https://agupubs.onlinelibrary.wiley.com/doi/10.1029/2023GL106580, 2023.

Gaertner, E., Rinker, J., Sethuraman, L., Zahle, F., Anderson, B., Barter, G., Abbas, N., Meng, F., Bortolotti, P., Skrzypinski, W., Scott, G., Feil, R., Bredmose, H., Dykes, K., Shields, M., Allen, C., and Viselli, A.: IEA Wind TCP Task 37: Definition of the IEA 15-Megawatt Offshore Reference Wind Turbine, National Renewable Energy Lab.(NREL), Golden, CO (United States), 54, https://www.nrel.gov/docs/fy20osti/75698.pdf, 2020.

Gaudet, B. J., García Medina, G., Krishnamurthy, R., Shaw, W. J., Sheridan, L. M., Yang, Z., Newsom, R. K., and Pekour, M.: Evaluation of Coupled Wind–Wave Model Simulations of Offshore Winds in the Mid-Atlantic Bight Using Lidar-Equipped Buoys, Monthly Weather Review, 150, 1377-1395, https://journals.ametsoc.org/view/journals/mwre/150/6/MWR-D-21-0166.1.xml, 2022.

850 GWEC: Global Wind Report 2024, Global Wind Energy Council, Brussels, Belgium., 168, https://gwec.net/global-wind-report-2024/, 2024a.

GWEC: Global Offshore Wind Report 2024, Global Wind Energy Council, Brussels, Belgium., 156, https://gwec.net/global-40ffshore-wind-report-2024/2024ba.2024a.

GWEC: Global Wind Report 2024, Global Wind Energy Council, Brussels, Belgium., 168, https://gwec.net/global-wind-report-2024/, 2024b.

Haidvogel, D. B., Arango, H., Budgell, W. P., Cornuelle, B. D., Curchitser, E., Di Lorenzo, E., Fennel, K., Geyer, W. R., Hermann, A. J., Lanerolle, L., Levin, J., McWilliams, J. C., Miller, A. J., Moore, A. M., Powell, T. M., Shchepetkin, A. F., Sherwood, C. R., Signell, R. P., Warner, J. C., and Wilkin, J.: Ocean forecasting in terrain-following coordinates: Formulation and skill assessment of the Regional Ocean Modeling System, Journal of Computational Physics, 227, 3595-3624, https://www.sciencedirect.com/science/article/abs/pii/S0021999107002549, 2008.

Hallowell, S. T., Myers, A. T., Arwade, S. R., Pang, W., Rawal, P., Hines, E. M., Hajjar, J. F., Qiao, C., Valamanesh, V., Wei, K., Carswell, W., and Fontana, C. M.: Hurricane risk assessment of offshore wind turbines, Renewable Energy, 125, 234-249, https://www.sciencedirect.com/science/article/pii/S0960148118302349, 2018.

Han, T., McCann, G., Mücke, T. A., and Freudenreich, K.: How can a wind turbine survive in tropical cyclone?, Renewable energy, 70, 3-10, https://www.sciencedirect.com/science/article/pii/S0960148114000949, 2014.

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- Hansen, T. A., Wilson, E. J., Fitts, J. P., Jansen, M., Beiter, P., Steffen, B., Xu, B., Guillet, J., Münster, M., and Kitzing, L.: Five grand challenges of offshore wind financing in the United States, Energy Research & Social Science, 107, 103329, https://www.sciencedirect.com/science/article/pii/S2214629623003894, 2024.
- Hashemi, M. R., Kresning, B., Hashemi, J., and Ginis, I.: Assessment of hurricane generated loads on offshore wind farms; a closer look at most extreme historical hurricanes in New England, Renewable Energy, 175, 593-609, https://www.sciencedirect.com/science/article/pii/S0960148121007187, 2021.
 - Hettmansperger, T. P. and Malin, J. S.: A modified Mood's test for location with no shape assumptions on the underlying distributions, Biometrika, 62, 527-529, https://www.jstor.org/stable/2335398?seq=1, 1975.
- Hong, S.-Y. and Lim, J.-O. J.: The WRF single-moment 6-class microphysics scheme (WSM6), Journal of the Korean Meteorological Society, 42, 129-151, https://www.researchgate.net/publication/331192569 Hongandlim-JKMS-2006, 2006.
 - Hu, G. and Franzke, C. L. E.: Evaluation of Daily Precipitation Extremes in Reanalysis and Gridded Observation-Based Data Sets Over Germany, Geophysical Research Letters, 47, e2020GL08962, https://agupubs.onlinelibrary.wiley.com/doi/10.1029/2020GL089624, 2020.
- Huffman, G. J., Bolvin, D. T., Joyce, R., Kelley, O. A., Nelkin, E. J., Portier, A., Stocker, E. F., Tan, J., Watters, D. C., and West, B. J.: IMERG V07 Release Notes, 17, https://gpm.nasa.gov/sites/default/files/2024-02/IMERG_V07_ReleaseNotes_240221.pdf, 2024.
 - IEC: IEC 61400-3-1:2019. Wind energy generation systems. Part 3–1: Design requirements for fixed offshore wind turbines, International Electrotechnical Commission, Geneva, Switzerland. ISBN: 9780580944345, Geneva, Switzerland, 2019a.
- IEC: IEC 61400-1 Edition 4.0 2019-02 Wind turbines Part 1: Design requirements, International Electrotechnical Commission, Geneva, Switzerland. ISBN 978-2-8322-6253-5, Geneva, Switzerland, 2019b.
 - Jacob, R., Larson, J., and Ong, E.: M × N Communication and Parallel Interpolation in CCSM Using the Model Coupling Toolkit, 33, https://www.mcs.anl.gov/research/projects/acpi/mct/mxnP1225.pdf, 2005.
- Jacobson, M. Z., Archer, C. L., and Kempton, W.: Taming hurricanes with arrays of offshore wind turbines, Nature Climate Change, 4, 195-200, https://www.nature.com/articles/nclimate2120, 2014.
- Jansen, M., Staffell, I., Kitzing, L., Quoilin, S., Wiggelinkhuizen, E., Bulder, B., Riepin, I., and Müsgens, F.: Offshore wind competitiveness in mature markets without subsidy, Nature Energy, 5, 614-622, https://www.nature.com/articles/s41560-020-0661-2, 2020.
 - Ju, S.-H., Hsu, H.-H., and Hsiao, T.-Y.: Three-dimensional wind fields of tropical cyclones for wind turbine structures, Ocean Engineering, 237, 109437, https://www.sciencedirect.com/science/article/pii/S0029801821008428, 2021.
- Kain, J. S.: The Kain–Fritsch Convective Parameterization: An Update, Journal of Applied Meteorology, 43, 170-181, https://journals.ametsoc.org/view/journals/apme/43/1/1520-0450 2004 043 0170 tkcpau 2.0.co 2.xml, 2004.
- Khaira, U. and Astitha, M.: Exploring the Real-Time WRF Forecast Skill for Four Tropical Storms, Isaias, Henri, Elsa and Irene, as They Impacted the Northeast United States, Remote Sensing, 15, 3219, https://www.mdpi.com/2072-4292/15/13/3219, 2023.
- [900] Koukoura, C., Brown, C., Natarajan, A., and Vesth, A.: Cross-wind fatigue analysis of a full scale offshore wind turbine in the case of wind–wave misalignment, Engineering Structures, 120, 147-157, https://www.sciencedirect.com/science/article/pii/S0141029616301389, 2016.
- Kresning, B., Hashemi, M. R., and Gallucci, C.: Simulation of Hurricane Loading for Proposed Offshore Windfarms off the US Northeast Coast, Journal of Physics: Conference Series, 12 pp, https://iopscience.iop.org/article/10.1088/1742-905 6596/1452/1/012026, 2020.
- Kresning, B., Hashemi, M. R., Shirvani, A., and Hashemi, J.: Uncertainty of extreme wind and wave loads for marine renewable energy farms in hurricane-prone regions, Renewable Energy, 220, 119570, https://www.sciencedirect.com/science/article/pii/S0960148123014854, 2024.

- Lackmann, G. M.: Hurricane Sandy before 1900 and after 2100, Bulletin of the American Meteorological Society, 96, 547-910 560, https://journals.ametsoc.org/view/journals/bams/96/4/bams-d-14-00123.1.xml, 2015.
- Landsea, C. W. and Franklin, J. L.: Atlantic Hurricane Database Uncertainty and Presentation of a New Database Format, Monthly Weather Review, 141, 3576-3592, https://journals.ametsoc.org/view/journals/mwre/141/10/mwr-d-12-00254.1.xml, 2013.
- Larsén, X. G. and Ott, S.: Adjusted spectral correction method for calculating extreme winds in tropical-cyclone-affected water areas, Wind Energy Science, 7, 2457-2468, https://wes.copernicus.org/articles/7/2457/2022/, 2022.
- Larsén, X. G., Ott, S., Badger, J., Hahmann, A. N., and Mann, J.: Recipes for Correcting the Impact of Effective Mesoscale Resolution on the Estimation of Extreme Winds, Journal of Applied Meteorology and Climatology, 51, 521-533, https://journals.ametsoc.org/view/journals/apme/51/3/jamc-d-11-090.1.xml, 2012.
- Larsén, X. G., Du, J., Bolaños, R., Imberger, M., Kelly, M. C., Badger, M., and Larsen, S.: Estimation of offshore extreme wind from wind-wave coupled modeling, Wind Energy, 22, 1043–1057, https://onlinelibrary.wiley.com/doi/epdf/10.1002/we.2339, 2019.
- Larson, J., Jacob, R., and Ong, E.: The Model Coupling Toolkit: A New Fortran90 Toolkit for Building Multiphysics Parallel Coupled Models, The International Journal of High Performance Computing Applications, 19, 277-292, https://journals.sagepub.com/doi/10.1177/1094342005056115, 2005.
- 925 Li, J., Li, Z., Jiang, Y., and Tang, Y.: Typhoon Resistance Analysis of Offshore Wind Turbines: A Review, Atmosphere, 13, 451, https://www.mdpi.com/2073-4433/13/3/451, 2022.
 - Li, S. and Chen, C.: Air-sea interaction processes during hurricane Sandy: Coupled WRF-FVCOM model simulations, Progress in Oceanography, 206, 102855, https://www.sciencedirect.com/science/article/pii/S0079661122001148, 2022.
- Liu, N., Ling, T., Wang, H., Zhang, Y., Gao, Z., and Wang, Y.: Numerical Simulation of Typhoon Muifa (2011) Using a Coupled Ocean-Atmosphere-Wave-Sediment Transport (COAWST) Modeling System, Journal of Ocean University of China, 14, 199-209, https://link.springer.com/article/10.1007/s11802-015-2415-5, 2015.
- Marks, F. D., Black, P. G., Montgomery, M. T., and Burpee, R. W.: Structure of the Eye and Eyewall of Hurricane Hugo (1989), Monthly Weather Review, 136, 1237-1259, https://journals.ametsoc.org/view/journals/mwre/136/4/2007mwr2073.1.xml, 2008.
- Martín del Campo, J. O., Pozos-Estrada, A., and Pozos-Estrada, O.: Development of fragility curves of land-based wind turbines with tuned mass dampers under cyclone and seismic loading, Wind Energy, 24, 737-753, https://onlinelibrary.wiley.com/doi/full/10.1002/we.2600, 2021.
- Marvel, K., Kravitz, B., and Caldeira, K.: Geophysical limits to global wind power, Nature Climate Change, 3, 118-121, https://www.nature.com/articles/nclimate1683, 2013.
- 940 McElman, S., Verma, A., and Goupee, A.: Quantifying Tropical Cyclone-Generated Waves in Extreme-Value-Derived Design for Offshore Wind, Wind Energy Science Discussions, 2024, https://wes.copernicus.org/preprints/wes-2024-129/, 2024.
- Mlawer, E. J., Taubman, S. J., Brown, P. D., Iacono, M. J., and Clough, S. A.: Radiative transfer for inhomogeneous atmospheres: RRTM, a validated correlated-k model for the longwave, Journal of Geophysical Research: Atmospheres, 102, 16663-16682, https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/97JD00237, 1997.
- 945 Mooney, P. A., Gill, D. O., Mulligan, F. J., and Bruyère, C. L.: Hurricane simulation using different representations of atmosphere–ocean interaction: the case of Irene (2011), Atmospheric Science Letters, 17, 415-421, https://rmets.onlinelibrary.wiley.com/doi/10.1002/asl.673, 2016.
- Mooney, P. A., Mulligan, F. J., Bruyère, C. L., Parker, C. L., and Gill, D. O.: Investigating the performance of coupled WRF-ROMS simulations of Hurricane Irene (2011) in a regional climate modeling framework, Atmospheric Research, 215, 57-74, https://www.sciencedirect.com/science/article/pii/S0169809517313091, 2019.
- Morin, G., Boudreault, M., and García-Franco, J. L.: A Global Multi-Source Tropical Cyclone Precipitation (MSTCP) Dataset, Scientific Data, 11, 609, https://www.nature.com/articles/s41597-024-03395-w, 2024.

- Mudd, L. A. and Vickery, P. J.: Gulf of Mexico Hurricane Risk Assessment for Offshore Wind Energy Science Discussions, 2024, https://wes.copernicus.org/preprints/wes-2024-123/, 2024.
- Mukherjee, P. and Ramakrishnan, B.: On the understanding of very severe cyclone storm Ockhi with the WRF-ARW model, Environmental Research: Climate, 1, 015002, https://iopscience.iop.org/article/10.1088/2752-5295/ac6adb, 2022.
- Müller, S., Larsén, X. G., and Verelst, D. R.: Tropical cyclone low-level wind speed, shear, and veer: sensitivity to the boundary layer parametrization in the Weather Research and Forecasting model, Wind Energy Science, 9, 1153-1171, https://wes.copernicus.org/articles/9/1153/2024/, 2024.
- 960 Nakanishi, M. and Niino, H.: An Improved Mellor–Yamada Level-3 Model: Its Numerical Stability and Application to a Regional Prediction of Advection Fog, Boundary-Layer Meteorology, 119, 397-407, https://link.springer.com/article/10.1007/s10546-005-9030-8, 2006.
 - NDBC: Handbook of Automated Data Quality Control Checks and Procedures, 78 https://www.ndbc.noaa.gov/publications/NDBCHandbookofAutomatedDataQualityControl2009handbook.pdf, 2009.
- 965 Olson, J. B., Smirnova, T., Kenyon, J. S., Turner, D. D., Brown, J. M., Zheng, W., and Green, B. W.: A description of the MYNN surface-layer scheme, NOAA, https://repository.library.noaa.gov/view/noaa/30605, 2021.
 - Nie, Y. and Sun, J.: Evaluation of High-Resolution Precipitation Products over Southwest China, Journal of Hydrometeorology, 21, 2691-2712, https://journals.ametsoc.org/view/journals/hydr/21/11/JHM-D-20-0045.1.xml, 2020.
- Oost, W. A., Komen, G. J., Jacobs, C. M. J., and Van Oort, C.: New evidence for a relation between wind stress and wave ages from measurements during ASGAMAGE, Boundary-Layer Meteorology, 103, 409-438, https://link.springer.com/article/10.1023/A:1014913624535, 2002.
- Pan, Y., Yan, C., and Archer, C. L.: Precipitation reduction during Hurricane Harvey with simulated offshore wind farms, Environmental Research Letters, 13, 084007, https://iopscience.iop.org/article/10.1088/1748-9326/aad245, 2018.
- Petrović, V. and Bottasso, C. L.: Wind turbine optimal control during storms, Journal of Physics: Conference Series, 012052, https://iopscience.iop.org/article/10.1088/1742-6596/524/1/012052, 2014.
- Porchetta, S., Muñoz-Esparza, D., Munters, W., van Beeck, J., and van Lipzig, N.: Impact of ocean waves on offshore wind farm power production, Renewable Energy, 180, 1179-1193, https://www.sciencedirect.com/science/article/pii/S0960148121012799, 2021.
- Porchetta, S., Temel, O., Muñoz-Esparza, D., Reuder, J., Monbaliu, J., Van Beeck, J., and van Lipzig, N.: A new roughness length parameterization accounting for wind–wave (mis) alignment, Atmospheric Chemistry and Physics, 19, 6681-6700, https://kuleuven.limo.libis.be/discovery/search?query=any.contains,LIRIAS2807444&tab=LIRIAS&search_scope=lirias_profile&vid=32KUL_KUL:Lirias&offset=0, 2019.
- Porchetta, S., Temel, O., Warner, J. C., Muñoz-Esparza, D., Monbaliu, J., van Beeck, J., and van Lipzig, N.: Evaluation of a roughness length parametrization accounting for wind—wave alignment in a coupled atmosphere—wave model, Quarterly Journal of the Royal Meteorological Society, 147, 825-846, https://rmets.onlinelibrary.wiley.com/doi/10.1002/qj.3948, 2020.
 - Powell, M. D. and Reinhold, T. A.: Tropical Cyclone Destructive Potential by Integrated Kinetic Energy, Bulletin of the American Meteorological Society, 88, 513-526, https://journals.ametsoc.org/view/journals/bams/88/4/bams-88-4-513.xml, 2007.
- Pryor, S. C. and Barthelmie, R. J.: A global assessment of extreme wind speeds for wind energy applications, Nature Energy, 6, 268-276, https://www.nature.com/articles/s41560-020-00773-7, 2021.
 - Pryor, S. C. and Barthelmie, R. J.: Power Production, Inter-and Intra-Array Wake Losses from the US East Coast Offshore Wind shadows impact planning of large offshore wind farms, Applied Energy Lease Areas, Energies, 17, 1063, 359, 122755, https://www.mdpisciencedirect.com/1996-1073/17/5/1063-science/article/pii/S0306261924001387, 2024a.
- Pryor, S. C. and Barthelmie, R. J.: Power Production, Inter-and Intra-Array Wake Losses from the US East Coast Offshore Wind shadows impact planning of large offshore wind farms, Applied Energy, 359, 122755 Lease Areas, Energies, 17, 1063, https://www.sciencedirect.com/science/article/pii/S0306261924001387https://www.mdpi.com/1996-1073/17/5/1063, 2024b.

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Field Code Changed

- Pryor, S. C., Barthelmie, R. J., and Shepherd, T. J.: Wind power production from very large offshore wind farms, Joule, 5, 2663-2686, https://www.sciencedirect.com/science/article/pii/S254243512100430X, 2021.
- and Malkus, J.: Some Aspects of Hurricane Daisy, 1958, Tellus, 13, 181-213, 1000 https://onlinelibrary.wiley.com/doi/10.1111/j.2153-3490.1961.tb00077.x, 1961.
 - Rose, S., Jaramillo, P., Small, M. J., Grossmann, I., and Apt, J.: Hurricane Risk to Offshore Wind Turbines Along the U.S. Coast, Carnegie Mellon University, https://www.cmu.edu/ceic/assets/docs/publications/working-papers/ceic-12-01.pdf, 2012a.
- $Rose, S., Jaramillo, P., Small, M. J., Grossmann, I., and Apt, J.: Quantifying the hurricane risk to offshore wind turbines, Proceedings of the National Academy of Sciences, 109, 3247-3252, \\ $\frac{https://www.pnas.org/doi/full/10.1073/pnas.1111769109?doi=10.1073\%2Fpnas.1111769109}{2012b}.$
 - Schreck III, C. J., Klotzbach, P. J., and Bell, M. M.: Optimal Climate Normals for North Atlantic Hurricane Activity, Geophysical Research Letters, 48, https://agupubs.onlinelibrary.wiley.com/doi/10.1029/2021GL092864, 2021.
- Shchepetkin, A. F. and McWilliams, J. C.: The regional oceanic modeling system (ROMS): a split-explicit, free-surface, 1010 topography-following-coordinate oceanic model. Ocean Modelling, 347-404. https://www.sciencedirect.com/science/article/pii/S1463500304000484, 2005.
- Shchepetkin, A. F. and McWilliams, J. C.: Correction and commentary for "Ocean forecasting in terrain-following coordinates: Formulation and skill assessment of the regional ocean modeling system" by Haidvogel et al., J. Comp. Phys. 227, pp. 3595-Journal of Computational Physics, 228.
- https://www.sciencedirect.com/science/article/pii/S0021999109004872, 2009. 1015
 - Skamarock, W. C.: Evaluating Mesoscale NWP Models Using Kinetic Energy Spectra, Monthly Weather Review, 132, 3019-3032, https://journals.ametsoc.org/view/journals/mwre/132/12/mwr2830.1.xml, 2004.
- Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D. O., Liu, Z., Berner, J., Wang, W., Powers, J. G., Duda, M. G., Barker, D. M., and Huang, X.-Y.: A Description of the Advanced Research WRF Version 4, 162, https://opensky.ucar.edu/islandora/object/technotes%3A576, 2019. 1020
- Stewart, G. M. and Lackner, M. A.: The impact of passive tuned mass dampers and wind-wave misalignment on offshore wind turbine loads, Engineering structures, 73, 54-61, https://www.sciencedirect.com/science/article/pii/S0141029614002673,
- Taylor, P. K. and Yelland, M. J.: The dependence of Sea Surface Roughness on the Height and Steepness of the Waves, Journal 1025 of Physical Oceanography, 31, 572-590, https://journals.ametsoc.org/view/journals/phoc/31/2/1520-<u>0485_2001_031_0572_tdossr_2.0.co_2.xml</u>, 2001.
- Tewari, M., Chen, F., Wang, W., Dudhia, J., LeMone, M. A., Mitchell, K., Ek, M., Gayno, G., Wegiel, J., and Cuenca, R. H.: Implementation and verification of the unified Noah land surface model in the WRF model (Formerly Paper Number 17.5), Proceedings of the 20th Conference on Weather Analysis and Forecasting/16th Conference on Numerical Weather Prediction, Seattle, WA, USA, 6 pp, https://ams.confex.com/ams/84Annual/techprogram/paper 69061.htm, 2004.
- Valamanesh, V., Myers, A. T., and Arwade, S. R.: Multivariate analysis of extreme metocean conditions for offshore wind turbines, Structural Safety, 55, 60-69, https://www.sciencedirect.com/science/article/pii/S0167473015000296, 2015.
- Valamanesh, V., Myers, A. T., Hajjar, J. F., and Arwade, S. R.: Probabilistic Modeling of Joint Hurricane-induced Wind and Wave Hazards to Offshore Wind Farms on the Atlantic Coast, International Conference on Structural Safety and Reliability
- University, 1035 (ICOSSAR). Columbia New York. $\underline{https://people.umass.edu/\!\!\sim\!\!arwade/owt_hurricanes/valamanesh_myers_hajjar.pdf, 2013.$
- Wang, J., Hendricks, E., Rozoff, C. M., Churchfield, M., Zhu, L., Feng, S., Pringle, W. J., Biswas, M., Haupt, S. E., Deskos, G., Jung, C., Xue, P., Berg, L. K., Bryan, G., Kosovic, B., and Kotamarthi, R.: Modeling and observations of North Atlantic cyclones: Implications for US Offshore wind energy, Journal of Renewable and Sustainable Energy, 16, https://pubs.aip.org/aip/jrse/article-abstract/16/5/052702/3316837/Modeling-and-observations-of-North-Atlantic, 2024.

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- Warner, J. C., Perlin, N., and Skyllingstad, E. D.: Using the Model Coupling Toolkit to couple earth system models, Environmental Modelling & Software, 23, 1240-1249, https://www.sciencedirect.com/science/article/pii/S1364815208000388, 2008.
- Warner, J. C., Armstrong, B., He, R., and Zambon, J. B.: Development of a Coupled Ocean–Atmosphere–Wave–Sediment Transport (COAWST) Modeling System, Ocean Modelling, 35, 230-244, https://www.sciencedirect.com/science/article/pii/S1463500310001113, 2010.
- Wiser, R., Rand, J., Seel, J., Beiter, P., Baker, E., Lantz, E., and Gilman, P.: Expert elicitation survey predicts 37% to 49% declines in wind energy costs by 2050, Nature Energy, 6, 555-565, https://www.nature.com/articles/s41560-021-00810-z, 2021.
- 1050 Xie, L., Yan, T., Pietrafesa, L. J., Morrison, J. M., and Karl, T.: Climatology and Interannual Variability of North Atlantic Hurricane Tracks, Journal of Climate, 18, 5370-5381, https://journals.ametsoc.org/view/journals/clim/18/24/jcli3560.1.xml, 2005
 - Yihe, X.: 'Super typhoon' devastates wind farm on Chinese coast, RECHARGE, https://www.rechargenews.com/wind/super-typhoon-devastates-wind-farm-on-chinese-coast/2-1-1706161, 2024.
- Zambon, J. B., He, R., and Warner, J. C.: Investigation of hurricane Ivan using the coupled ocean atmosphere wave sediment transport (COAWST) model, Ocean Dynamics, 64, 1535-1554Tropical to extratropical: Marine environmental changes associated with Superstorm Sandy prior to its landfall, Geophysical Research Letters, 41, 8935-8943, https://link.springeragupubs.onlinelibrary.wiley.com/articledoi/full/10.1007/s10236-014-0777-71002/2014GL061357, 2014a.
- Zambon, J. B., He, R., and Warner, J. C.: Tropical to extratropical: Marine environmental changes associated with Superstorm Sandy prior to its landfall, Geophysical Research Letters, 41, 8935-8943, https://agupubs.onlinelibrary.wiley.com/doi/full/10.1002/2014GL061357, 2014b.
- Zambon, J. B., He, R., and Warner, J. C.: Investigation of hurricane Ivan using the coupled ocean-atmosphere-wave-sediment transport (COAWST) model, Ocean Dynamics, 64, 1535-1554, https://link.springer.com/article/10.1007/s10236-014-0777-7, 2014b.

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